

Watching Volcanic Eruptions From Space:

How we use Satellites to Warn Aviation of the Threat from Ash Clouds

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Himawari-8 I 2018-06-13 03:00 UTC I Left: True colour I Right: True colour with IR ash detection Processed by AIRES I Data courtesy JMA









Sentinel-2 True-colour (RGB) image ~10 m resolution

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* Cloud height and eruption rate





*Volcanoes: distribution, eruptions and hazards

Distribution
Historic Eruptions
Recent Eruptions
Hazards









Move towards ash mass loadings rather than concentrations (maybe this will become "law") Conclusion 7/16 — Definitions of visible ash and discernible ash for operational use. That:

Visible ash be defined as:

"volcanic ash observed by the human eye and not be defined quantitatively by the observer"

Discernible ash be defined as:

"volcanic ash detected by defined impacts on/in aircraft or by agreed in-situ and/or remote-sensing techniques"





Volcanic ash cloud appearance in MODIS True-colour imagery



Brightness Temperature Differences

How do you identify volcanic eruptions in satellite data?



The appearance of a plume that is discoulored and appearing to come from a point source is suggestive. These are fires.

Context is important







This image shows a volcanic plume from Chaiten, southern Chile.

The plume appears "white" – similar to surrounding meteorological cloud but from this "true-colour" MODIS image it is difficult to judge how much of the plume is volcanic





True-colour with BTD overlay



Examples

Using spectral IR data helps

AVHRR visible channel

ASH CLOUD

Mt Ruapehu eruption, 17 June 1996

Context helps but if this eruption occurred at night then visible cues are not available



Brightness temperature difference - the 'reverse' absorption effect



Smoke from forest fires or oil fires or industrial plumes **DO NOT** show the reverse absorption effect



What causes the effect?

The SiO₂ molecule exhibits internal crystal lattice vibrations that preferentially absorb radiation at certain infrared wavelengths. These are the restrahlen and Christiansen frequencies. The effects and locations of the vibrational frequencies are best seen in refractive index measurements.







There is a strong particle size and compositional effect







Hyperspectral Infrared

AIRS and IASI provide 100s of channels in the IR and suggest the possibility of deriving new quantitative aspects of volcanic emissions











Examples (Setting the threshold)



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MODIS True-colour image – Eyjafjallajökull



17 May 2010 13**:**25 UT

Eyja

Mass loadings







Large Eruptions





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Large eruptions

19910817 19910815 19910816 19910818 19910819 19910820 19910821 19910822 19910823 0 20 40 60 80 100 120 140 SO₂ / DU

G. M. Miles et al.: Retrieval of volcanic SO₂ from HIRS/2 using optimal estimation

 Table 3. Total erupted SO2 rounded estimates for Cerro Hudson.

Eruptive phase	TOMS SO_2^1	TOMS SO_2^2	HIRS/2 Prata fit ³	HIRS/2 OE ⁴
8-9 August	700 kT	-	300 kT	$500 \pm 150 \text{kT}$
12 August	600 kT	-	400 kT	$300 \pm 90 \mathrm{kT}$
15 August	2700 kT	2000 kT	1200 kT	$1500 \pm 400 \mathrm{kT}$

¹ Constantine et al. (2000), with errors estimated to be circa 30 %. ² This work, based on updated TOMS algorithm, for total mass as observed on 16 August (as region was poorly observed on the 15th) with consideration of pixel overlap within orbit. ³ After Prata et al. (2003), but data reproduced and sampled as OE HIRS/2 product is herein. ⁴ This work, with retrieved error.

https://www.atmos-meas-tech.net/10/2687/2017/

Hudson, Chile. August 1991 Probably 3rd* largest SO₂ emission in the satellite era

> *Pinatubo ~18Tg El Chichon ~6Tg Hudson ~4Tg Nabro ~4Tg Kasatochi ~2Tg





Large Eruptions

For large eruptions weather satellite data from 1979 onwards can be used to determine SO_2 mass loadings based on a channel at 7.3 µm. This has implications for analysing the effects of volcanic emissions on climate.





Large eruptions

Table 1

Volcanic gases measured	l or	potentially	detectable	from	space.
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		Volatile species									
Sensor ^a	H ₂ 0	C02	C0	50 ₂	R₂S H₂S	HCI	BrO	OCIO	CH ₃ CI	Timespan	Reference(s) ^b
TOMS*										1978-2005	1,2
SBUV* (P)										1978-present	3, 4
HIRS*										1978-present	5
GOME										1995-2003	6, 7, 8, 9
MODIS*										1999-present	10, 11
ASTER										1999-present	12, 13, 14
морит	_									1999-present	15
SCIAMACHY (L)										2002-2012	8, 16, 17, 18
MIPAS (L)										2002–2012	19
AIRS										2002-present	20, 21
ACE (L)										2003-present	22
SEVIRI										2004-present	23
OMI										2004-present	18, 24, 25, 26
MLS ⁺ (L)										1991-2001; 2004-present	27, 28, 29, 30
TES (P)										2004-present	31
GOME-2*										2006-present	18, 32, 33, 34
IASI*										2006-present	15, 35, 36
OMPS*										2011–present	37
VIIRS										2011-present	38
CrIS										2011-present	39
AHI										2015-present	40
GOSAT (P)										2009-present	41
0C0-2										2014-present	42

Red = detected in a volcanic cloud; *Light gray* = potentially detectable but not yet proven in a volcanic context and/or not viable for routine volcanic measurements (e.g., due to background interference).



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Review Multi-decadal satellite measurements of global volcanic

degassing

S.A. Carn $^{a,\ b}$ $\stackrel{o}{\sim}$ $\stackrel{o}{\boxtimes}$, L. Clarisse c , A.J. Prata d

https://www.sciencedirect.com/science/article/pii/S0377027316000032



1979 81 83 85 87 89 91 93 95 97 99 01 03 05 07 09 11 13 15 80 82 84 86 88 90 92 94 96 98 2000 02 04 06 08 10 12 14 Year



Large eruptions - Kelut



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https://www.eumetsat.int/website/home/Images/ImageLibrary/DAT_2169181.html





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How to use satellite imagery: Estimation of cloud height

Some current methods

Radiative Transfer	Cloud-top Temperature + Radiosonde profile	Cloud slicing	Optimal Estimation
Geometrical	Cloud shadow	Stereoscopy	
Direct Measurement	Lidar	Radiosonde	Radar
Model	Plume position and dispersion	Inverse modelling	



Cloud shadow

e.g. Prata and Grant (2001)



Reference: Prata, A. J. and I. F. Grant (2001) Retrieval of microphysical and morphological properties of volcanic ash plumes from satellite data: Application to Mt. Ruapehu, New Zealand., *Quarterly journal of the Royal Meteorological Society*, 127 (576B), 2153-2179.



Geometry: Parallax from two satellites

e.g. Merucci, L., Klemen Zakšek, K., Carboni, E. and S. Corradini (2016)



Reference: Merucci et al. (2016), Stereoscopic Estimation of Volcanic Ash Cloud-Top Height from Two Geostationary Satellites, Remote Sens., 8, 206; doi:10.3390/rs8030206



Other methods

Inversion:

Stohl, A., Prata, A.J., Eckhardt, S., Clarisse, L., Durant, A., Henne, S., Kristiansen, N.I., Minikin, A., Schumann, U., Seibert, P. et al. **(2011).** Determination of time- and height-resolved volcanic ash emissions and their use for quantitative ash dispersion modeling: The 2010 Eyjafjallajökull eruption. *Atmos. Chem. Phys.*, 11, 4333–4351.

Optimal Estimation:

Francis, P.N., Cooke, M.C. and Saunders, R.W., **(2012).** Retrieval of physical properties of volcanic ash using Meteosat: A case study from the 2010 Eyjafjallajökull eruption. *Journal of Geophysical Research: Atmospheres*, *117*(D20).

Cloud-slicing:

Holz, R.E., Ackerman, S., Antonelli, P., Nagle, F., Knuteson, R.O., McGill, M., Hlavka, D.L. and Hart, W.D., **(2006).** An improvement to the high-spectral-resolution CO₂-slicing cloud-top altitude retrieval. *Journal of Atmospheric and Oceanic Technology*, *23*(5), pp.653-670.

MISR Stereo:

Flower, V.J. and Kahn, R.A., **(2017).** Assessing the altitude and dispersion of volcanic plumes using MISR multi-angle imaging from space: Sixteen years of volcanic activity in the Kamchatka Peninsula, Russia. *Journal of Volcanology and Geothermal Research*, 337, pp.1-15.

Cloud-top/radiosonde:

Woods, A.W. and Self, S., (1992). Thermal disequilibrium at the top of volcanic clouds and its effect on estimates of the column height. *Nature*, *355*(6361), p.628.



Stereoscopy from ATSR/ATSR-2/AATSR/SLSTR

e.g. Prata and Turner (1997)







Lascar eruption







Nadir view

Forward view



Example: Eyjafjallajökull

8-May 2010 Highly dispersed ash cloud over ocean (>36 hrs old)



SEVIRI (actual measurements)











Stare at the image until a third stereo image appears between the pair (you may need to go "cross-eyed")



There is a thin veil of high cloud running across the top-left corner of the image The aircraft contrails are below this The ash cloud (dark) is below both



Height retrieval



Method

Use 1.6 μ m reflectance nadir and forward views. Only process pixels where BTD_n<0. The r² correlation between parts of the image is calculated in 8x8 pixel chunks. Some smoothing is applied for graphical representation.



Validation-Caliop/AIRS











Insights

Contrails show that aircraft were crossing the Atlantic while ash was in the vicinity



London VAAC advisory suggests ash up to FL350 (35,000 ft ~10.7 km)



NATS 29,000-41,000 ft (~8.8-12.5 km)



VA ADVISORY DTG: 20100508/1200Z VAAC: LONDON VOLCANO: EVJAFJALLAJOKULL 1702-02 PSN: N6338 W01937 42 AREA: ICELAND

SUMMIT ELEV: 1666M ADVISORY NR: 2010/092 INFO SOURCE: ICELAND MET OFFICE AVIATION COLOUR CODE: RED ERUPTION DETAILS: PLUME ERUPTION CONTINUES, HEIGHTS UP TO FL230.

RMK: ADVISORY NO. NOW UPDATED AND CORRECTED NXT ADVISORY: 20100508/1800Z



Prior art (using satellite data)

- Cloud shape recognition pioneering work of Sawada (motivation to identify volcanic clouds in single channel imagery) – 1980's
- Solène Pouget (and co-workers) 2010 onwards
- Convective clouds much work in atmospheric sciences based on rapid growth and precipitation onset



Geostationary Satellites

Geo-satellites ideal because of high data repetition







Example of rapid vertical (convective) development of meteorological cloud























































Can the rapid spread of the umbrella cloud be used to estimate the eruption rate of the vertical column?



After Agusto Neri



Satellites detect many eruption columns from early onset













Evolution of the 11 µm brightness temperature at 10 min intervals







The volcanic cloud is easily observed in Himawari-8 and other satellite data. Himawari-8 samples the region at 2 x 2 km² spatial resolution in the infrared every 10 minutes. Assuming that the eruption column rises at some vertical ascent rate, unaffected by cross-winds, reaches a neutral buoyancy level and then spreads horizontally, an estimate of the rate of radial spreading can be used to estimate the eruption rate at the top of the cloud, essentially as a consequence of conservation of mass. A common formulation for the increase of the radius of the umbrella cloud is (Sparks et al., 1997; Pouget et al., 2013):

$$r_t = \left(\frac{3\nu NQ_t}{2\pi}\right)^{1/3} t^{2/3}$$

r is the radius

t is time

Q is the volumetric eruption rate N is the Brunt-Vaisala frequency ν is a dimensionless constant











