

Occlusion – cyclogenesis II

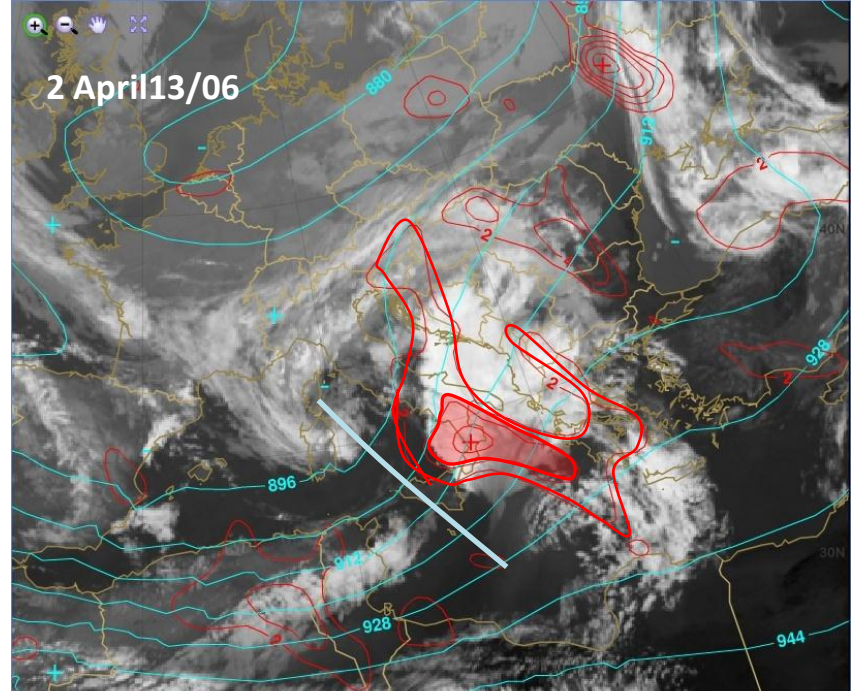
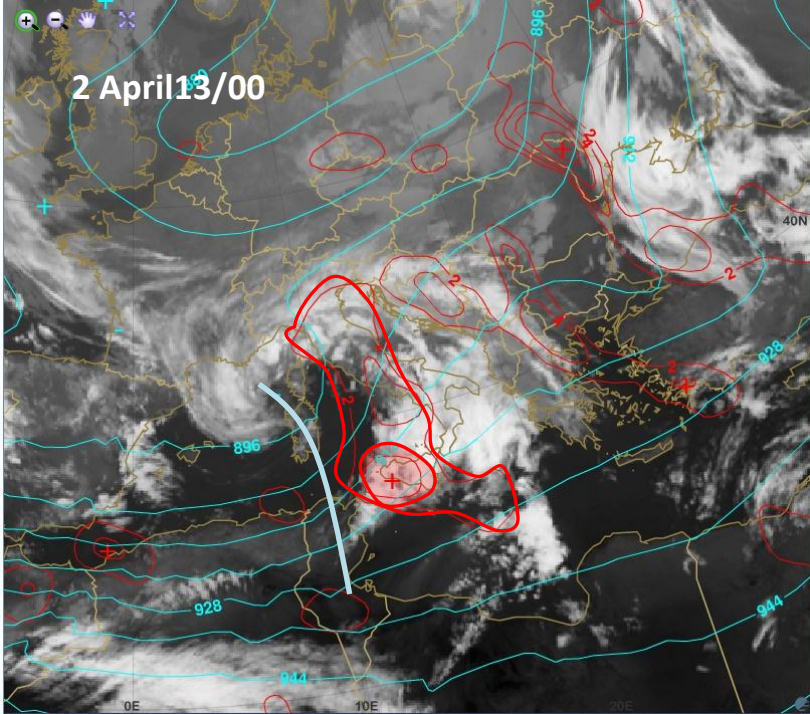
Upper level influences – Hoskins theory

Rapid cyclogenesis

Other types of cyclogenesis:

Instant occlusion

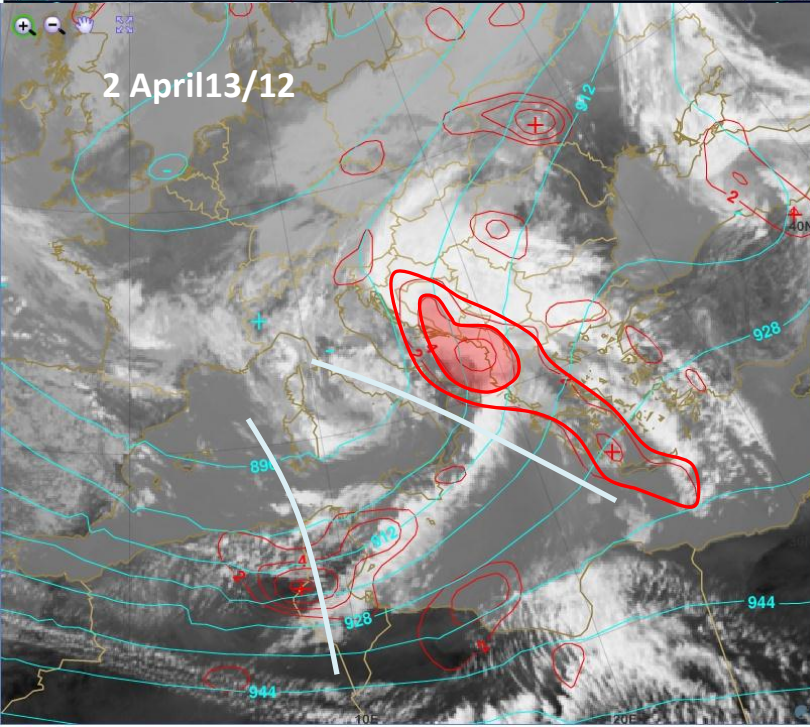
Cold air Development



WCB type:

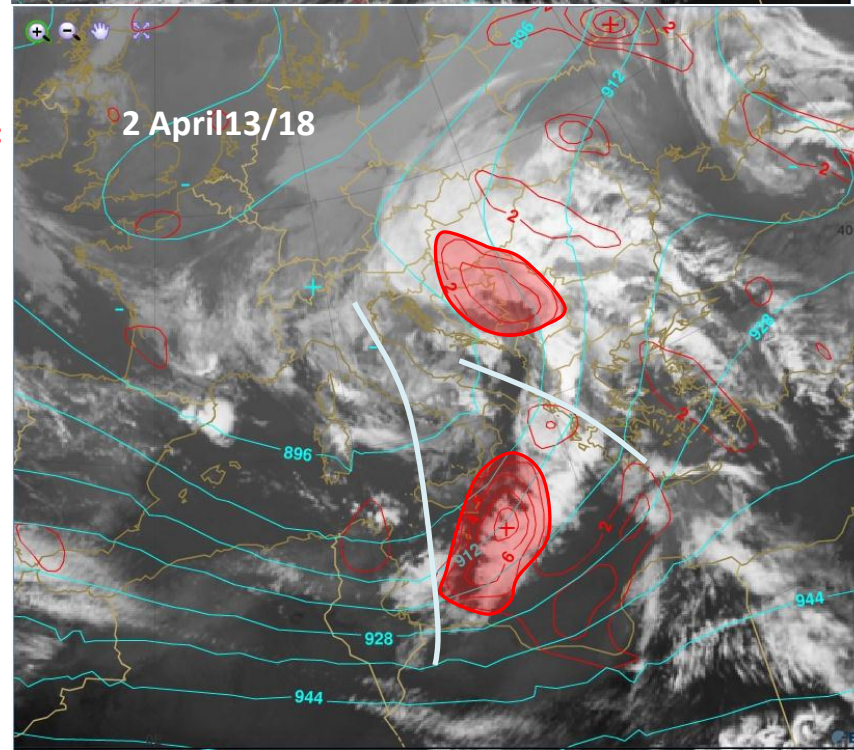
Upper level
Trough
300 hPa
+

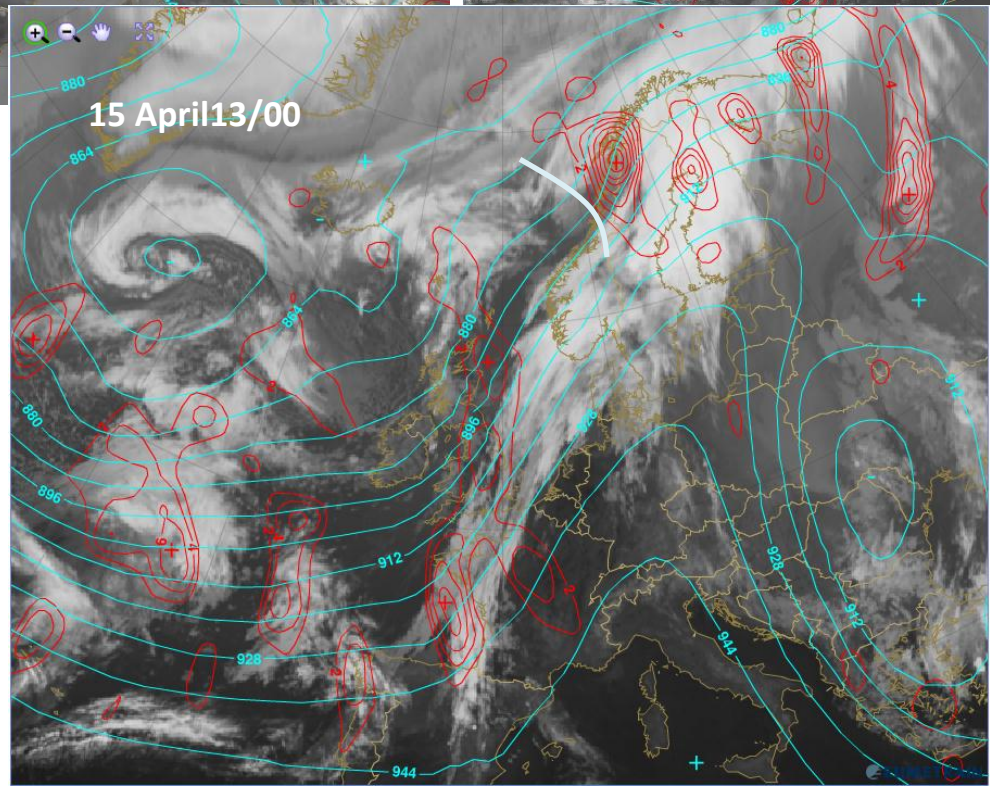
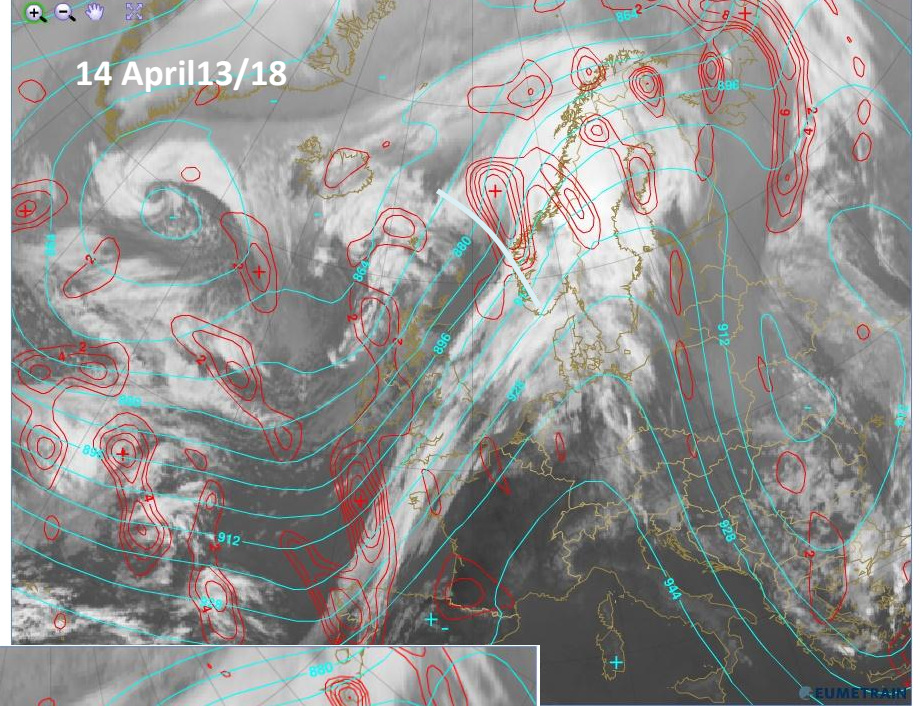
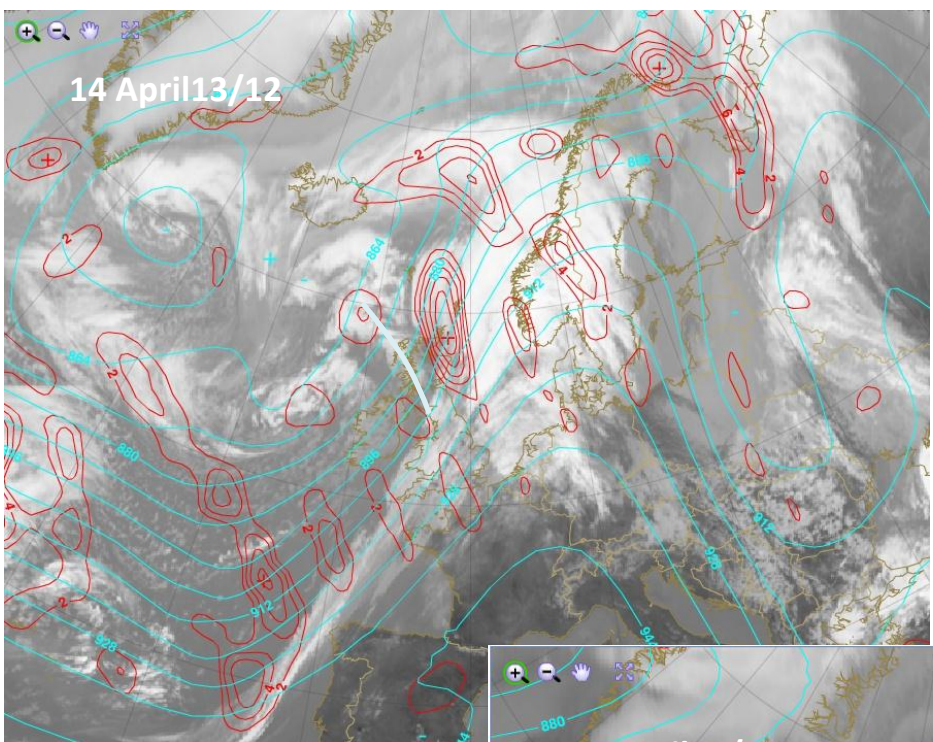
PVA 300hPa



Distinct
PVA maxima:

Why at this
Position?





CCB type:

Upper level
Trough 300 hPa
+
PVA 300 hPa

Distinct PVA max:
why so
intensive

Reflectance of occlusion process / cyclogenesis in derived upper level parameters

- For Fronts Temperature Advection is a dominating parameters
- For cyclogenesis also Vorticity and Vorticity Advection are dominating parameter

Vorticity

- $\zeta = \text{rot}_z \mathbf{v}_2 = \partial v / \partial x - \partial u / \partial y$
- $\eta = \zeta + f$ $f = 2 \Omega \sin \varphi$

- cyclonic: $\zeta > 0$ (NH) $\zeta < 0$ (SH)
- anticyclonic: $\zeta < 0$ (NH), $\zeta > 0$ (SH)

- Vorticity is a quality of the stream field
- Vorticity describes rotation movements

Vorticity is one of the parameters which are responsible for the formation of cloudiness

Vorticity Advection

- $VA = -\mathbf{v} \cdot \nabla \zeta$ Advection of rel. vorticity
- $VA = -\mathbf{v} \cdot \nabla \eta$ Advection of abs. vorticity

- $VA > 0$ positive Vorticity advection: PVA
- $VA < 0$ negative Vorticity advection: NVA

- PVA: it becomes more cyclonic (NH) or more anticyclonic (SH)
- NVA: it becomes more anticyclonic (NH) or more cyclonic (SH)

The Advection of Vorticity is one of the parameters which is responsible for the development of cloudiness

Omega equation in a quasi geostrophic system

$$\left(\sigma \nabla_p^2 + f_0^2 \frac{\partial^2}{\partial p^2} \right) \omega = -f_0 \frac{\partial}{\partial p} \left[-V_g \cdot \nabla_p (\zeta_g + f) \right] - \frac{R}{p} \nabla^2 \left[-V_g \cdot \nabla_p T \right] - \frac{R}{p c_p} \nabla^2 H$$

$\approx -\omega$ 1 2 3
Vorticity-
advection Temperature-
advection Diabatic
heating
1.) **2.)** **3.)**

c_p specific heat at constant pressure
 f Coriolis parameter
 p pressure
 H diabatic heat transfer
 R universal gas constant
 T temperature
 V_g geostrophic wind vector
 ζ_g relative geostrophic vorticity
 σ static stability
 ω vertical motion

- Assumptions for a Quasi-geostrophic Systems:**
 - 1.) there is hydrostatic equilibrium in the vertical
 - 2.) Pressure- and vorticity fields are in a geostrophic equilibrium
- The right side in this representation of the **omega equation** consists of three terms:
 - 1.) the vertical change of the vorticity advection**
 - 2.) the Laplacian of the temperature advection (the change of the horizontal gradient of the temperature advection)
 - 3.) the Laplacian of diabatic heat transfer (the change of the gradient of diabatic heat transfer)

Some basic considerations for vorticity and vorticity advection

- the vertical change of the vorticity advection

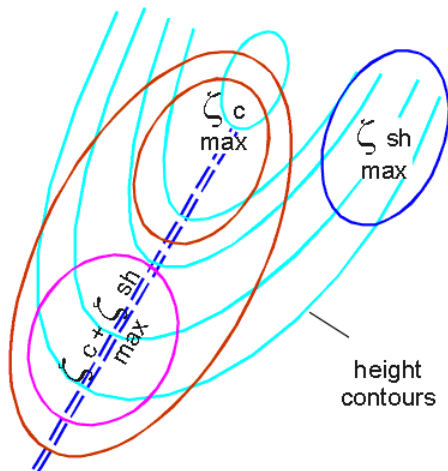
$$\partial/\partial p \cdot (-\mathbf{v}_p \cdot \nabla \eta)$$

- Upward motion is connected with
 - The increase of PVA (CVA) with height
 - The decrease of NVA (AVA) with height
- In the method of analysing satellite images together with numerical model a simplification is applied:
- Instead of the vertical change of Vorticity advection only the pronounced maxima of Vorticity advection in upper levels (500, 300 hPa) are taken into account
- This needs the assumption of:
 - Increase of wind speed with height
 - This is reasonable but every time fulfilled

The significance of processes in the upper troposphere for cyclogenesis

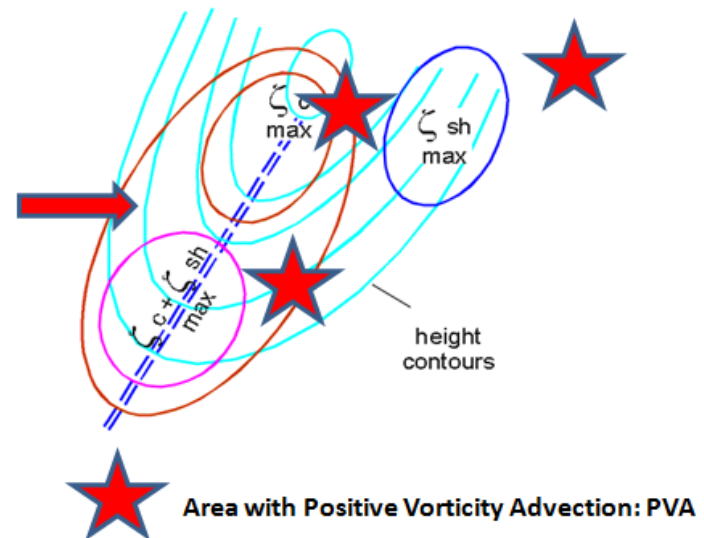
Vorticity (rotation)

- Vorticity consists of two contributions: curvature and shear vorticity
- Typical areas for Positive (Cyclonic) Vorticity values are (on the NH):
 - lows, troughs and jet streaks



Vorticity Advection

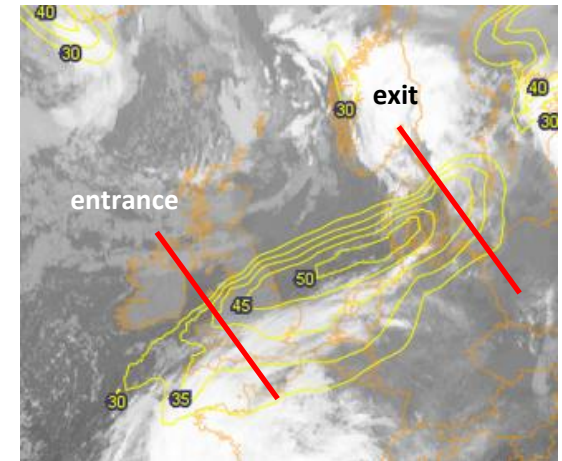
- If a weather system with high cyclonic values of vorticity is moving:
 - Downstream from Vorticity maxima becomes the troposphere more cyclonic →
 - Areas of Positive Vorticity Advection (PVA)



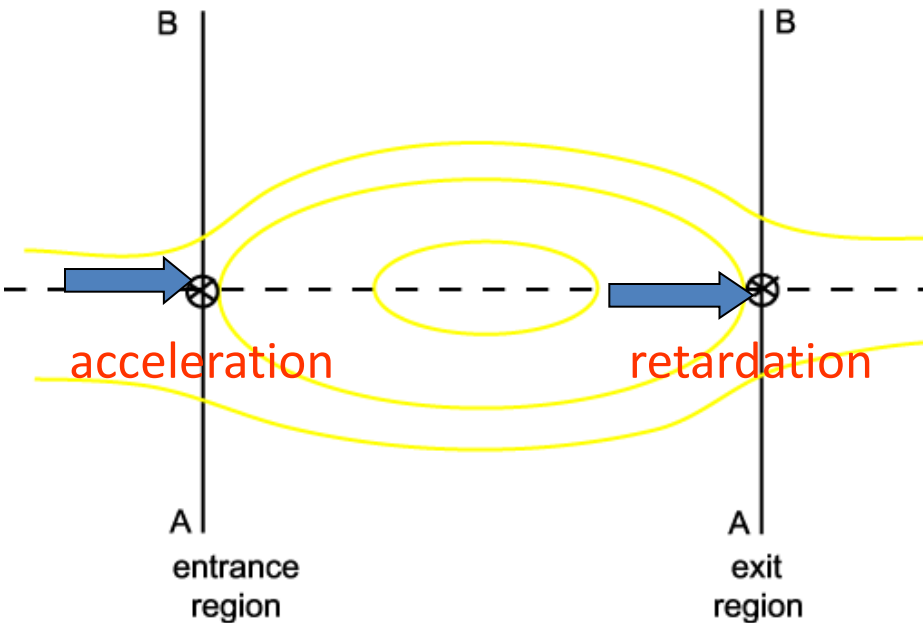
Jet streaks – 4-quadrant Model (Uccellini)

Consideration 1:

Jet streaks: Stream components in jet level



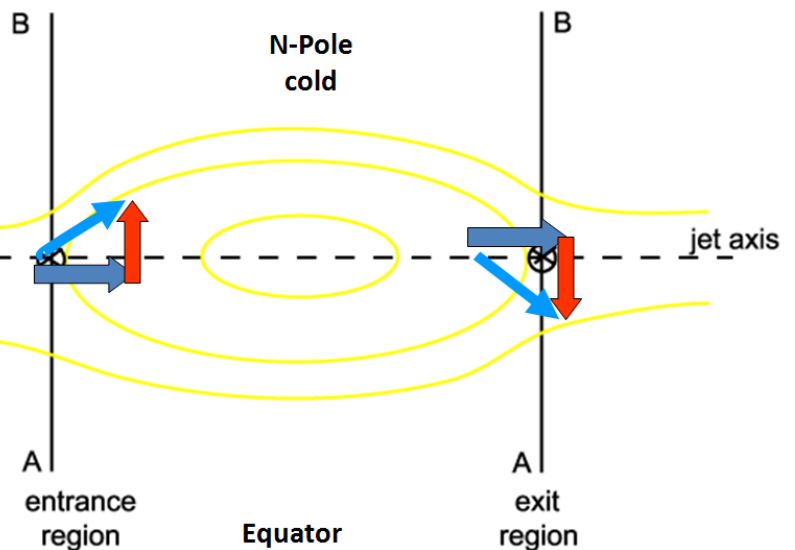
jet streak



Acceleration vector

Ageostrophic Wind vector

jet streak

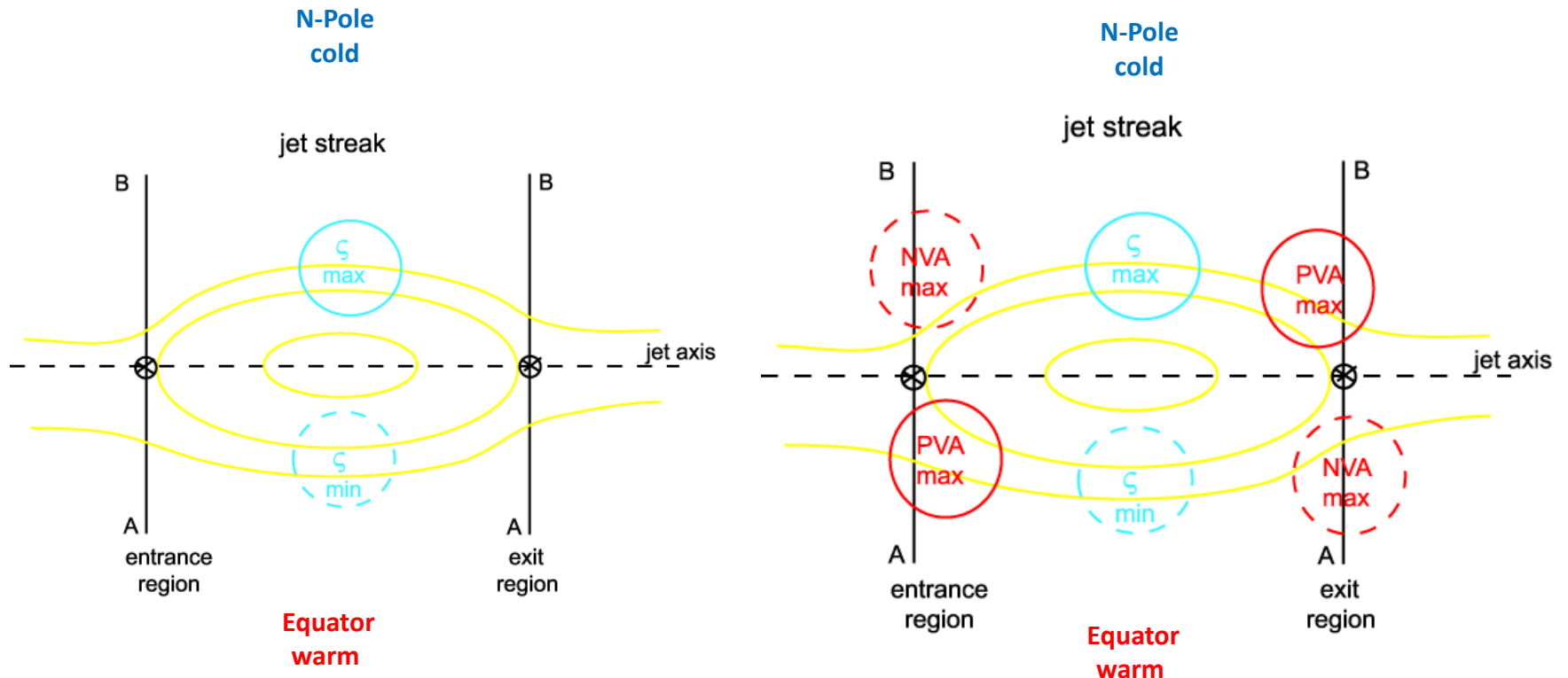


Resulting vector

Jet streaks – 4-quadrant Model (Uccellini)

Consideration 2:

Jet streaks: Distribution of vorticity and vorticity advection in jet level

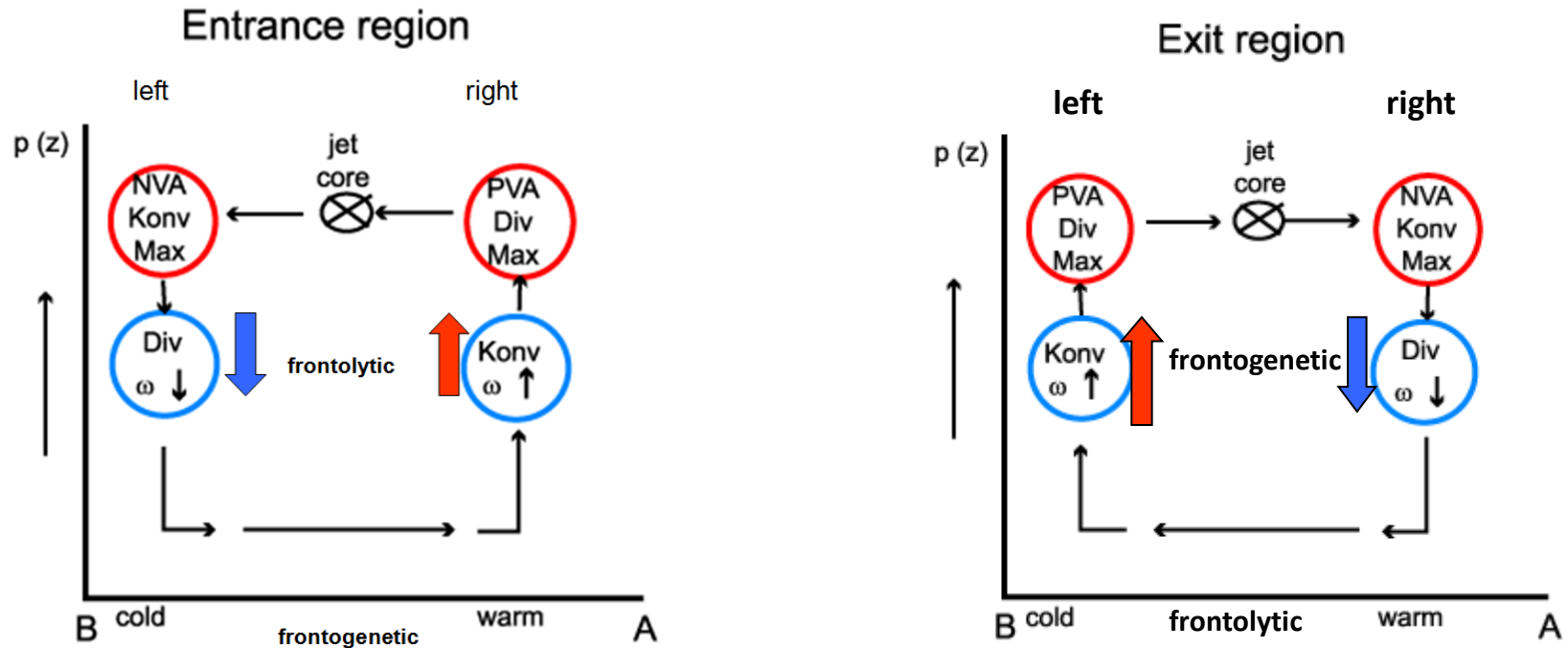


- Vorticity maximum on the cyclonic side of a jet streak maximum (increased shear)
- Vorticity minimum on the anticyclonic side

Jet streaks – 4-quadrant Model (Uccellini)

Consideration 3: Jet streaks:

Vertical cross section in entrance and exit region

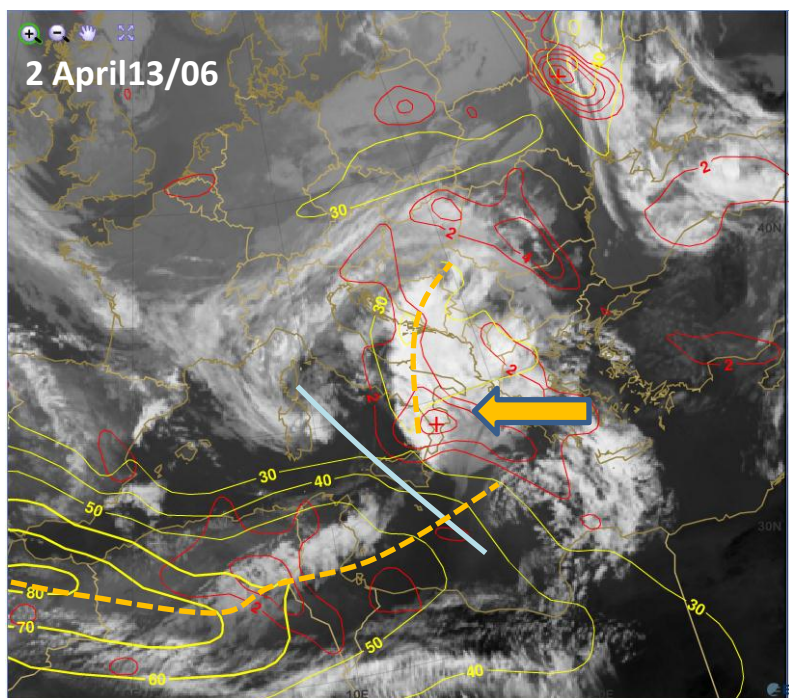
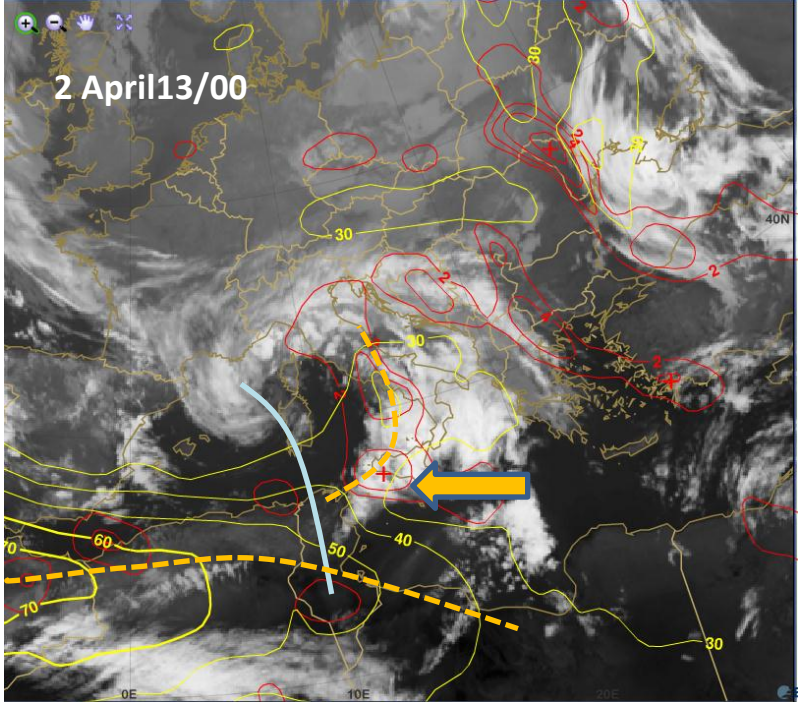


Entrance Region:

1. Frontogenetic in low levels
2. Frontolytic in middle levels
3. Upward motion in right entrance region
4. Preferred area for wave developments

Exit Region:

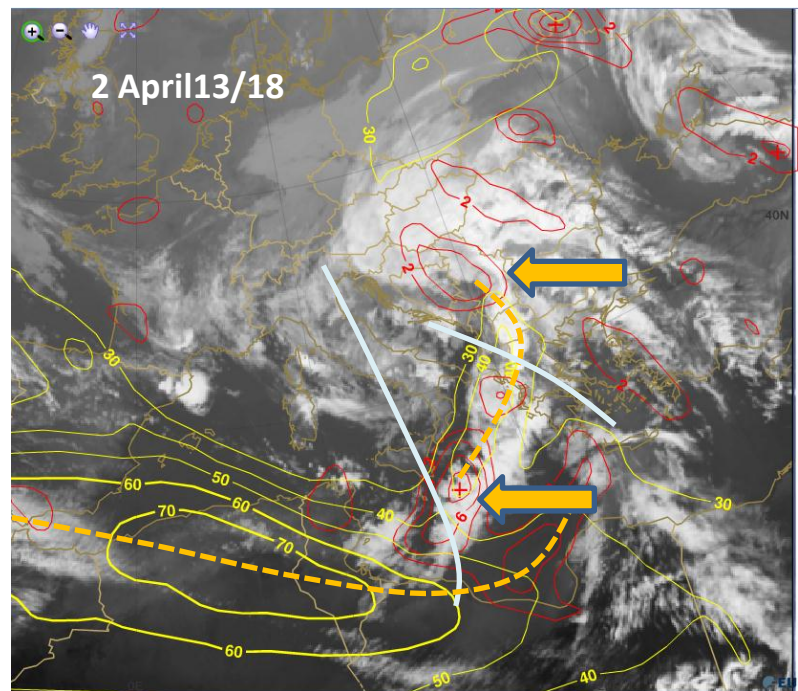
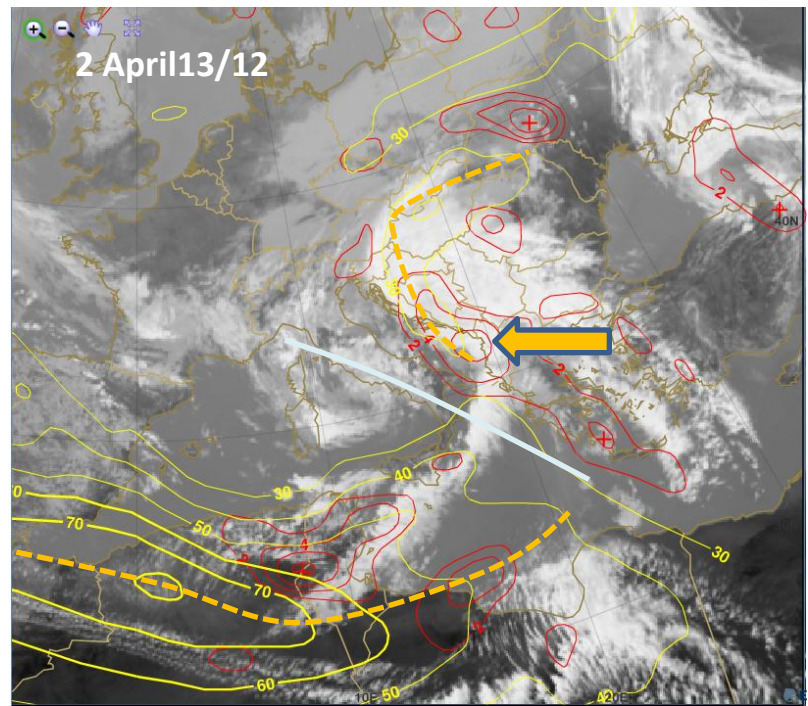
1. Frontolytic in low levels
2. Frontogenetic in middle levels
3. Upward motion in left exit region
4. Preferred area for rapid developments

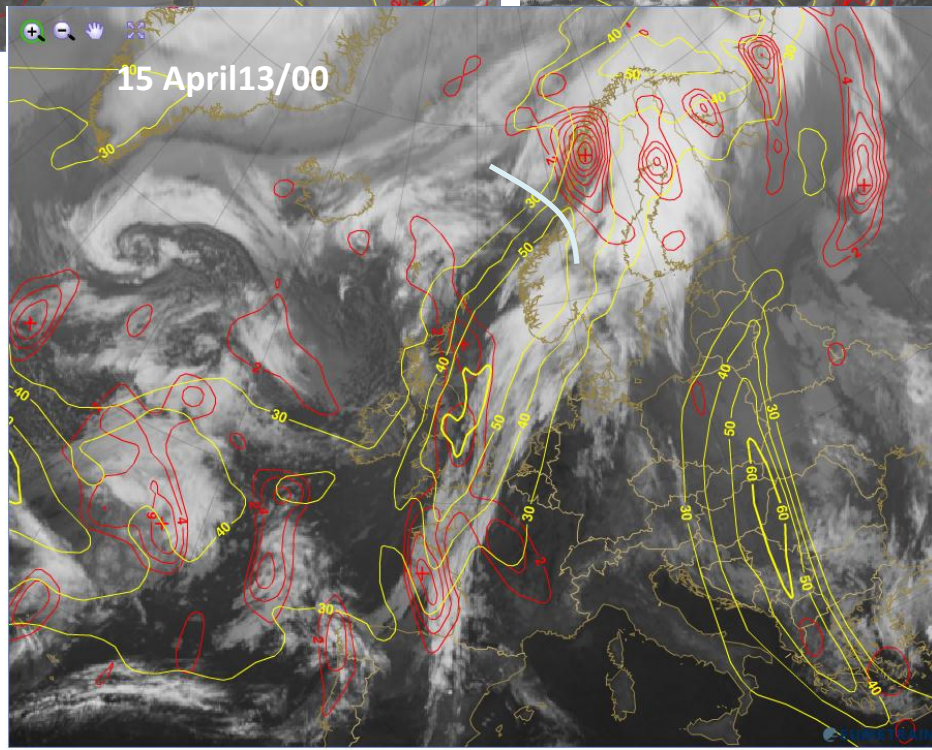
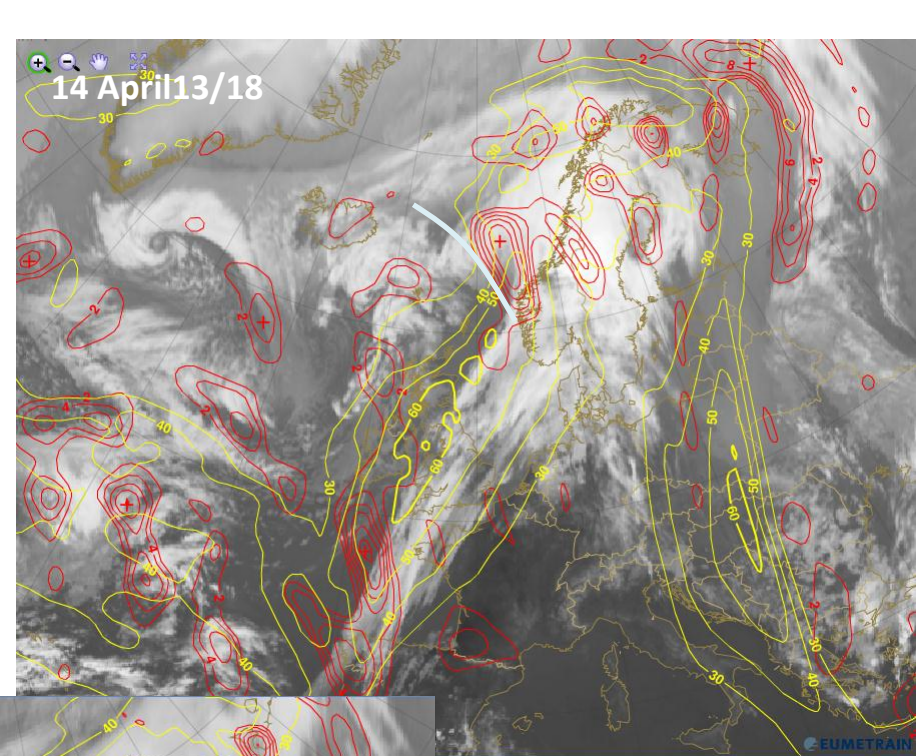
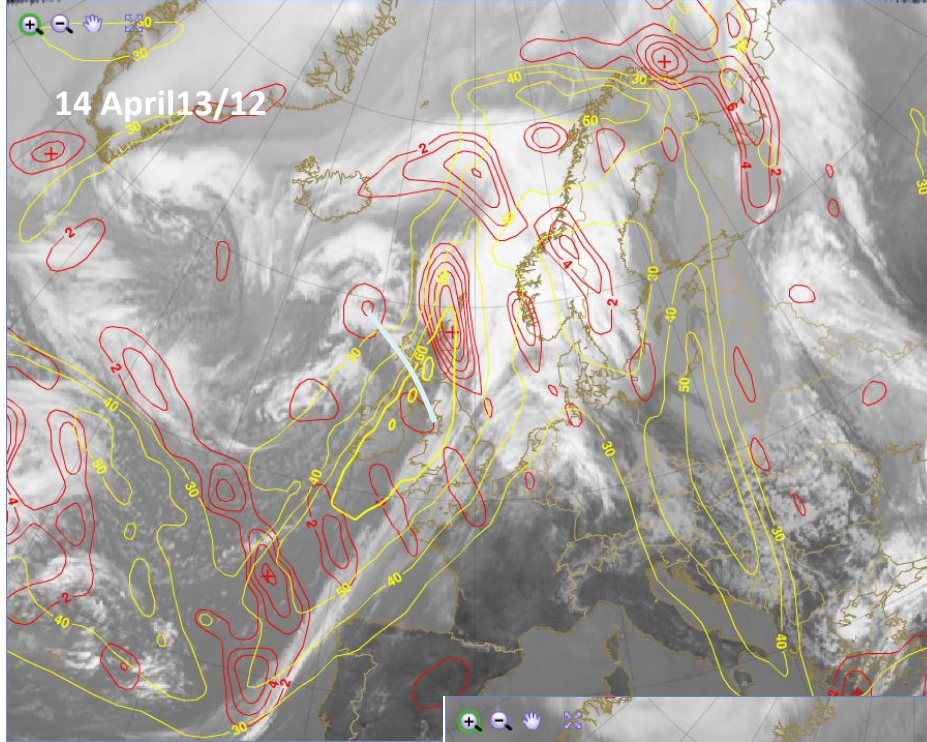


Wcb type:

Jet streak
+
PVA

Distinct
PVA maxima:
in REN + LEX
of two
involved jets



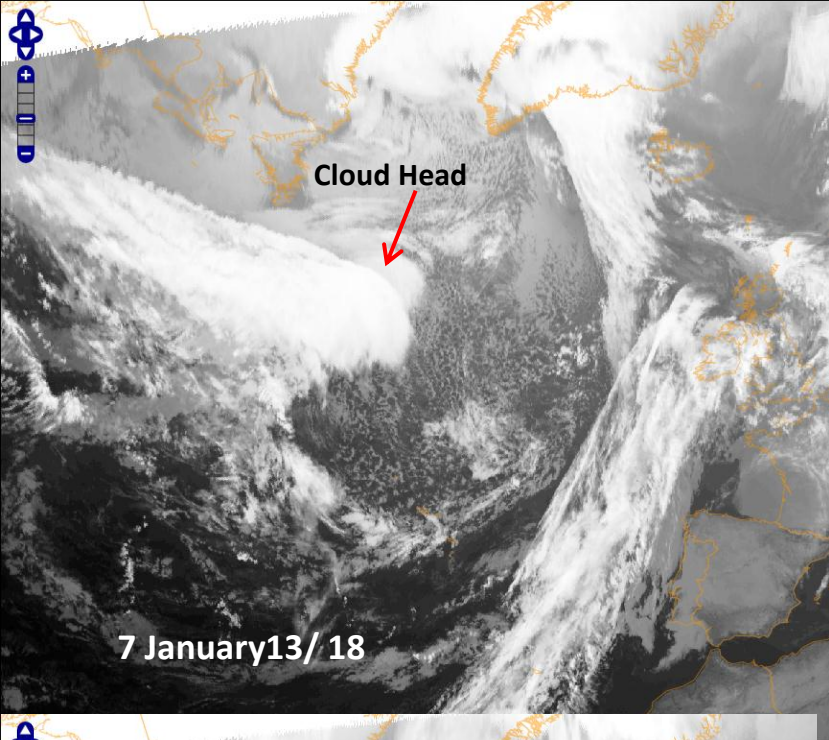


ccb type:

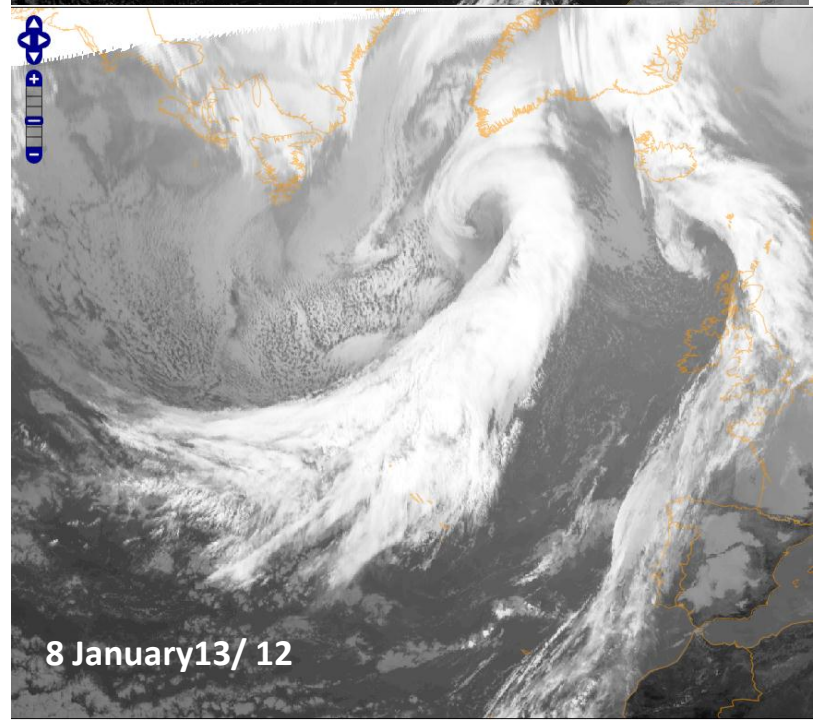
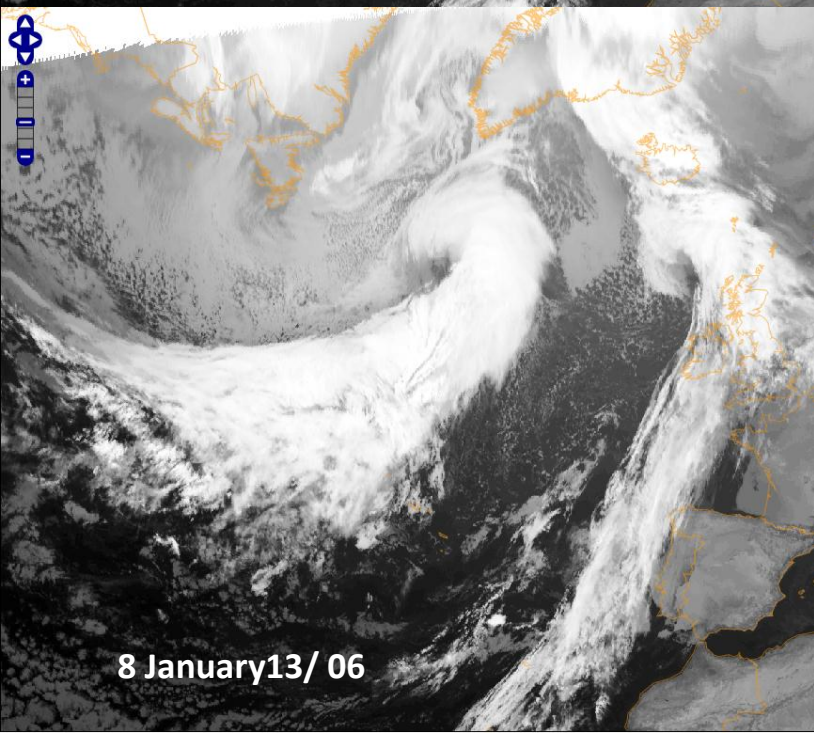
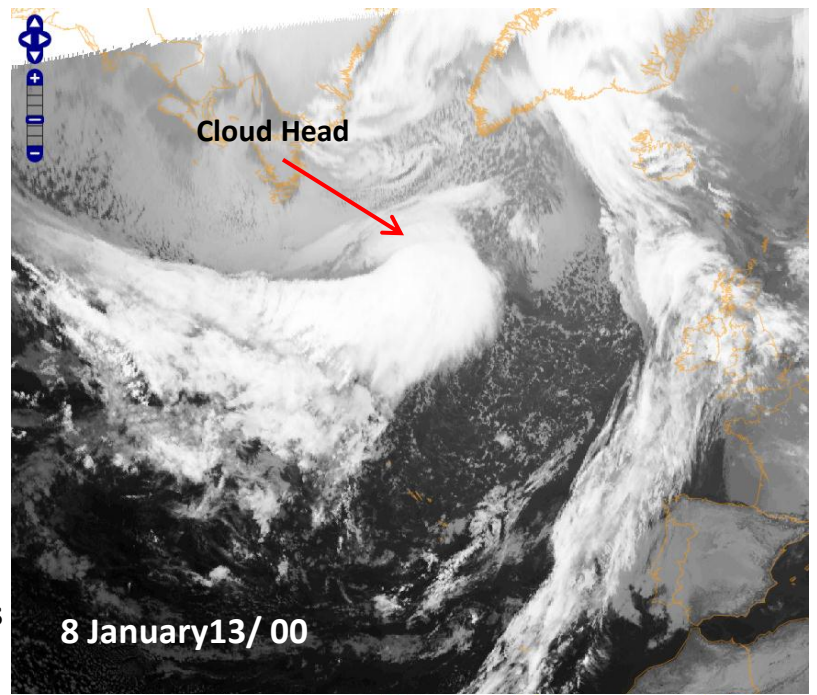
Upper level
Trough
+
PVA in LEX of
a jet streak

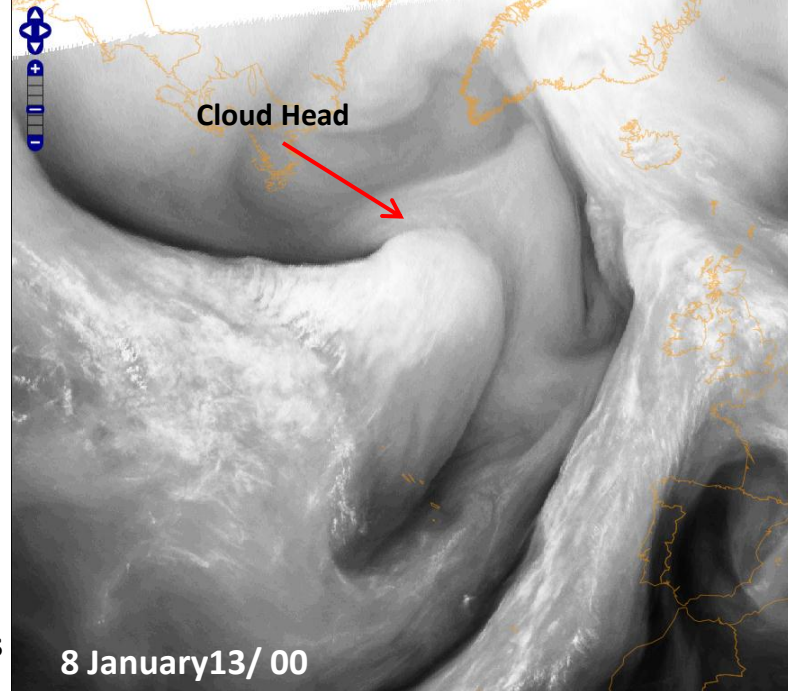
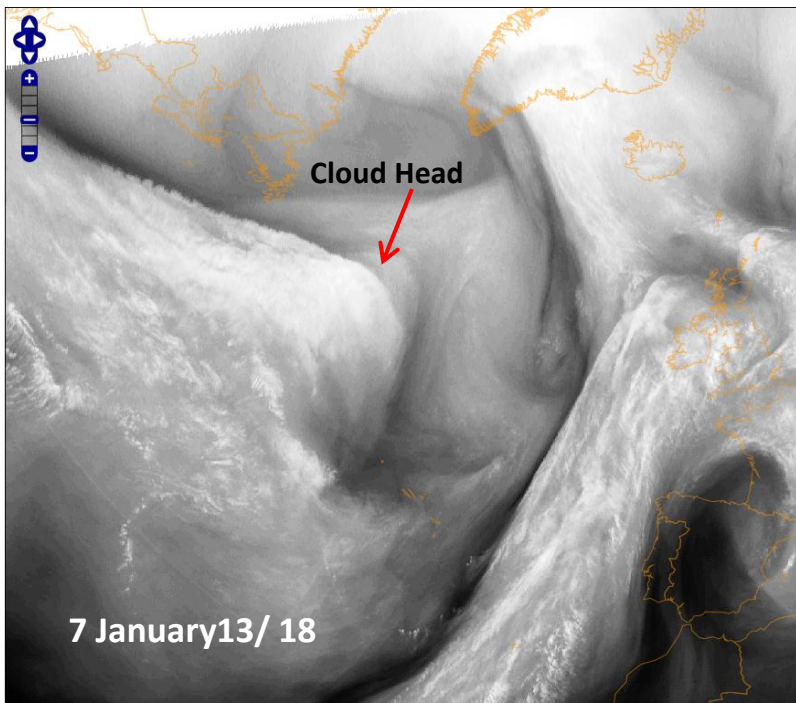
Rapid Cyclogenesis

- Intensive deepening of surface low within 12 hours
- Very typical structure in satellite images which show involvement of processes in upper levels

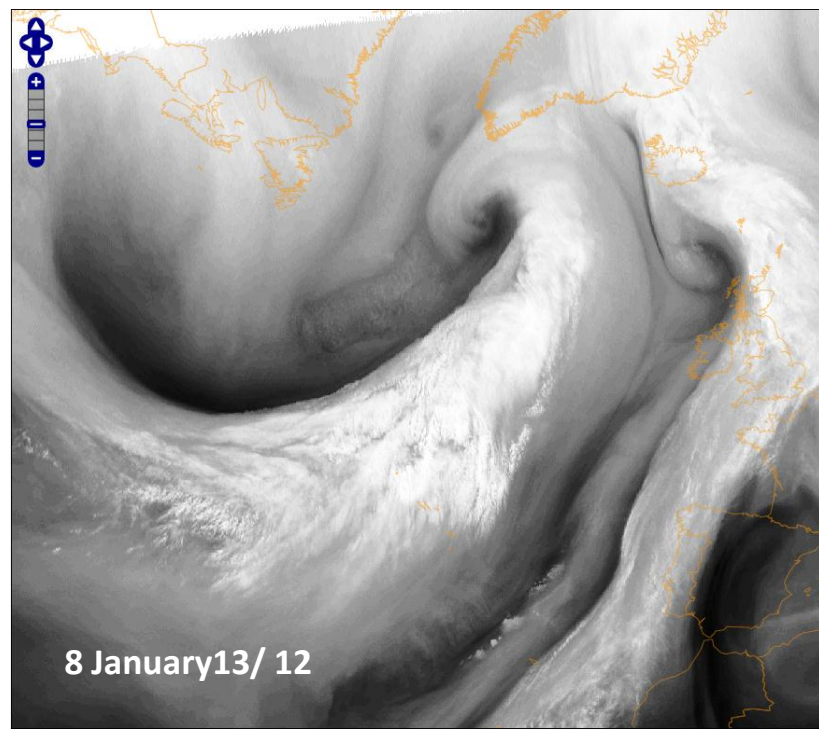
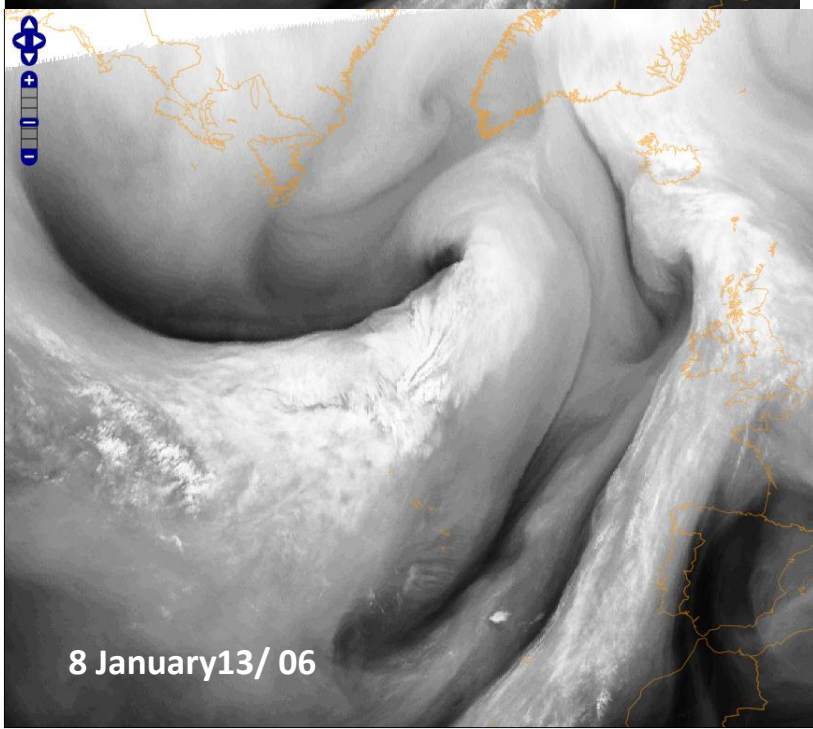


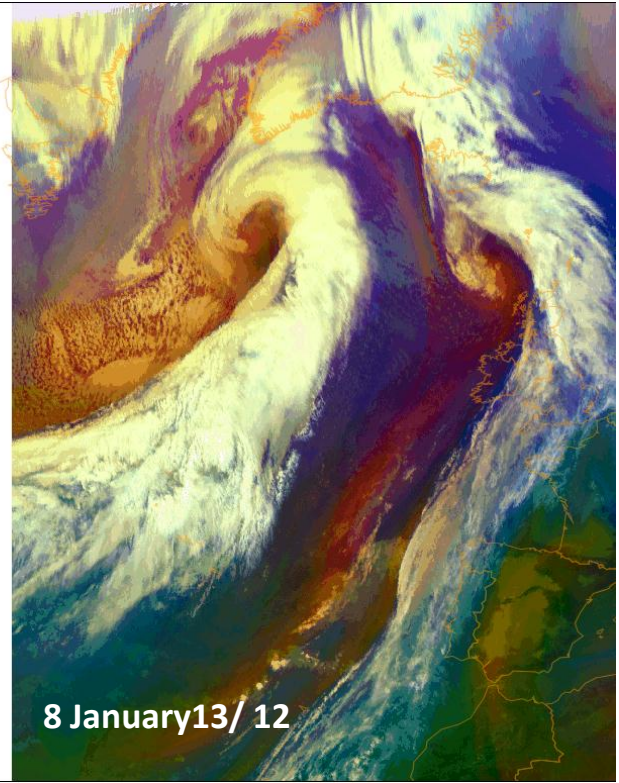
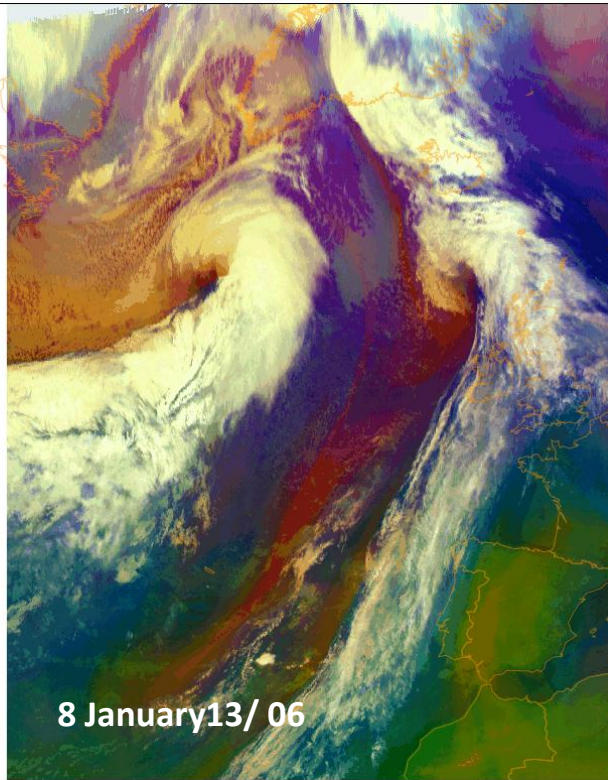
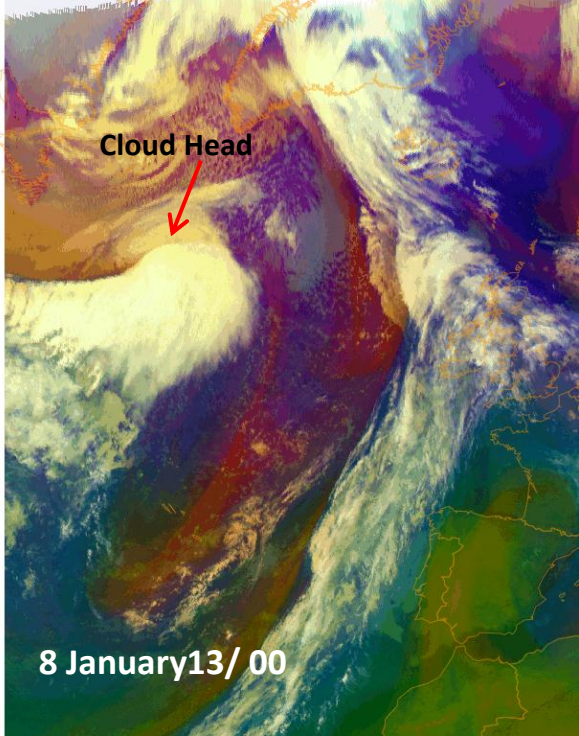
7 -8
January
2013
Rapid
Cyclogenesis
IR



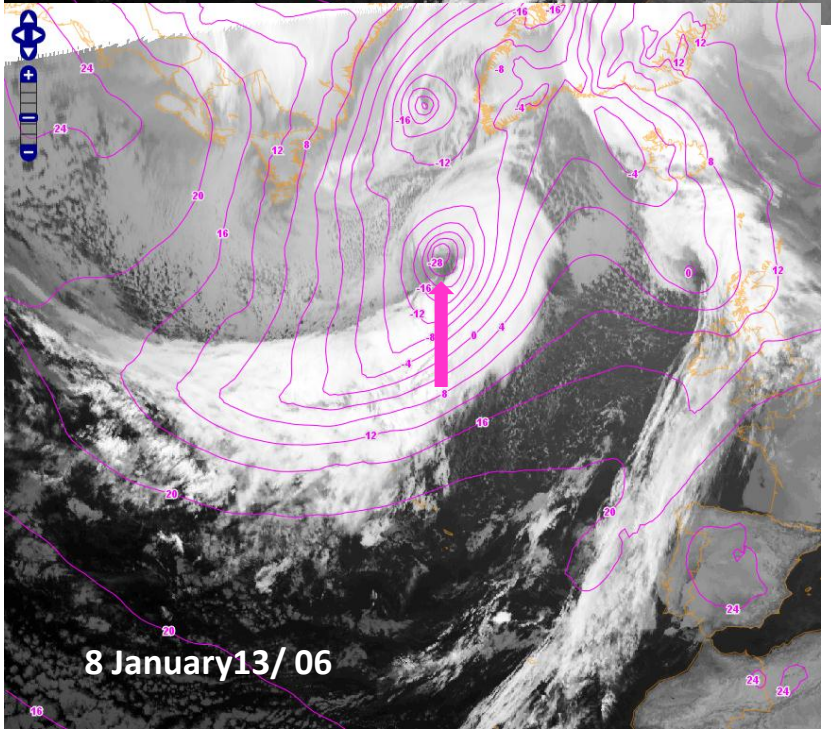
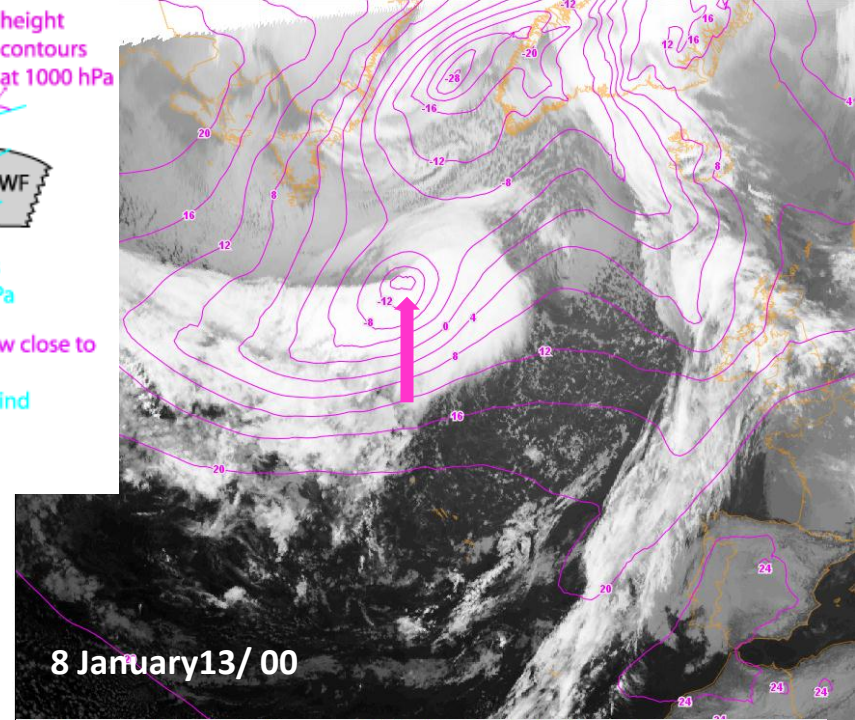
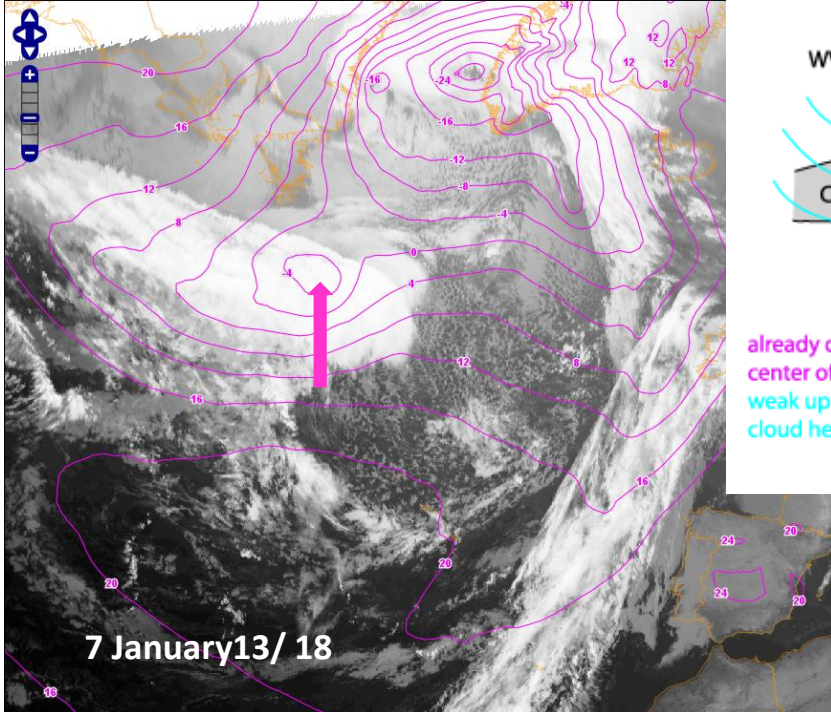


7 -8
January
2013
Rapid
Cyclogenesis
WV

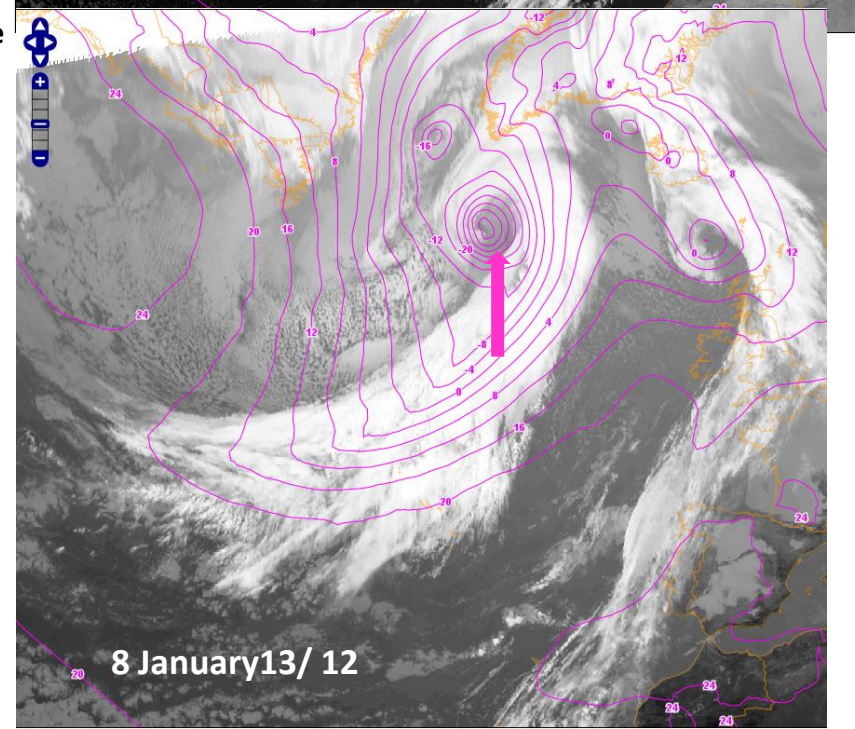


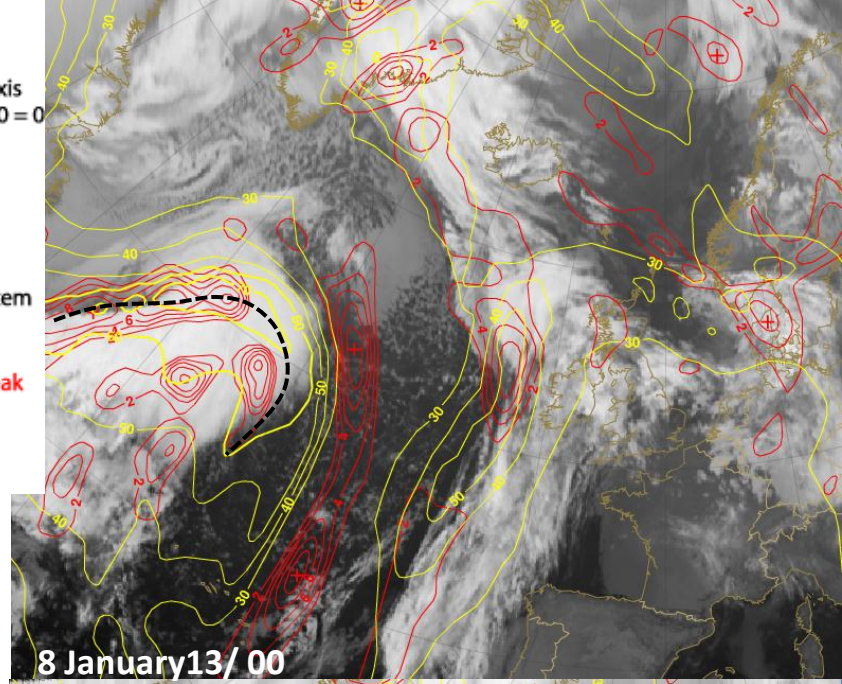
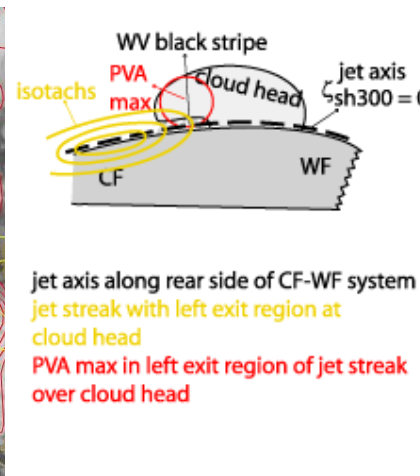
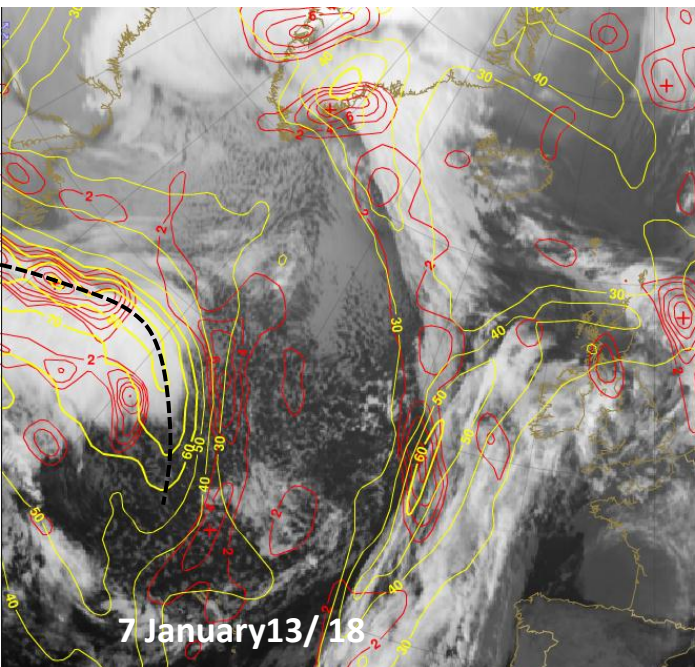


7 -8
January
2013
Rapid
Cyclogenesis
Air Mass RGB



Deepening of height at 1000 hPa (close surface Low) from -4 gpm to -36 gpm within 24 hours

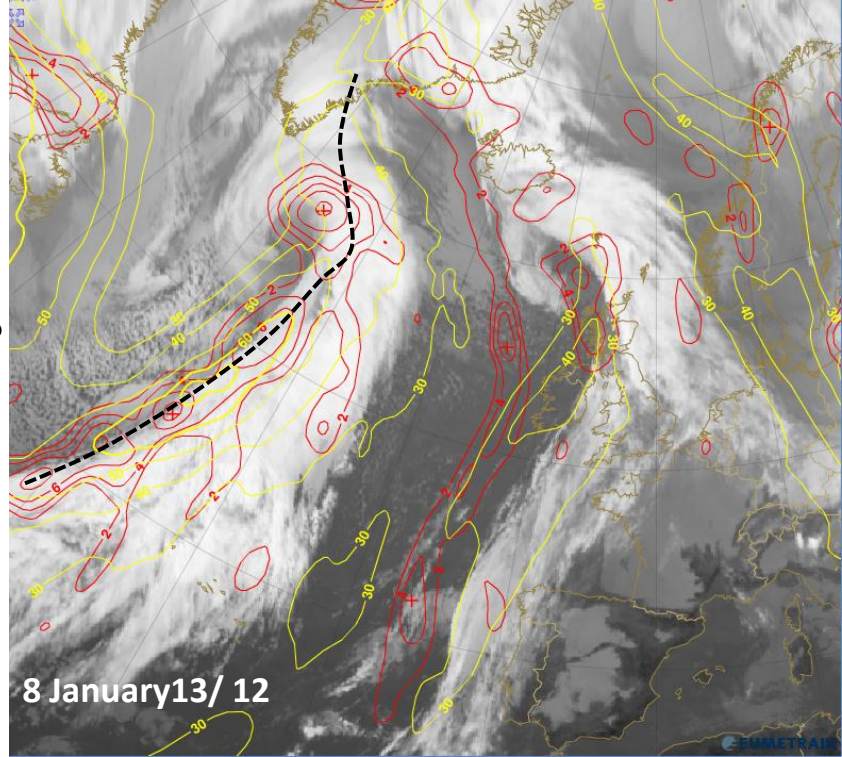
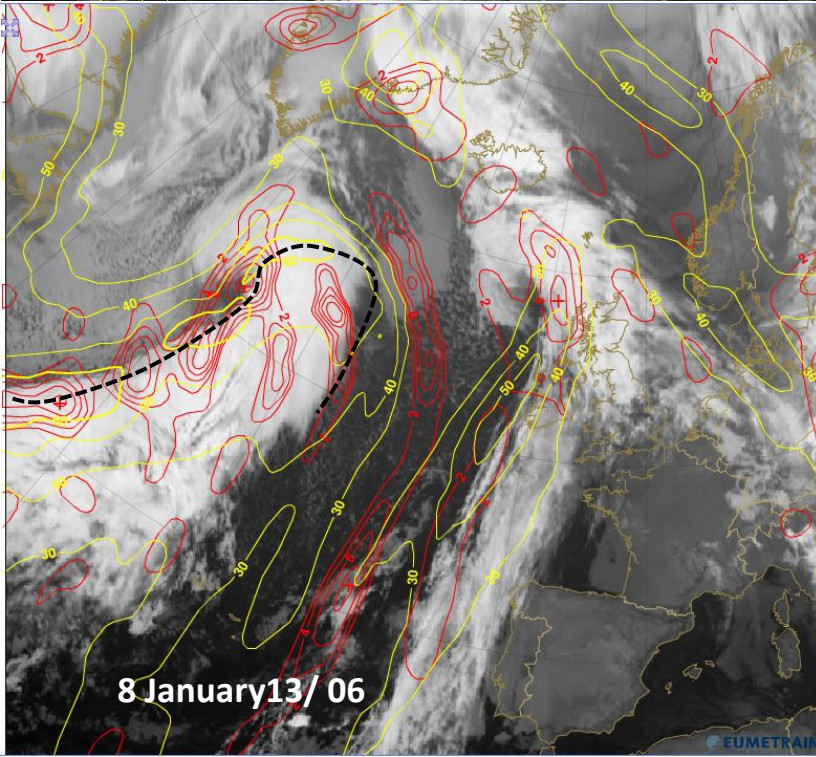




**Rapid
Cyclogenesis**

**PVA max
In LEX
Of intensive
jet streak**

**Why such
A rapid
cyclogenesis!?**



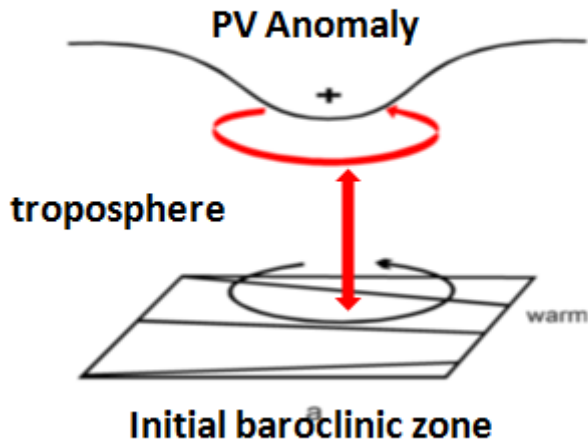
The significance of processes in the upper troposphere for cyclogenesis: Potential Vorticity (PV)

- $PV = -g (f + \zeta_{\theta}) \partial\theta / \partial p = \text{const}$
- PV (IPV) is constant for adiabatic processes; that means constant on isentropic surfaces
- Product of vorticity and stability is constant on isentropic surfaces \Rightarrow
- PV has very high values in stratosphere ≥ 2 units (NH); ≤ -2 units (SH)
- If stratospheric air with the inherent PV protrudes downward from stratosphere into troposphere - static stability decreases and (because of conservation) cyclonic vorticity has to increase there
- This is a motor for enhancing cyclogenesis

The significance of processes in the upper troposphere for cyclogenesis: Potential Vorticity (PV)

PV Anomalies - Hoskins theory

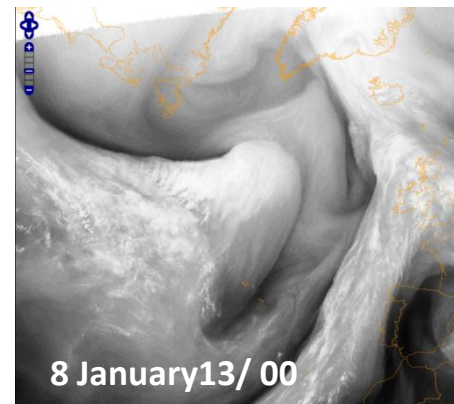
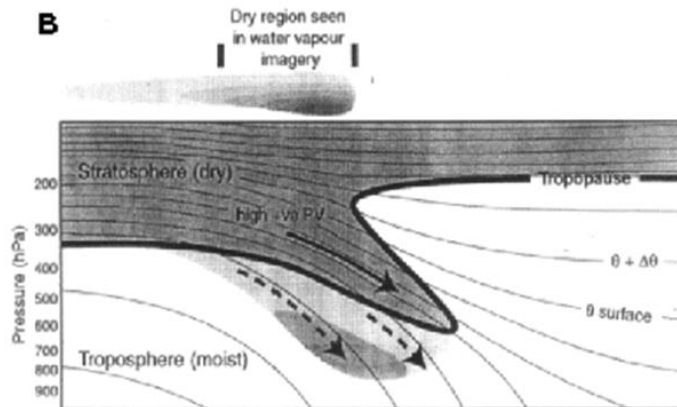
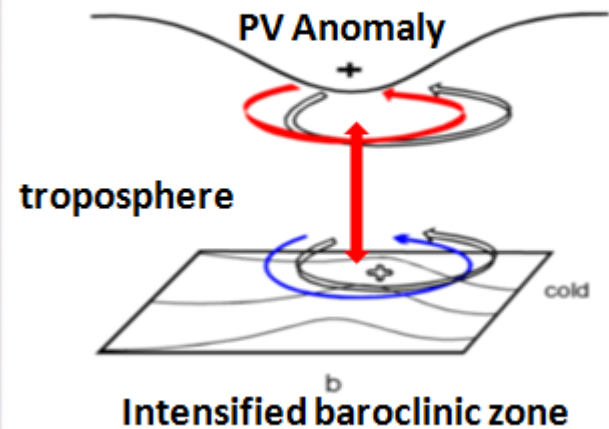
stratosphere



Interaction and mutual intensification of cyclonic circulation:

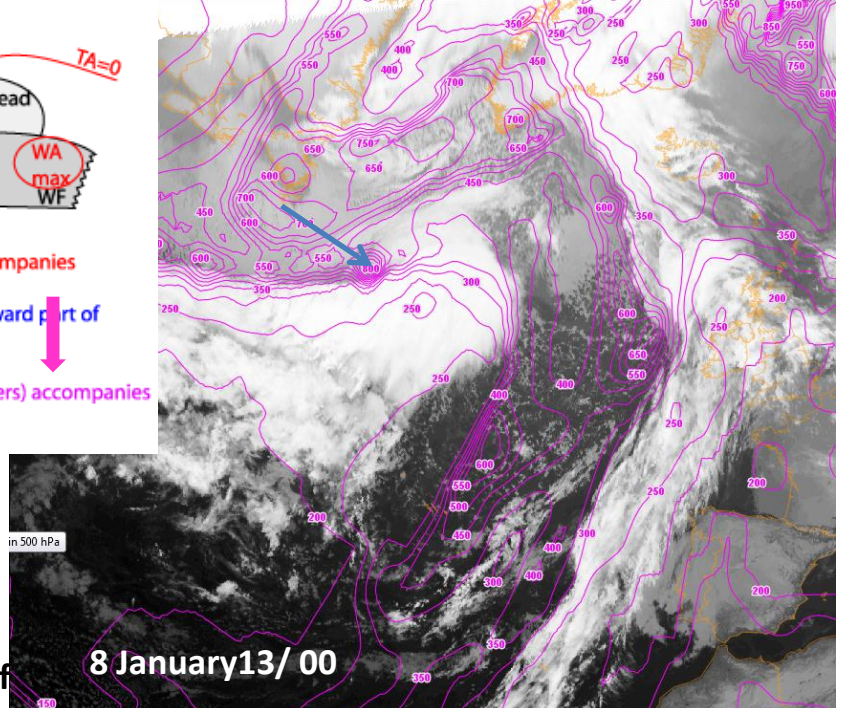
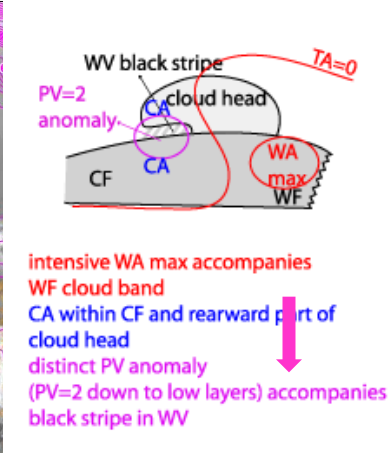
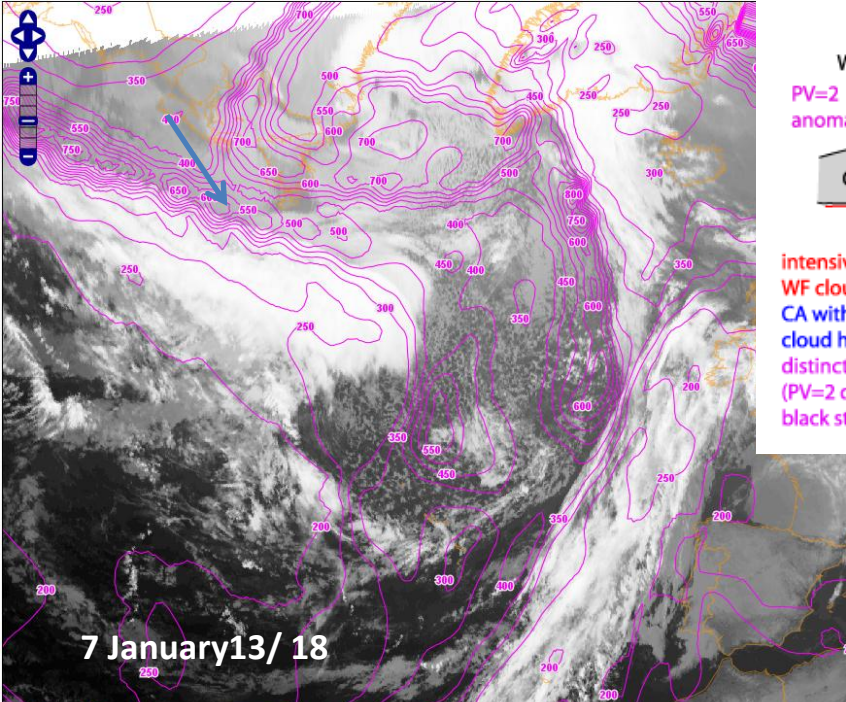
Rapid cyclogenesis

stratosphere

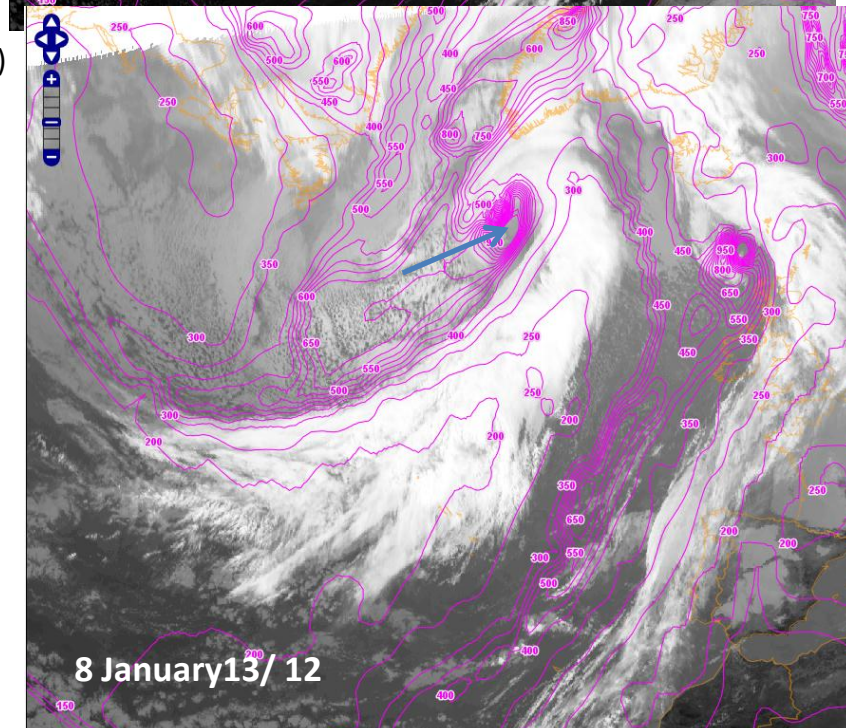
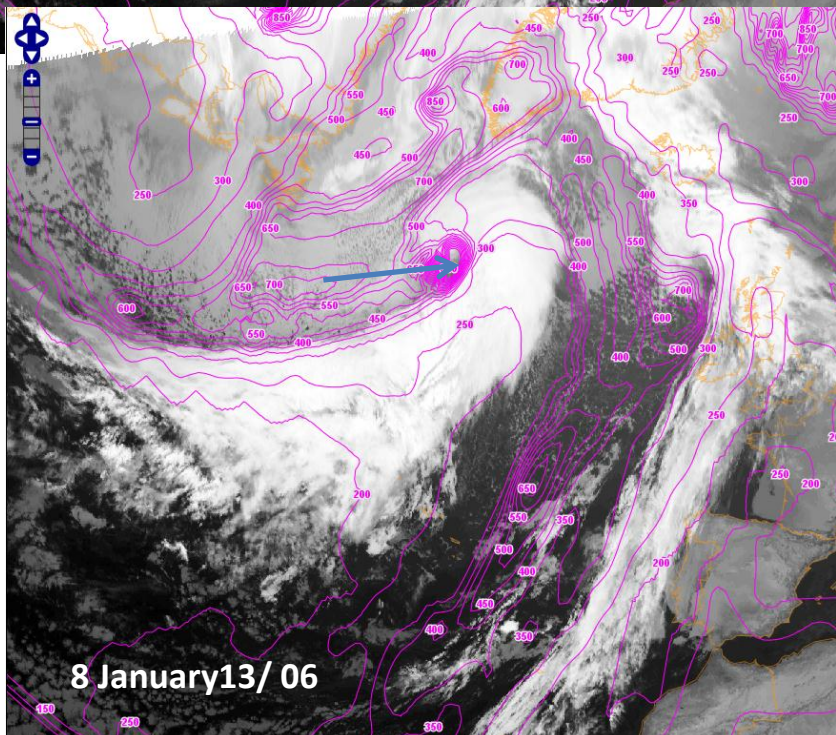


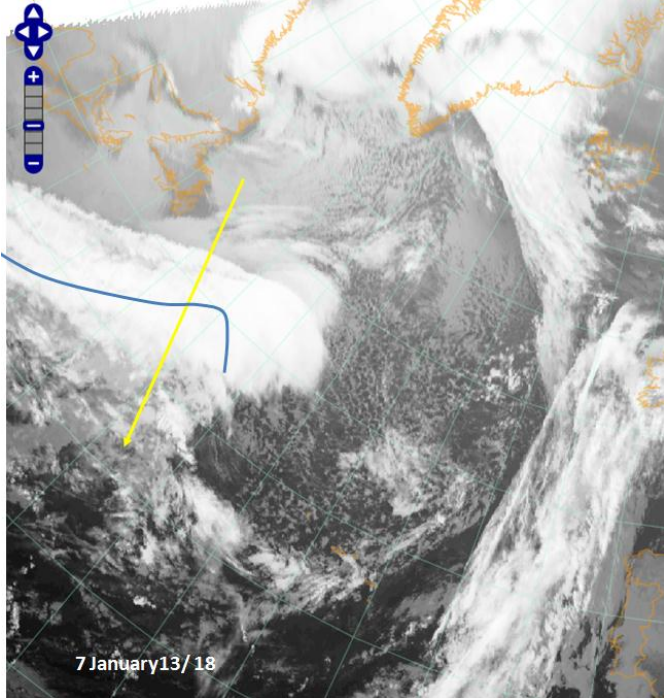
Importance of PV and PV anomaly

- The height of PV ≥ 2 units (-2 on SH) shows how far the air from stratosphere has protruded downward
- In case of classical cyclogenesis the typical height is 400 – 500 hPa
- In case of a rapid cyclogenesis it is down to 600 – 700 hPa and at same time the intensification of the lower tropospheric PV values takes place
- Consequently rapidly a “tunnel” of high PV values from tropopause down to surface develops
- Vertical cross sections can show the downward protrusion below the frontal surface

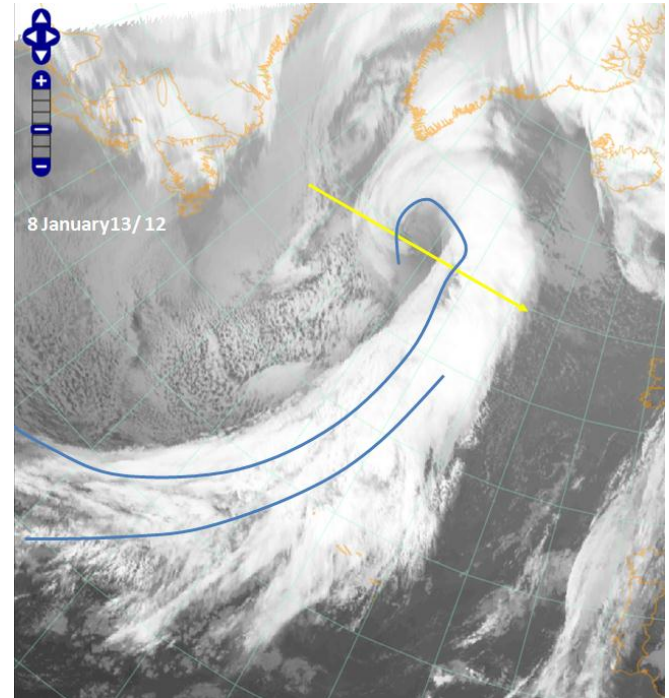
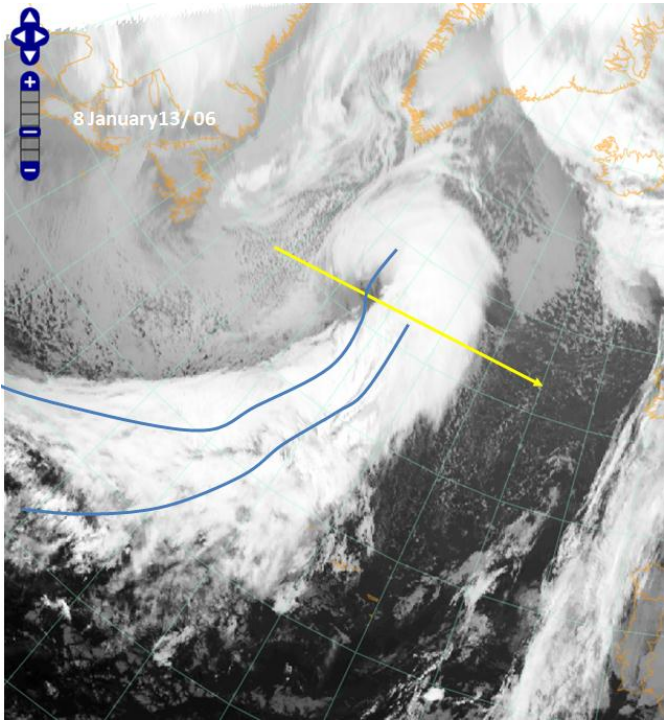
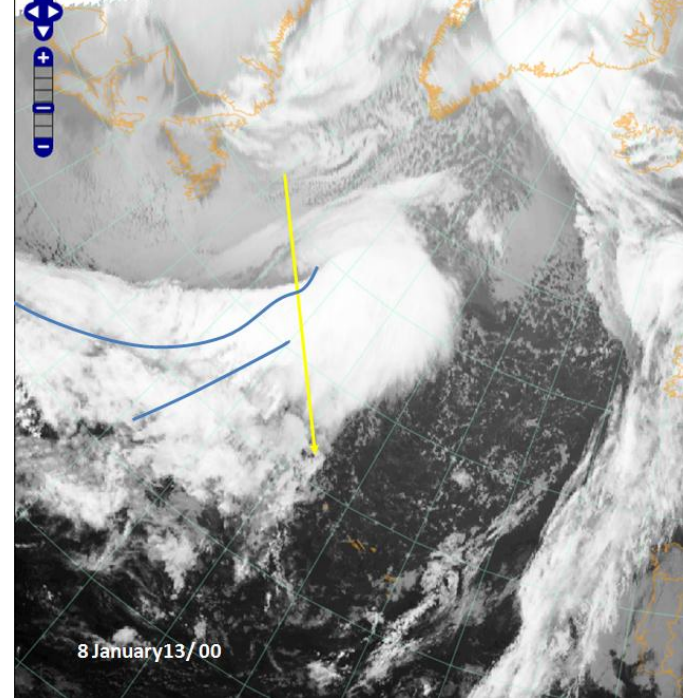


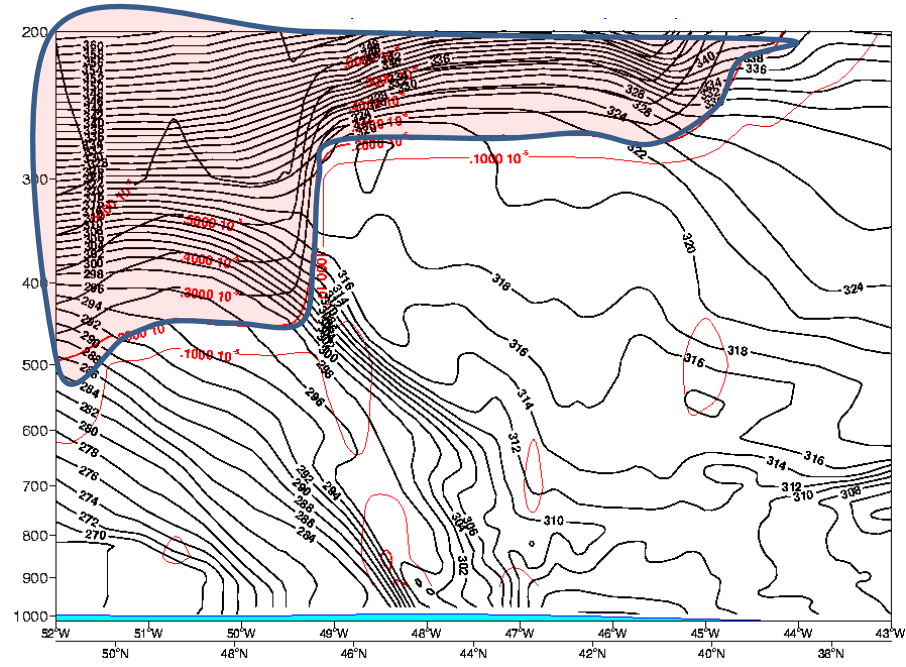
Height of PV (magenta)





**Vertical cross section
lines
through the centre
of the
Rapid cyclogenesis**





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**Rapid
Cyrogenesis:
Downward
Protrusion
of**

stratospheric

Air

**“tropopause
folding”**

of

8 January13/ 00

stratospheric

Air

**“tropopause
folding”**

of

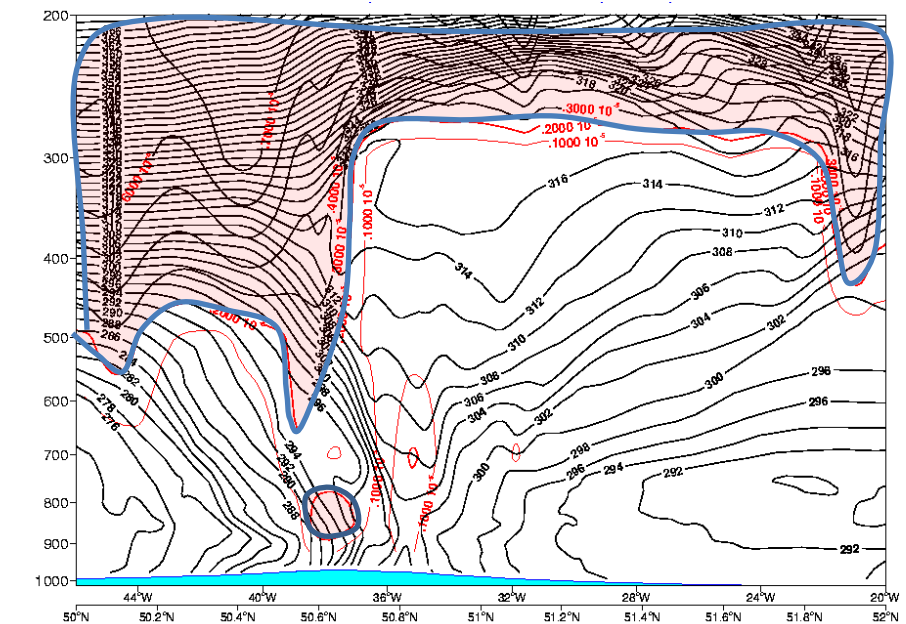
8 January13/ 12

stratospheric

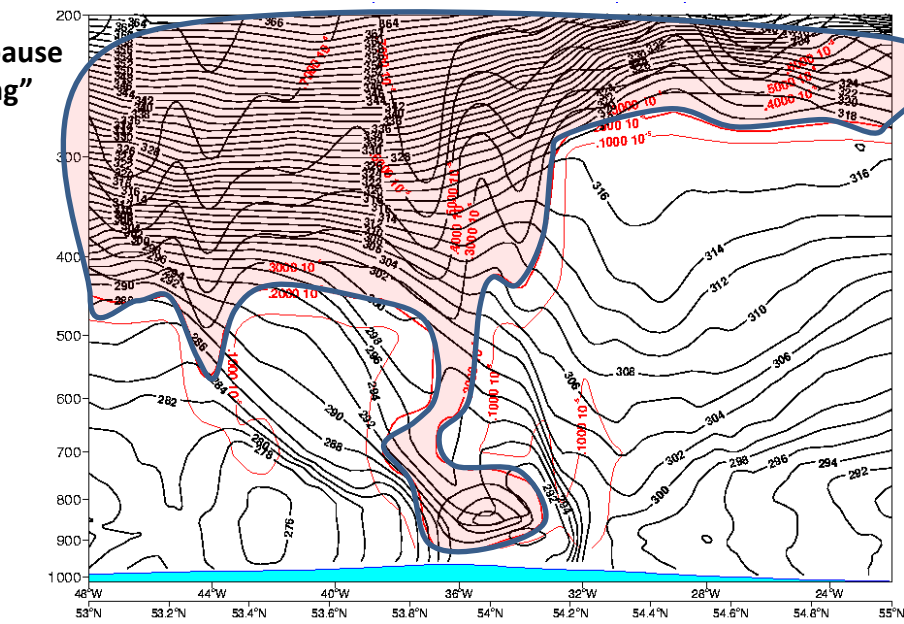
Air

**“tropopause
folding”**

of

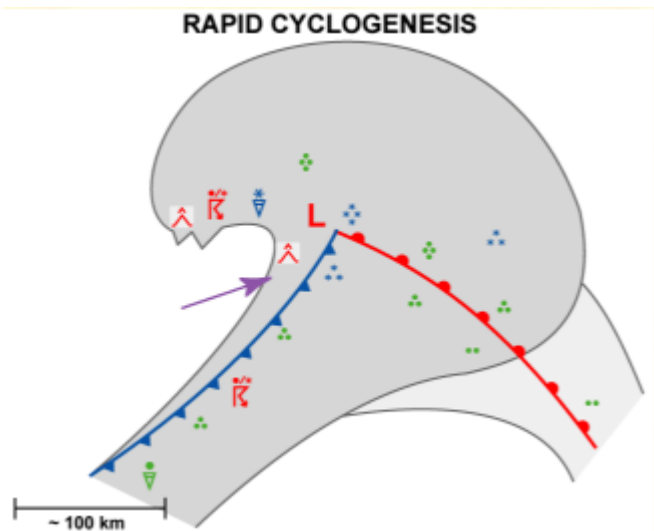
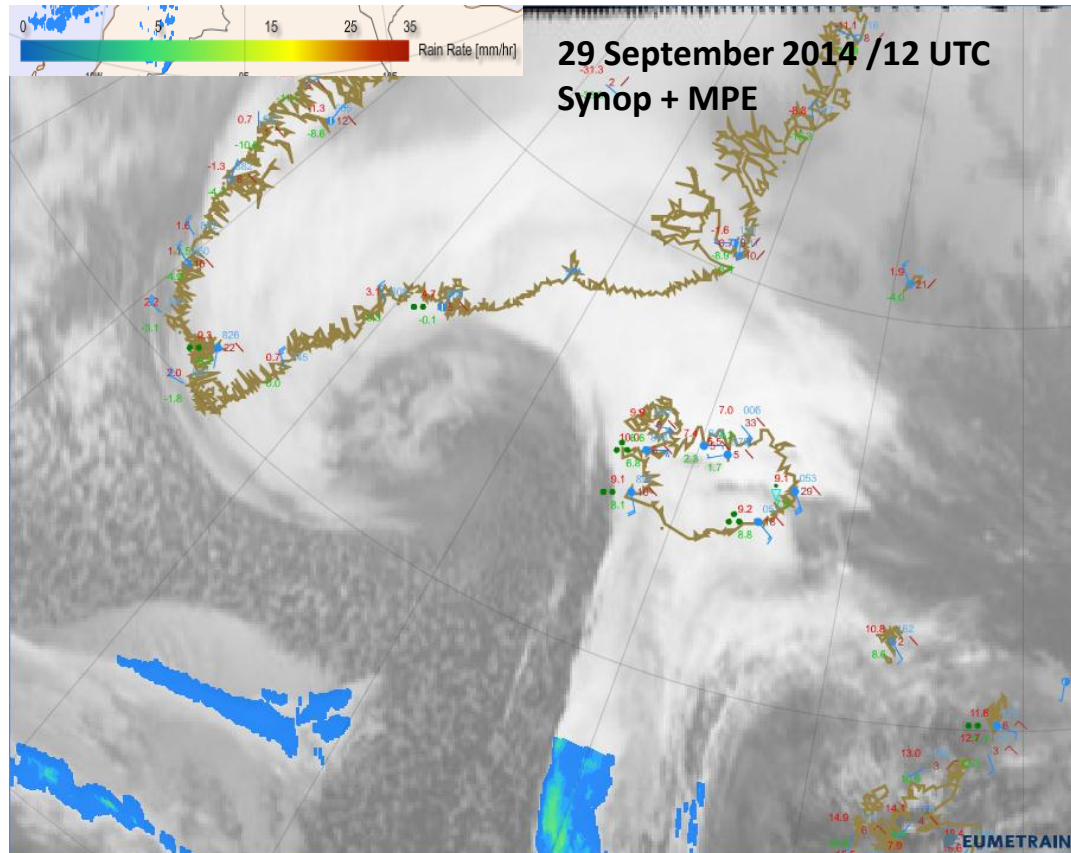


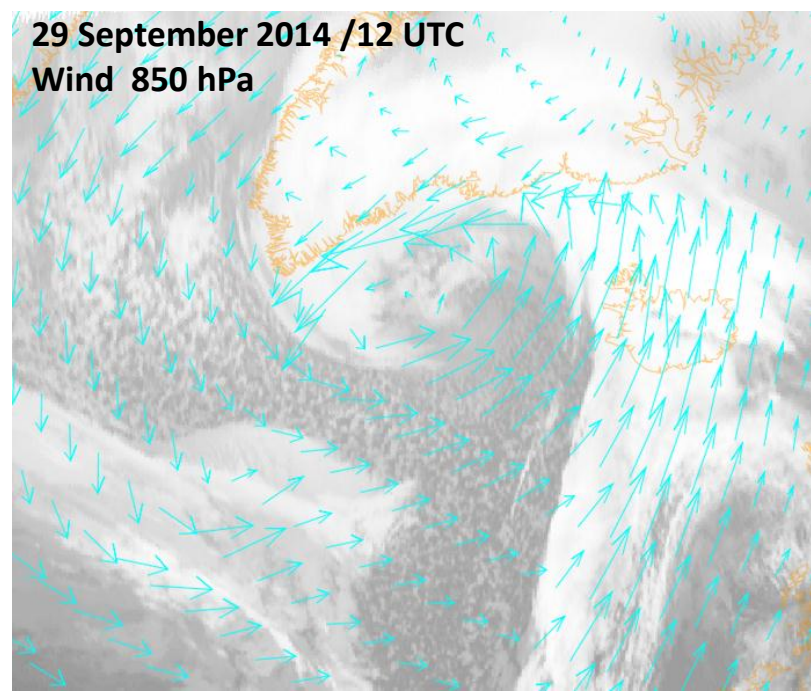
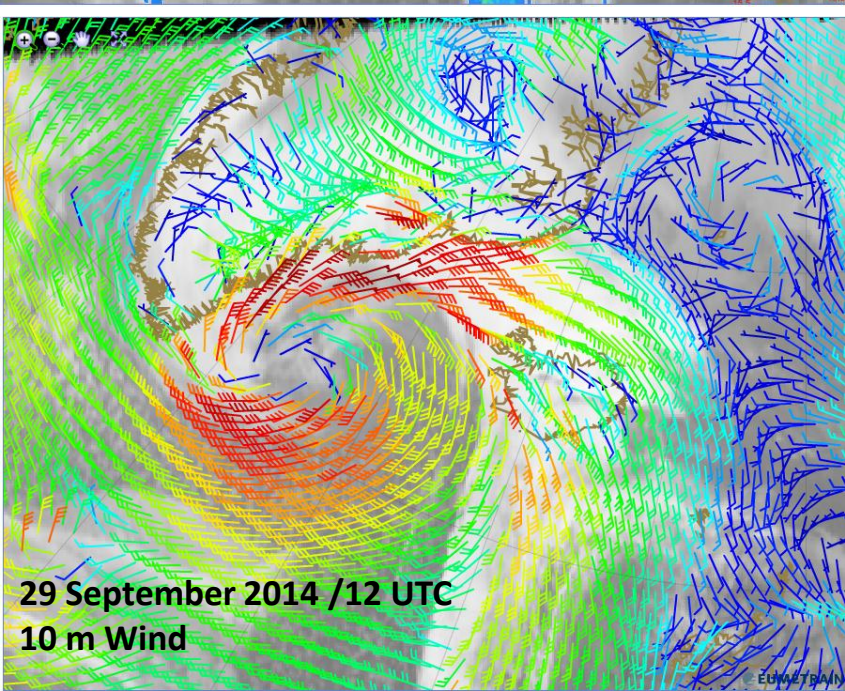
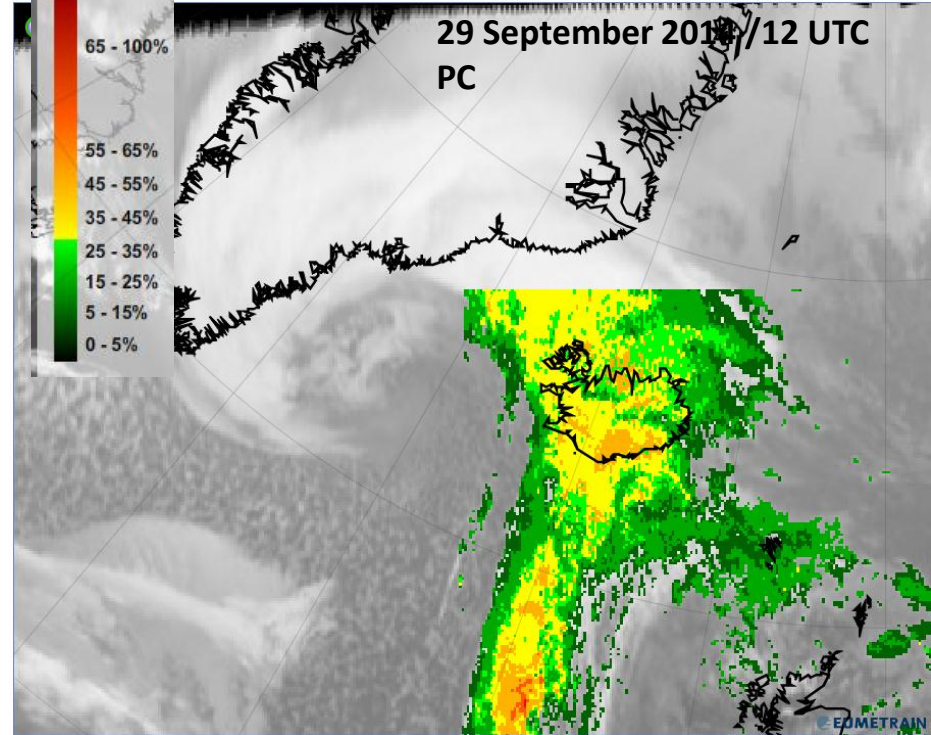
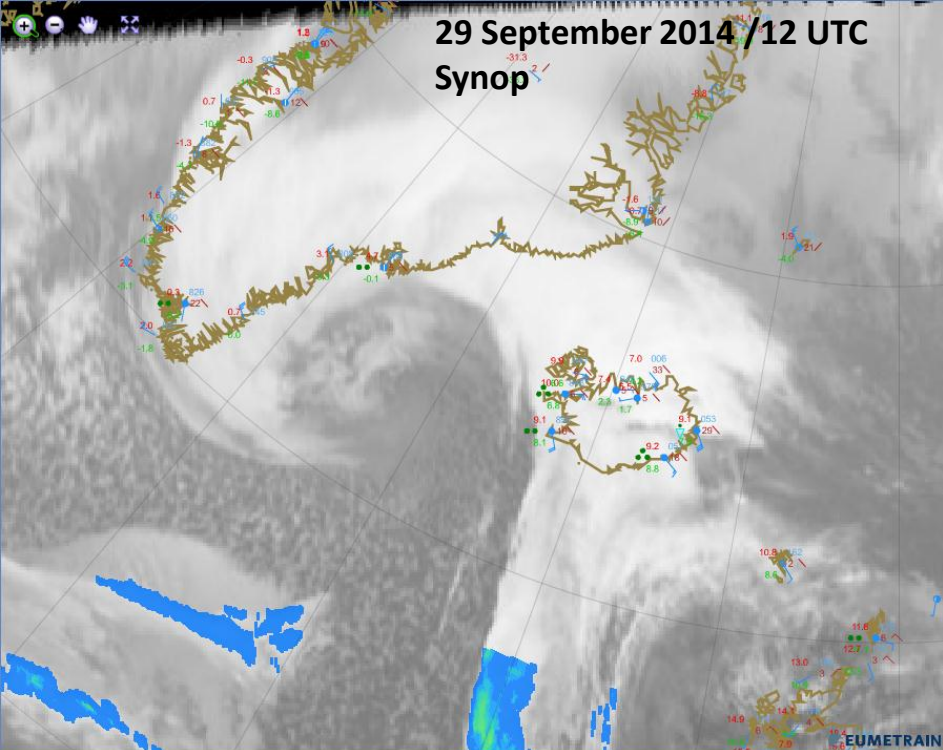
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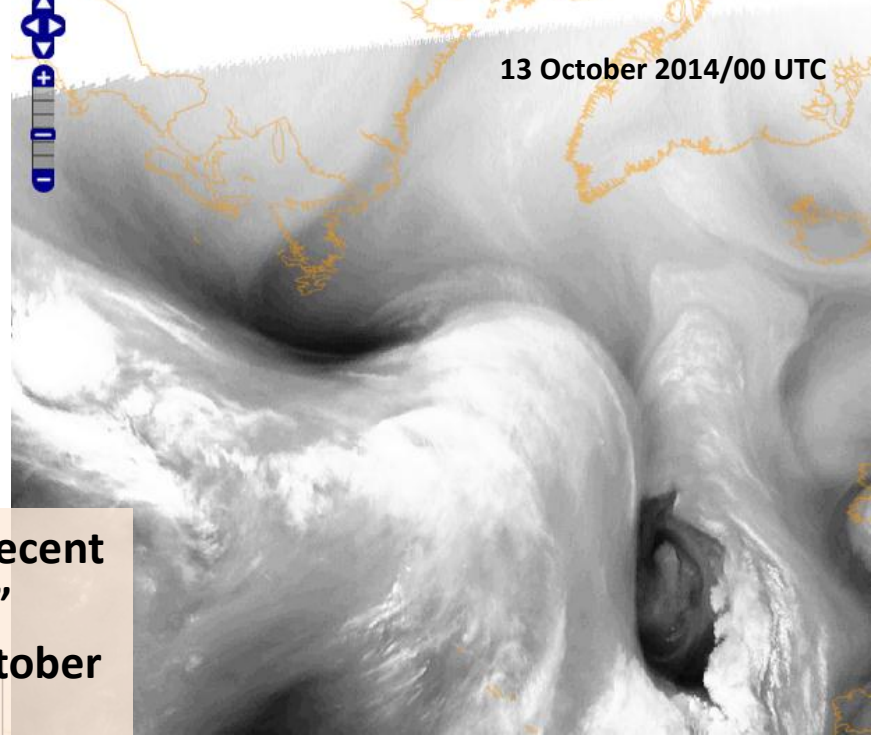
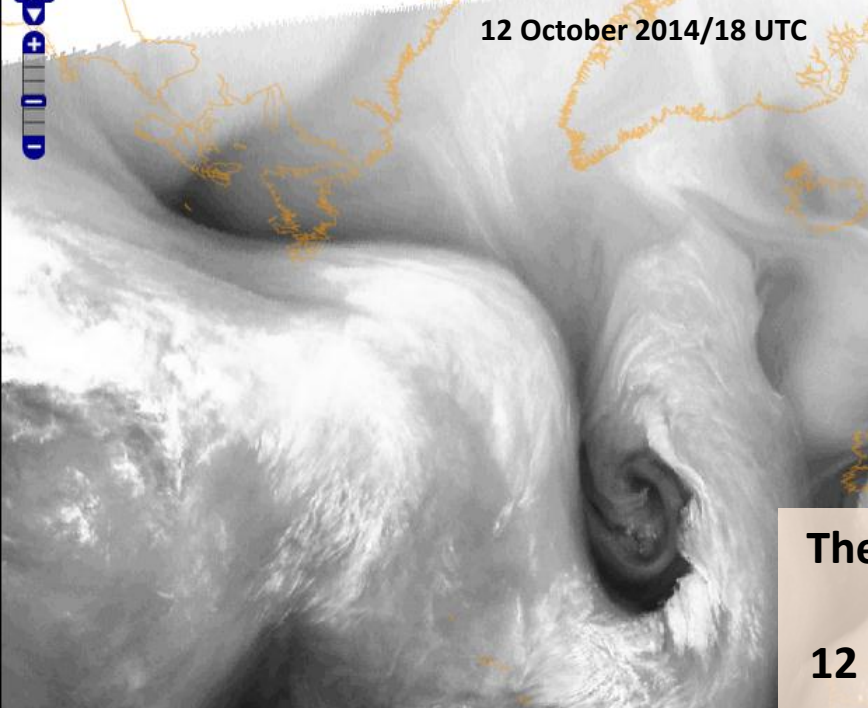


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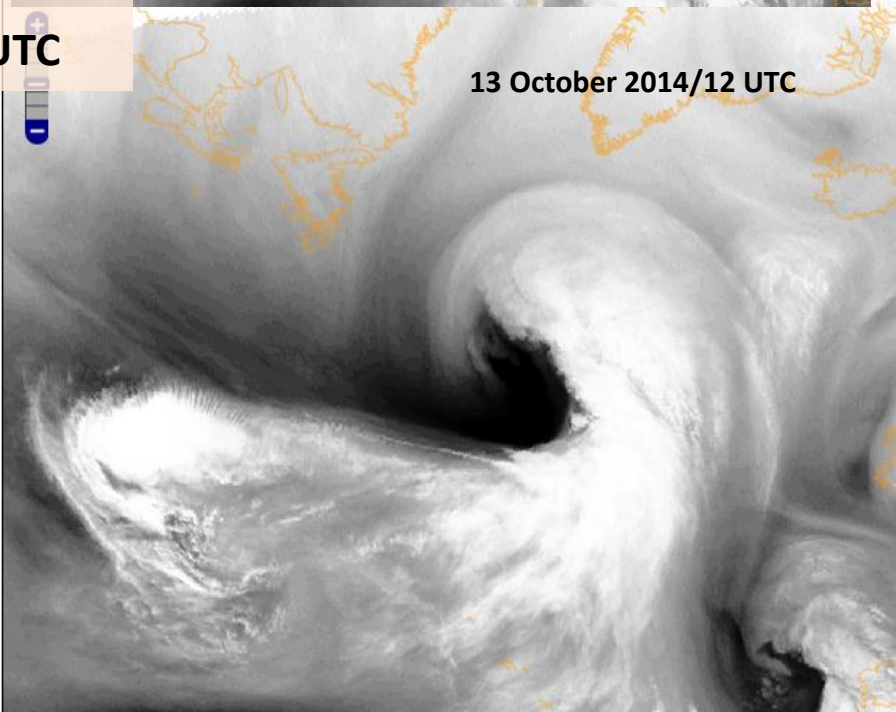
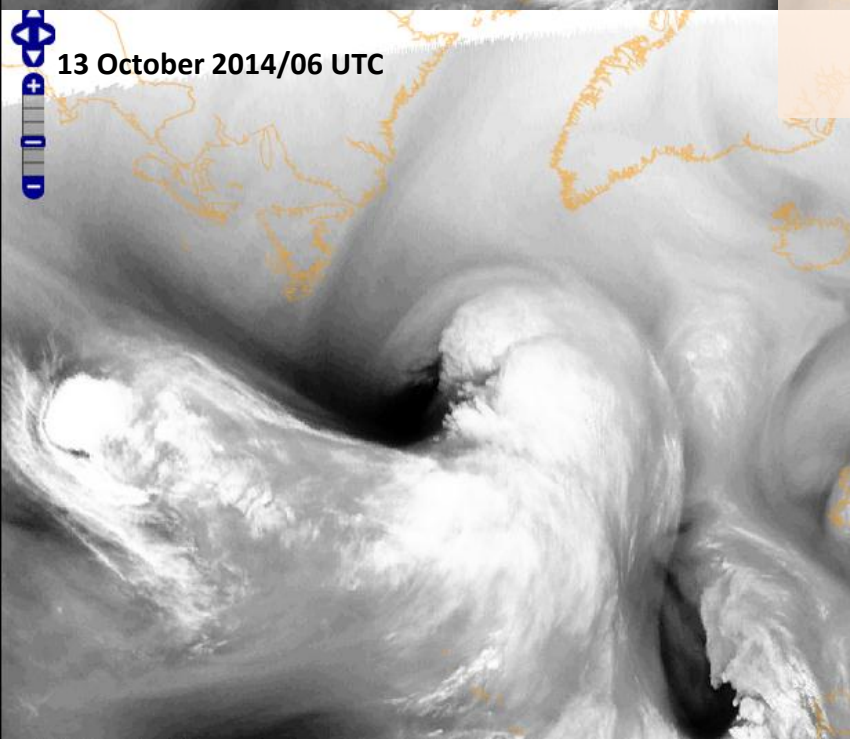
Parameter	Description
Precipitation	<ul style="list-style-type: none"> • Intense precipitation associated with the Warm Conveyor Belt. • Thunderstorms within the inner edge of the cloud head at the western side of the surface low.
Temperature	<ul style="list-style-type: none"> • Strong rise in surface temperature within the area of Warm Conveyor Belt.
Wind (incl. gusts)	<ul style="list-style-type: none"> • Strong winds within the area of the cloud head. In extreme cases winds reaching hurricane force. • Very strong gusts in the transition zone between dark and white in the WV channels near the cold front and the cloud head
Other relevant information	<ul style="list-style-type: none"> • Very strong pressure falls and rises. In the initial stage, falls ahead exceed the rises behind.







The most recent
"Racy"
12 – 13 October
2014
18 – 12 UTC

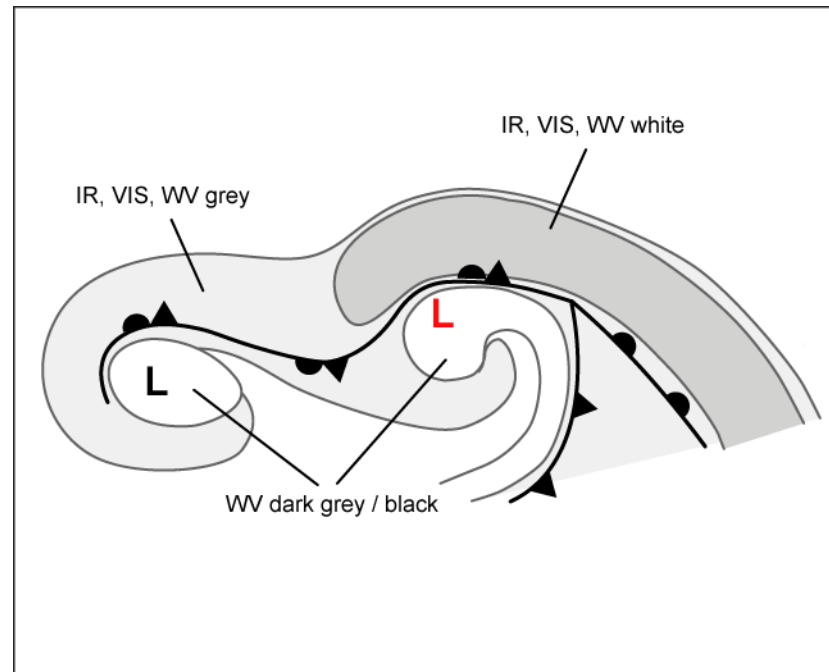


Special Developments at and with occlusion cloud bands

- Substructures developing at occlusion cloud band after some time
 - Development of secondary low centers at occlusion cloud band
- Development of occlusion cloud bands deviating from the classical theories:
 - Back bent occlusion
 - Cold Air Development
 - Instant occlusion

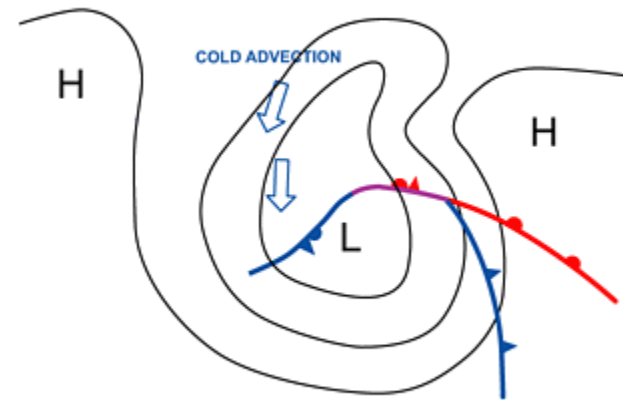
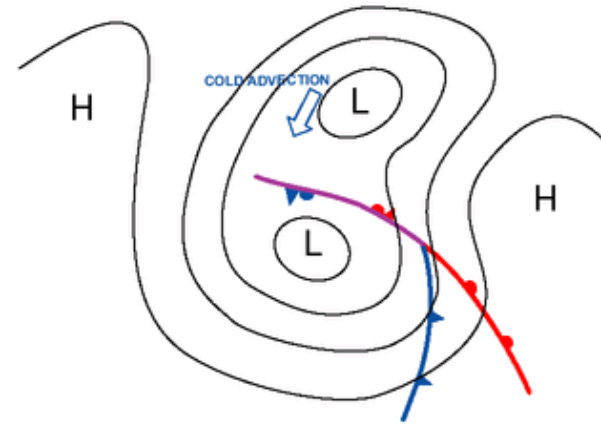
Secondary Low Centers at Occlusion Cloud bands

- A secondary Low develops at an occlusion cloud band if there is :
 - a distinct maximum of vorticity both shear and curvature components
 - a jet streak on the cyclonic side of the occluded front and parallel to it
 - Wind velocity in the jet streak are at least 50 m/s



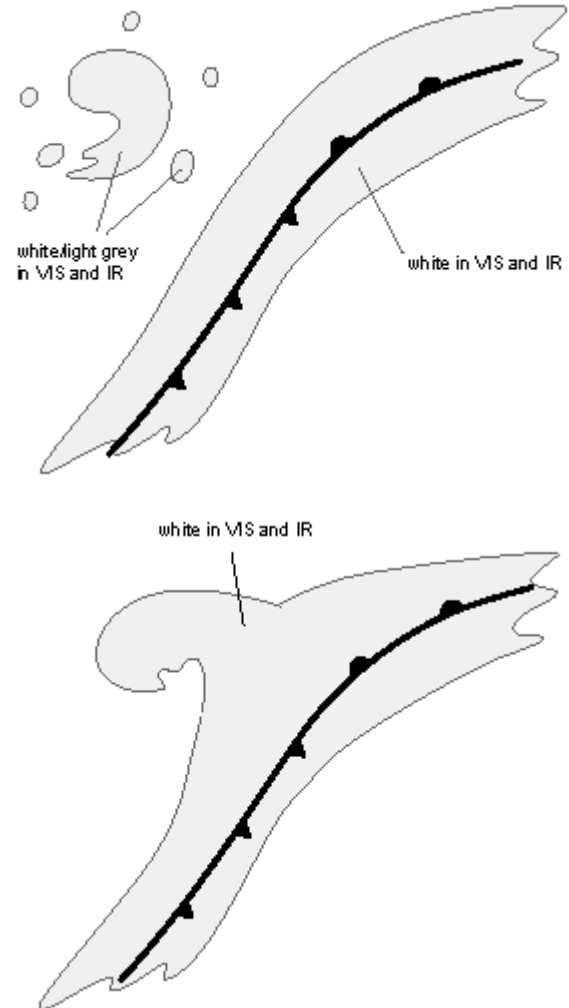
Backbent Occlusion

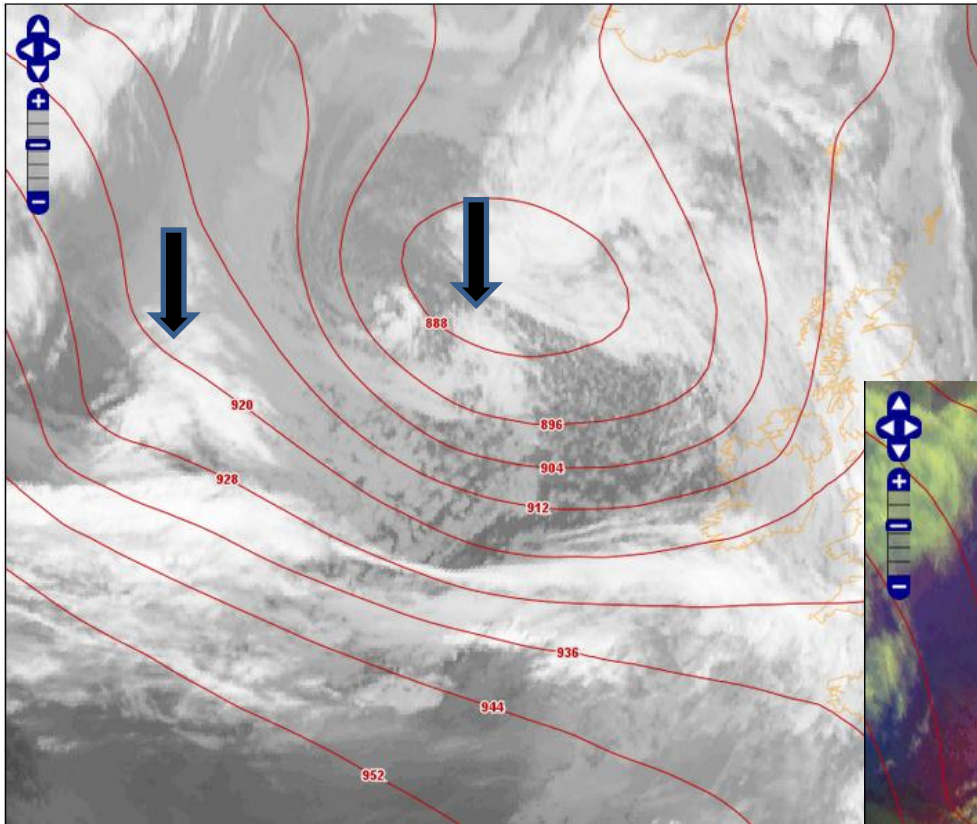
- Describes a development where the curved occlusion spiral changes to a cyclonic curvature
- This part takes the qualities of a Cold front
- Therefore the former occlusion propagates eastward as a cold front
- The main reason for this is a protruding cold air mass at the rear of the former occlusion band
- Consequently upper level height fields and temperature advection are the main key parameters



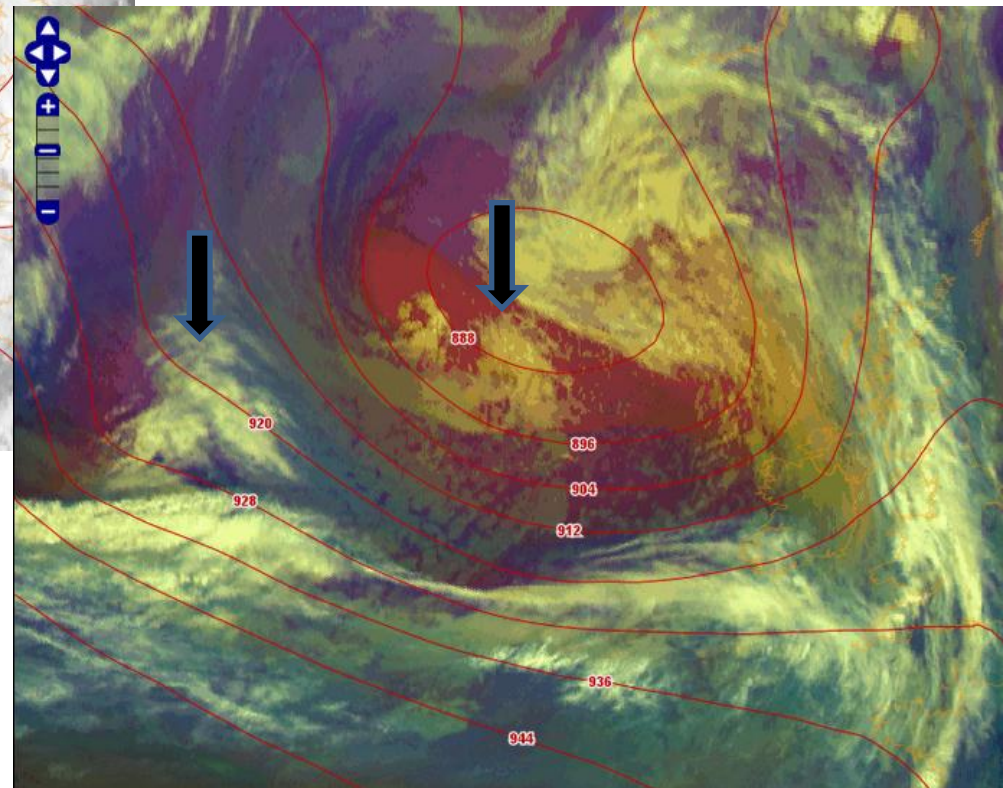
Instant Occlusion

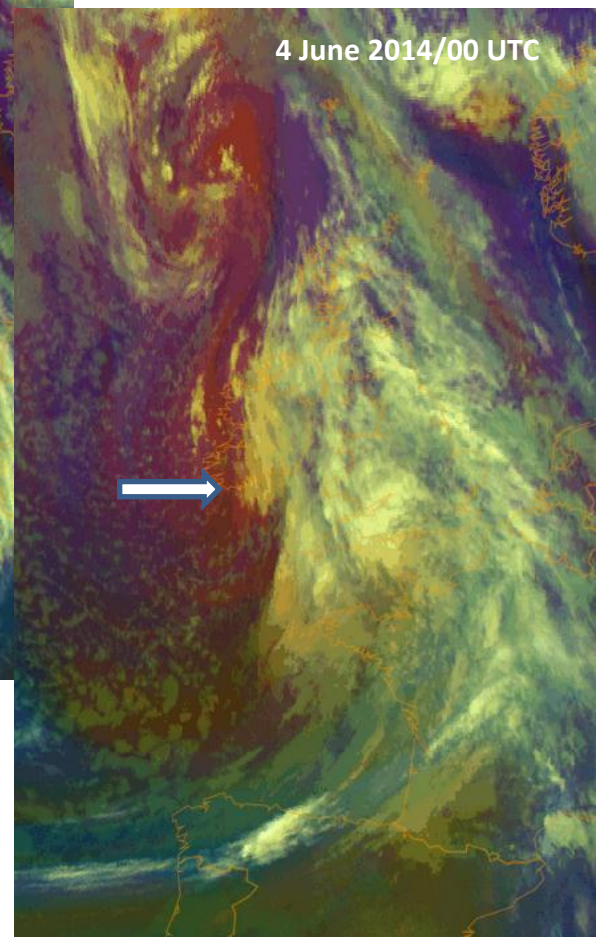
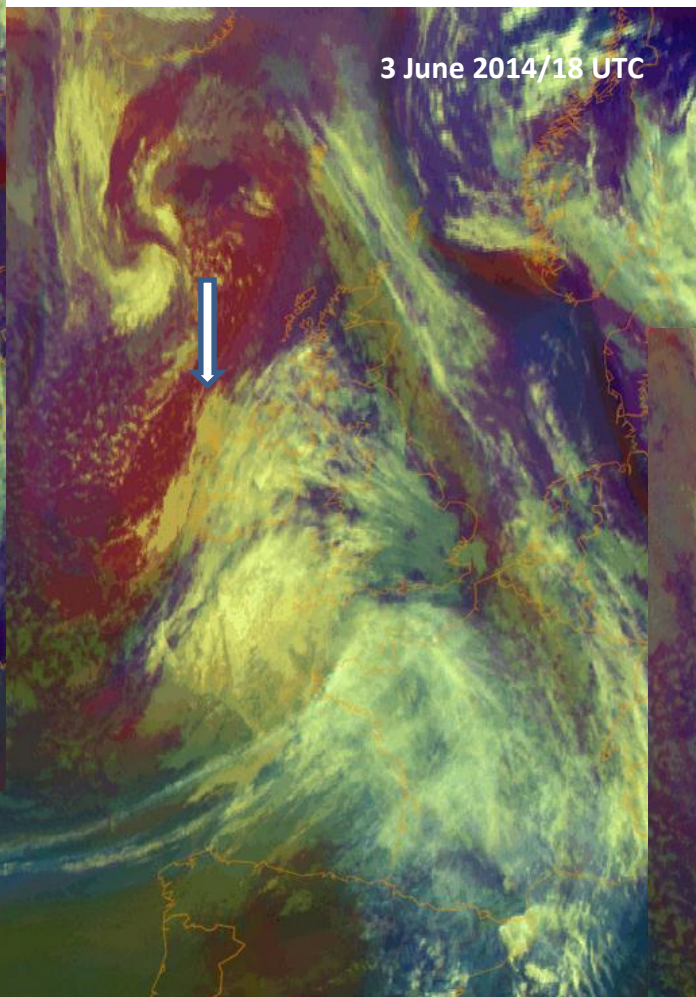
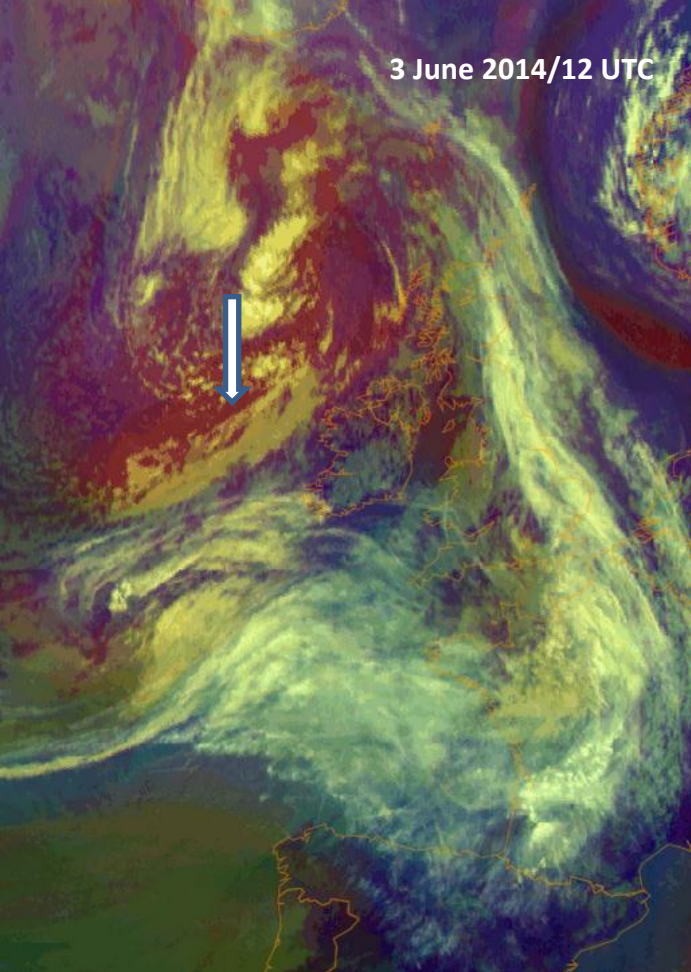
- Describes a development where a comma features in the cold air behind a frontal cloud band approaches the cloud band and finally merges with the cloud band
- Consequently the appearance of an occlusion spiral happens suddenly – “Instant occlusion”
- Different displacement velocities are the pre-condition for such a process



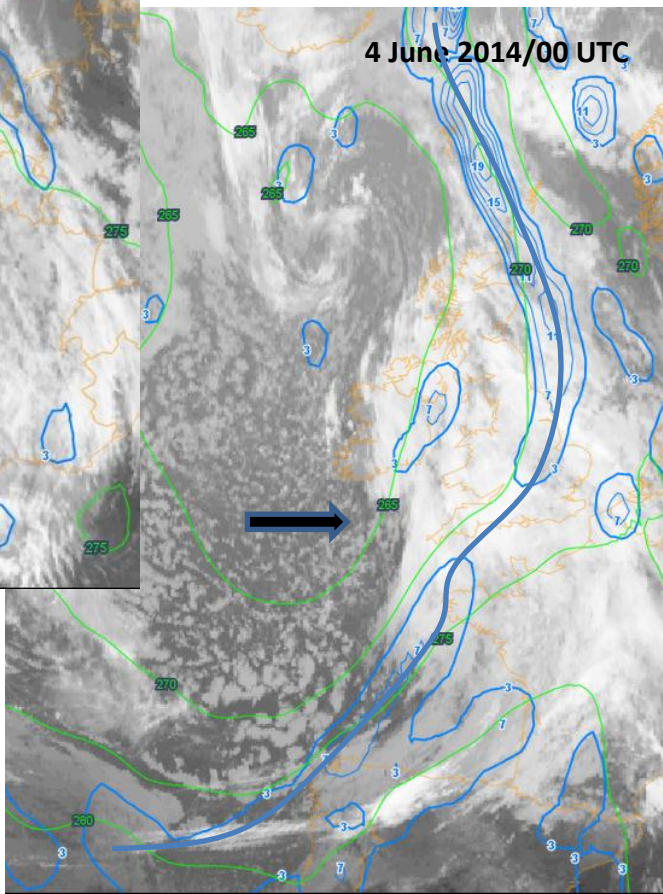
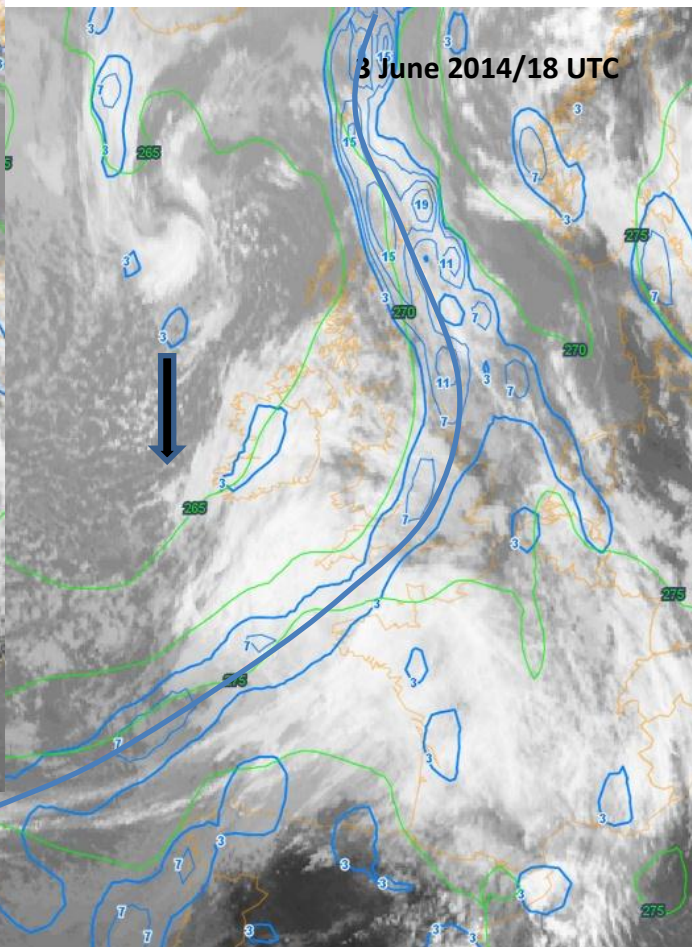
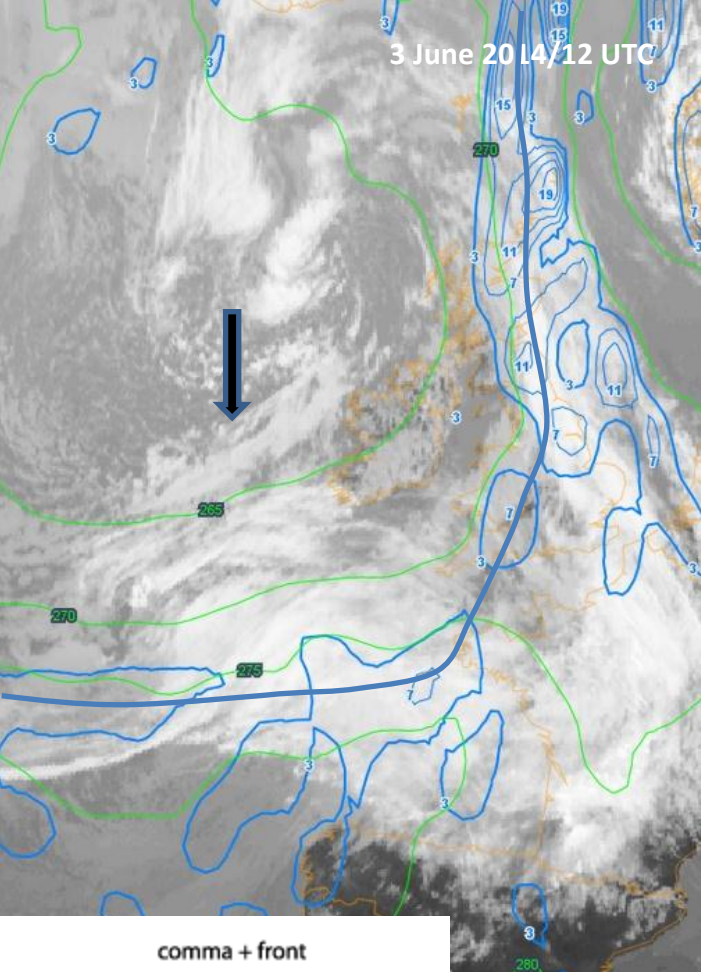


Initial stage
3 June 2014/00 UTC

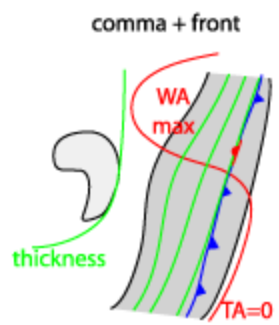




**3 – 4 June 2014
Instant Occlusion
Airmass RGB
images**

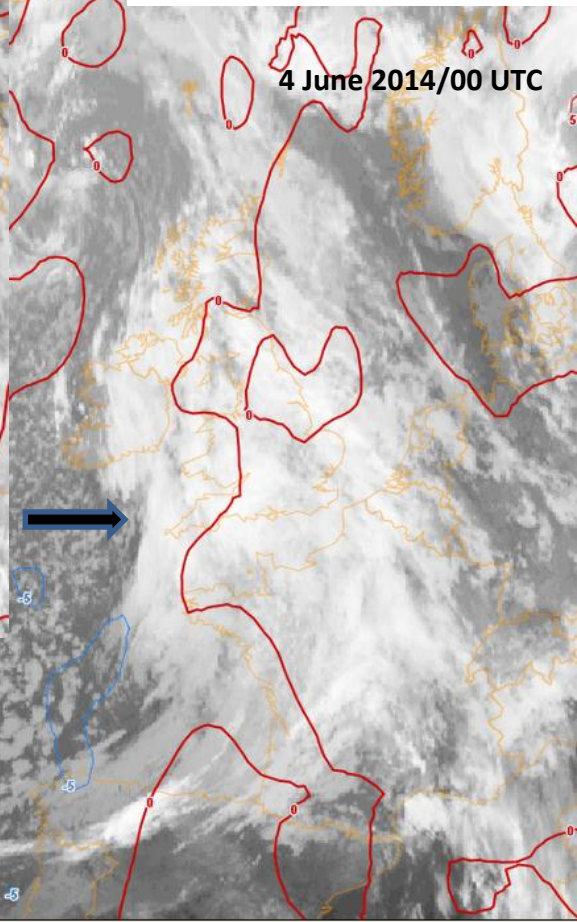
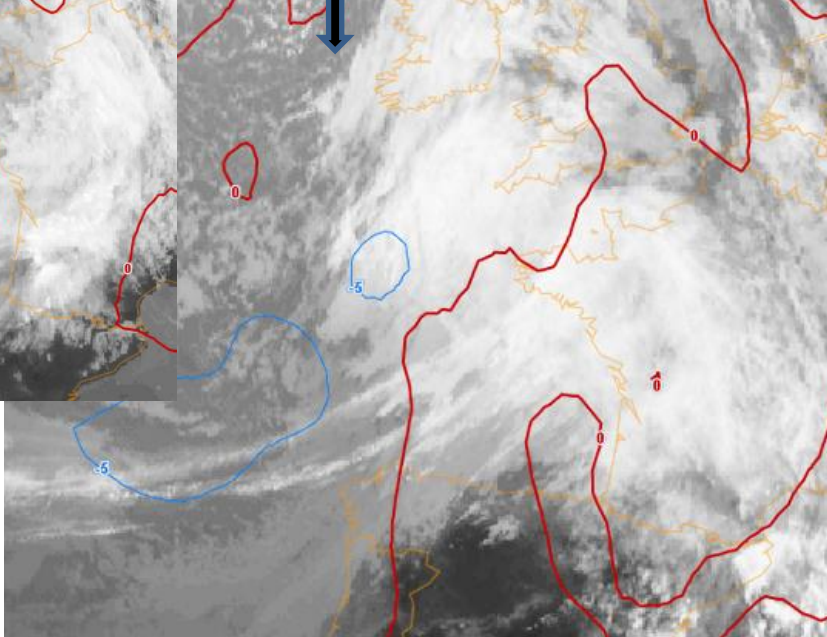
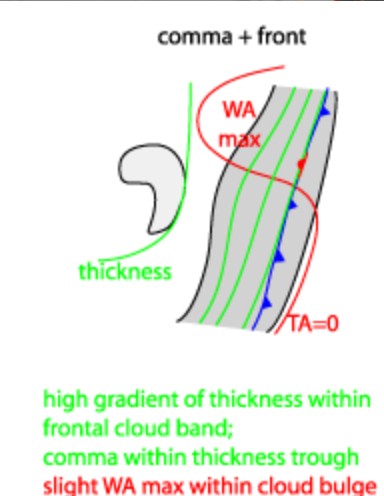
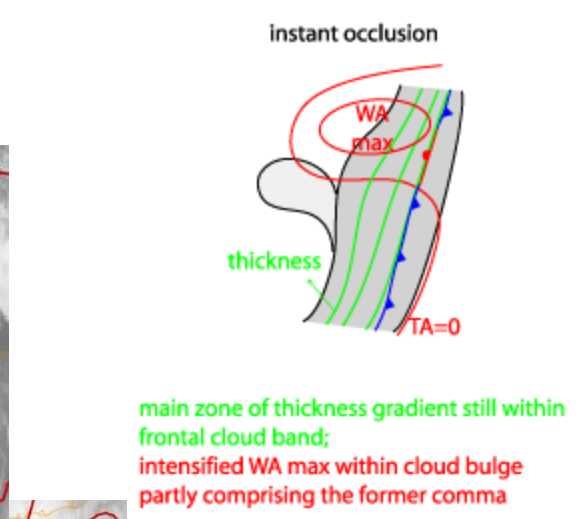
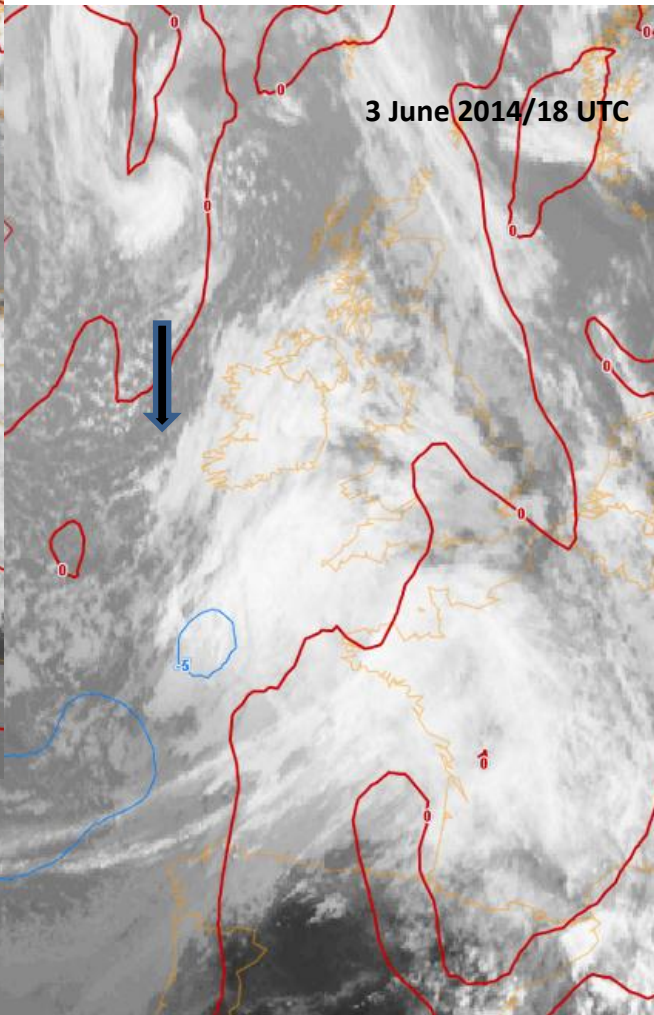
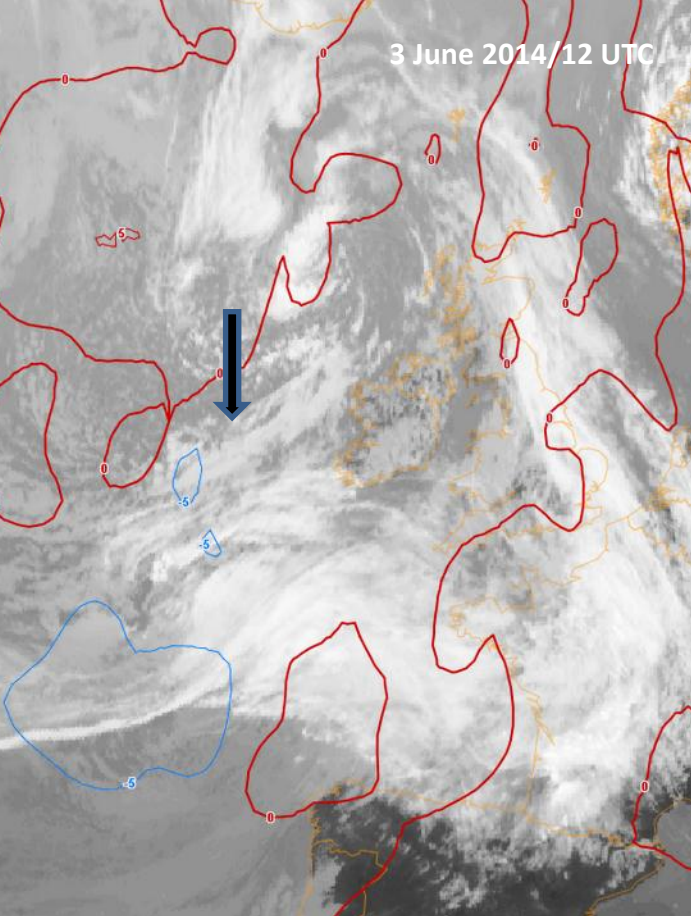


main zone of thickness gradient still within frontal cloud band;
intensified WA max within cloud bulge partly comprising the former comma

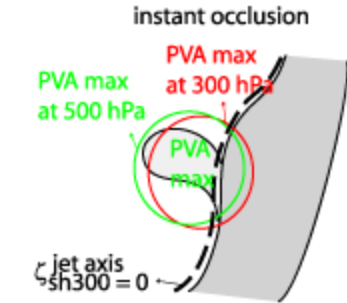
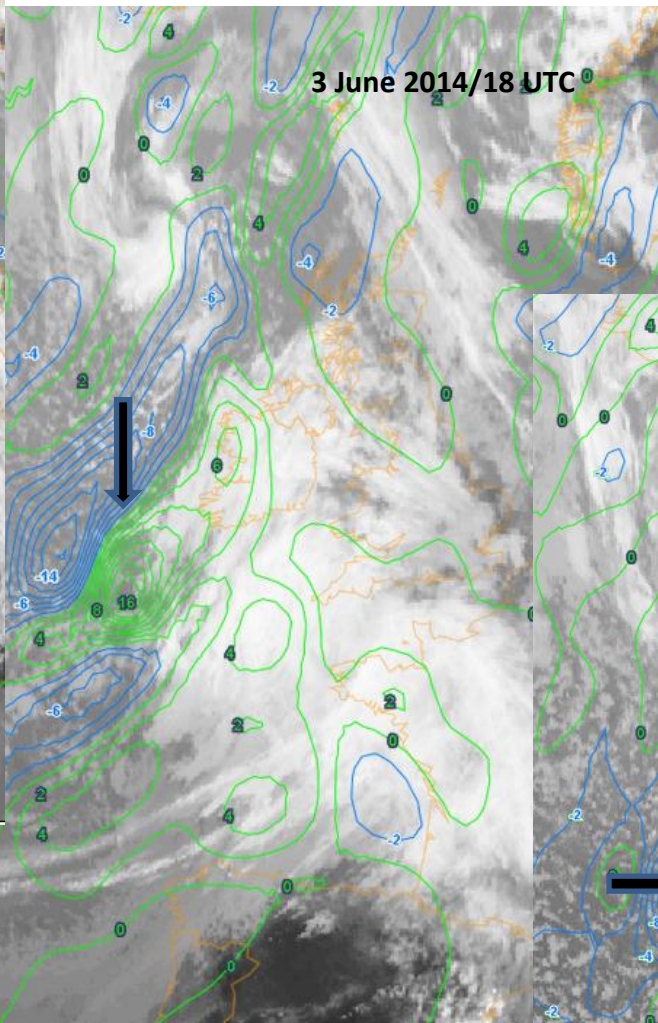
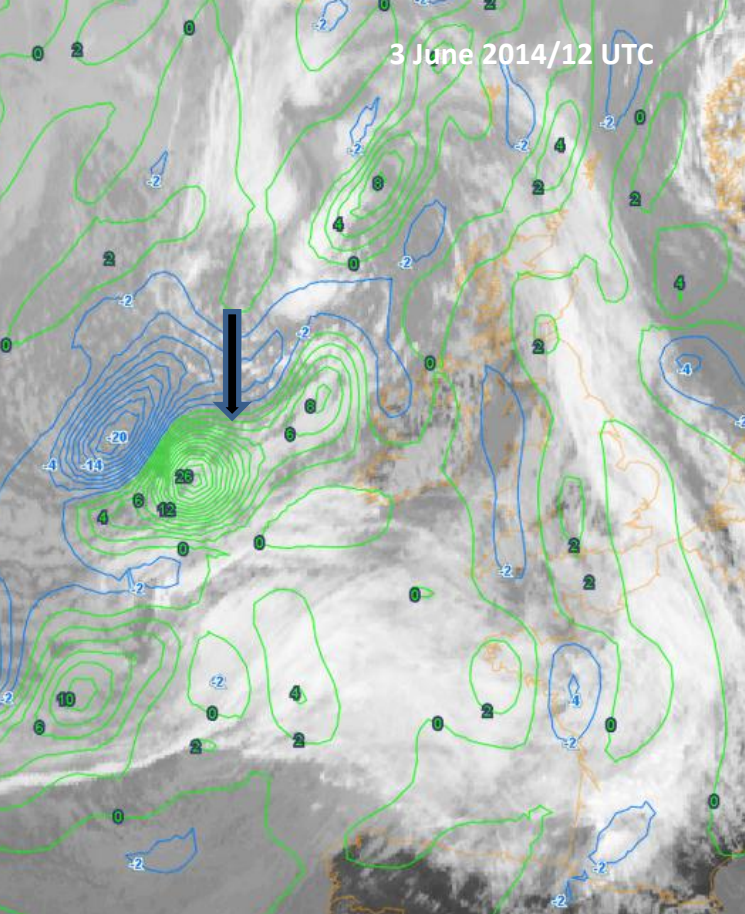


high gradient of thickness within frontal cloud band;
comma within thickness trough
slight WA max within cloud bulge

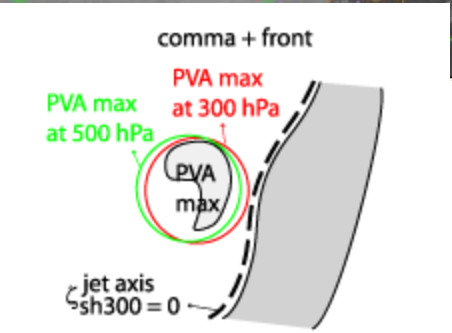
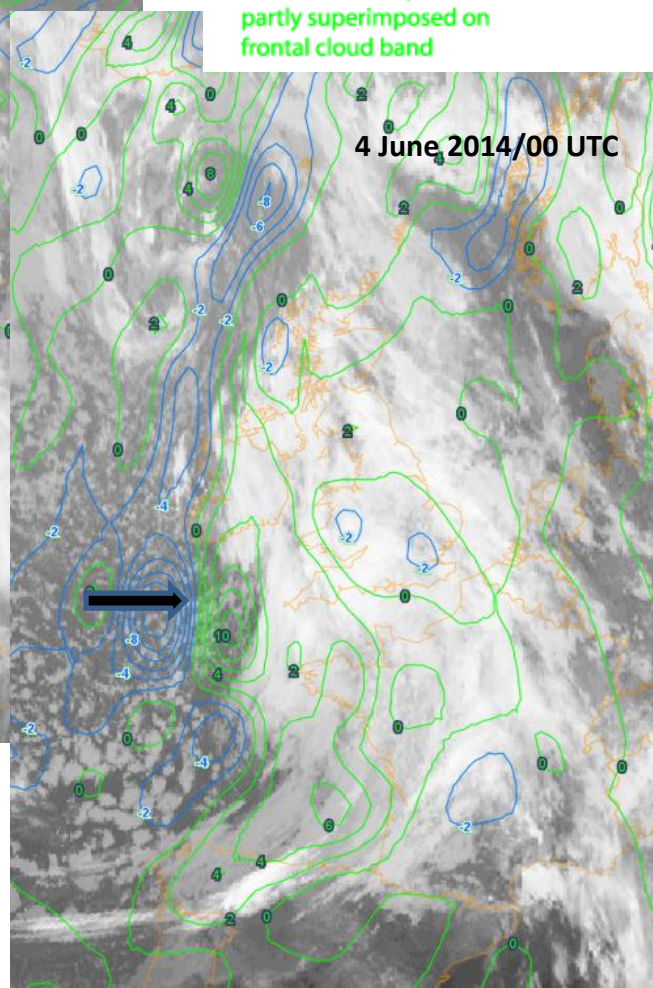
3 – 4 June 2014
Instant Occlusion
Green: equiv. Thickness
Blue: TFP



3 – 4 June 2014
Instant Occlusion
Red/blue: temperature advection 700 hPa
Red: WA, blue: CA



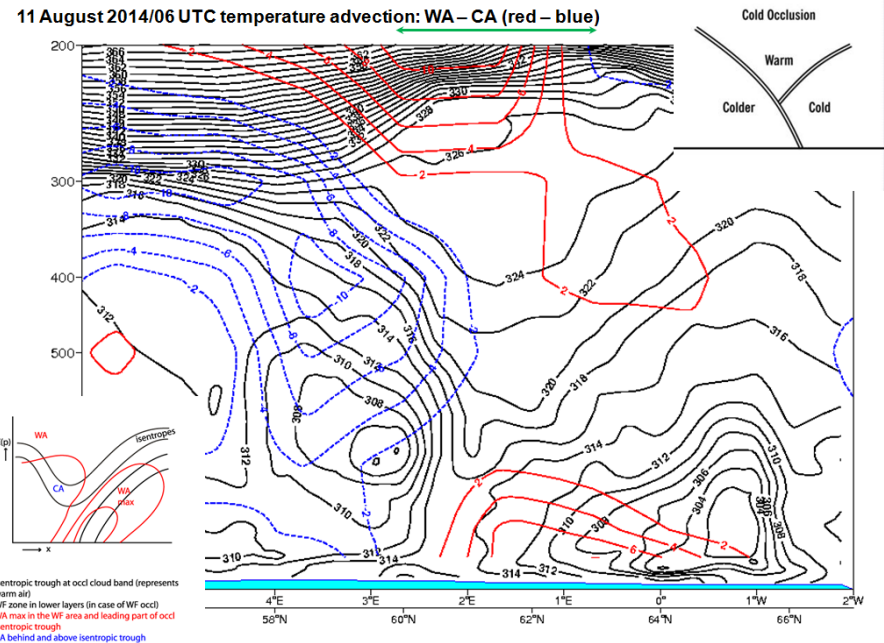
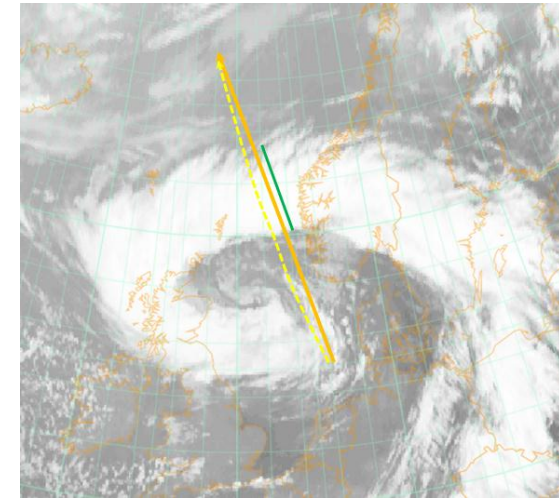
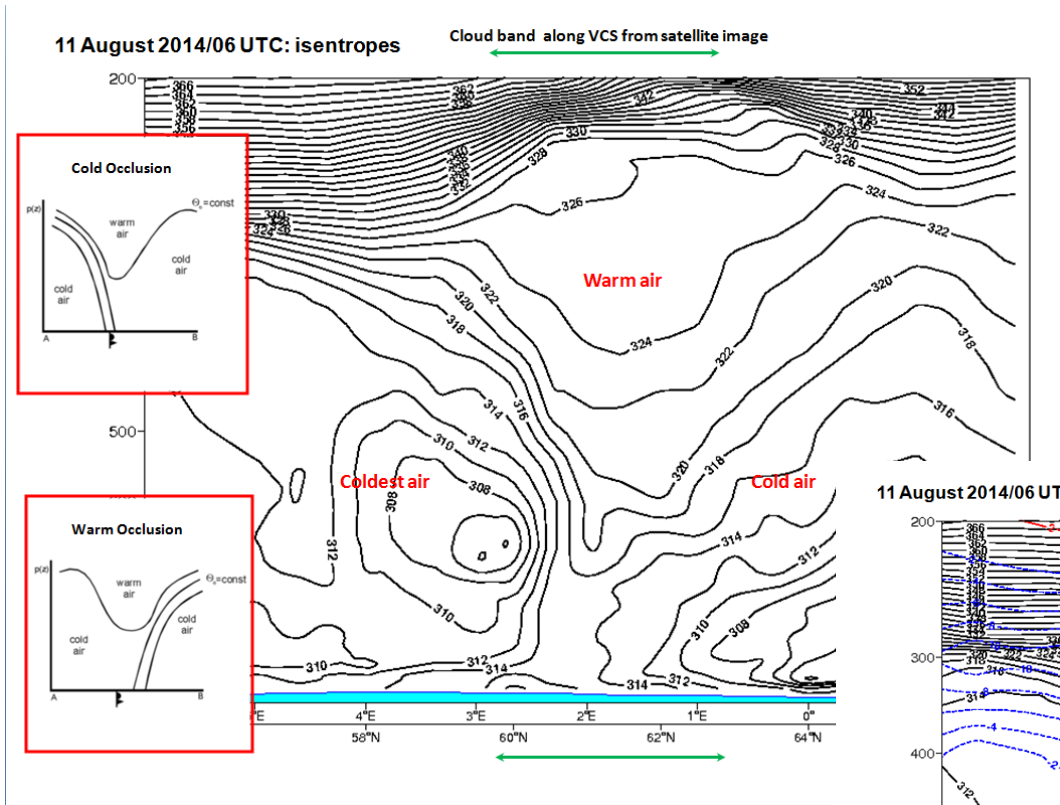
jet axis along rear side of frontal cloud band
PVA max accompanies comma and is partly superimposed on frontal cloud band



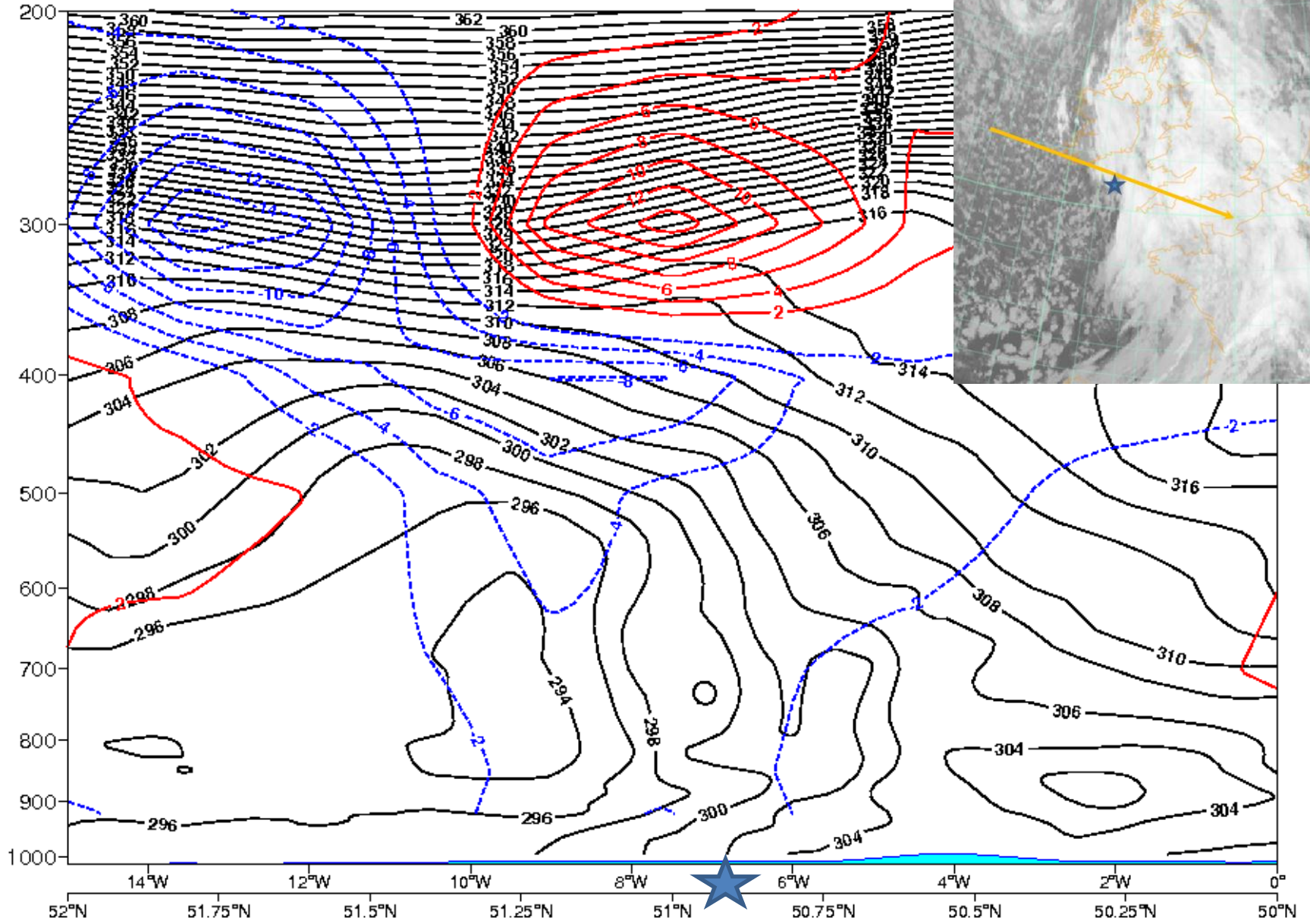
jet axis along rear side of frontal cloud band
PVA max at 500 hPa and 300 hPa above comma

3 – 4 June 2014
Instant Occlusion
Green/blue: Vorticity advection 700 hPa
Green: PVA, blue: NVA

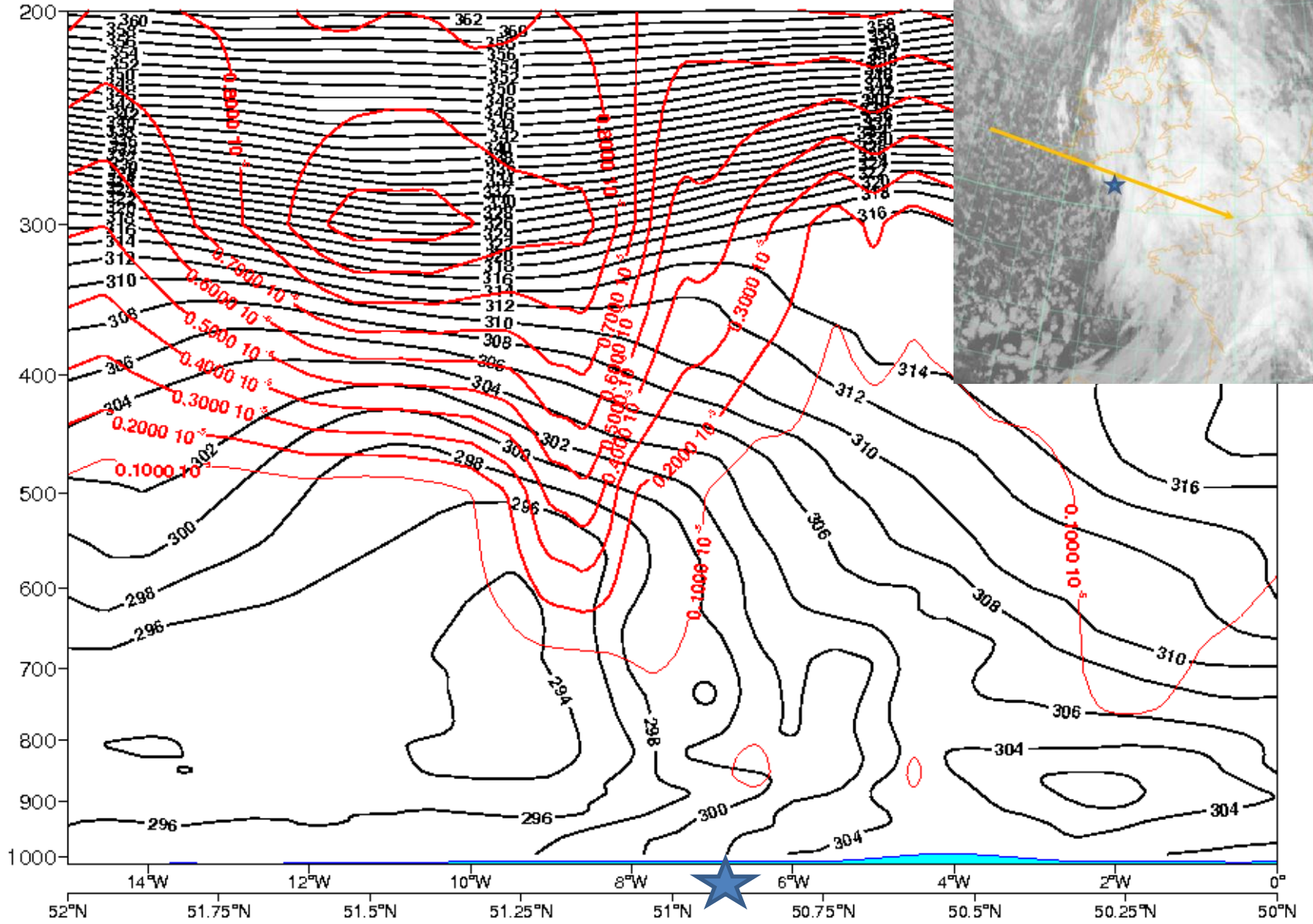
What typical features do we expect for a classical occlusion process ?



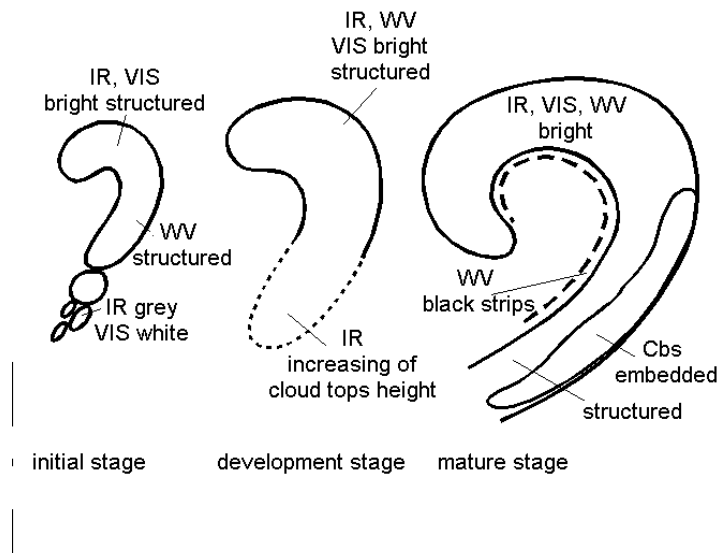
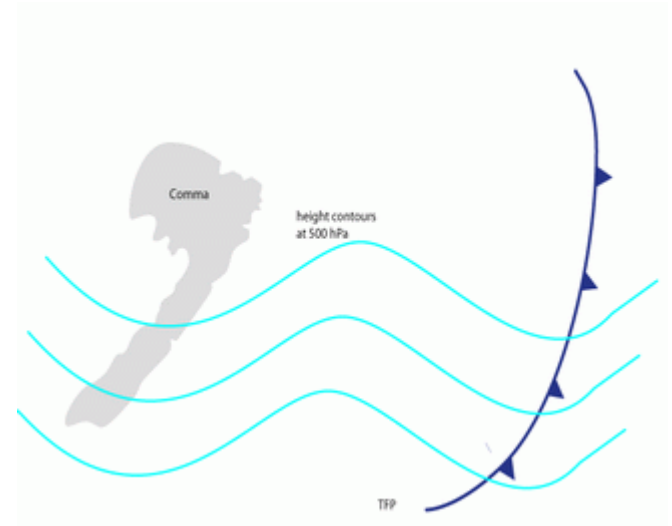
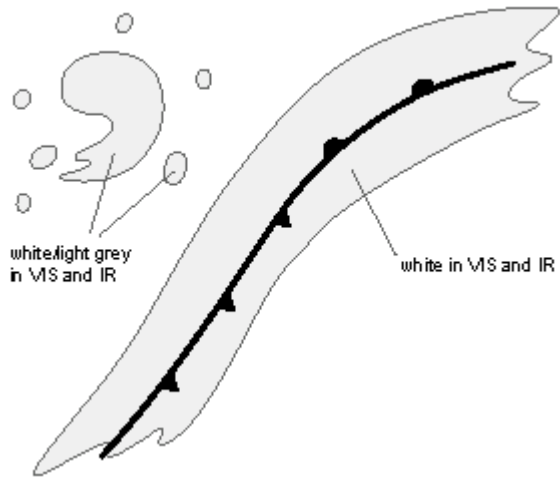
4 June 14/00

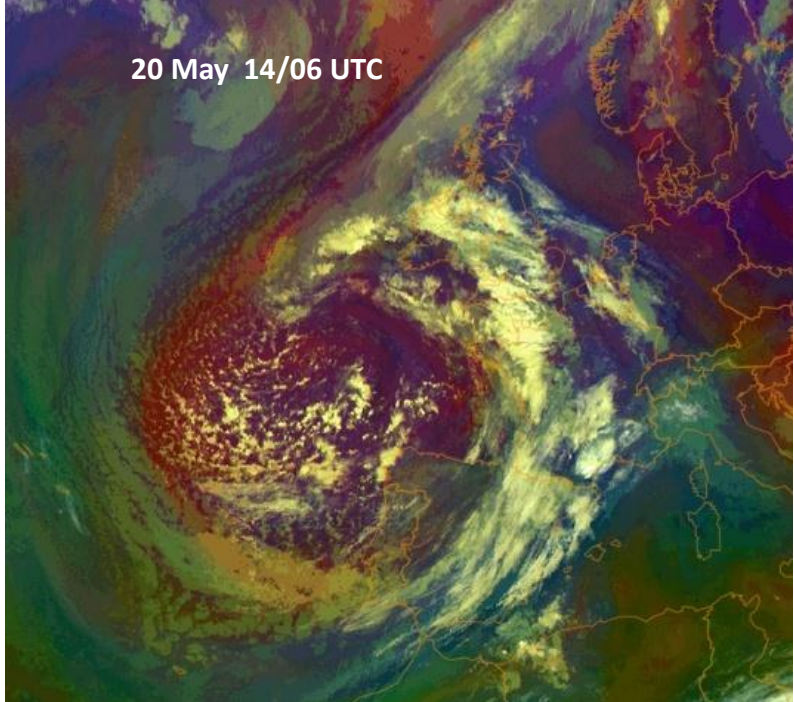


4 June 14/00



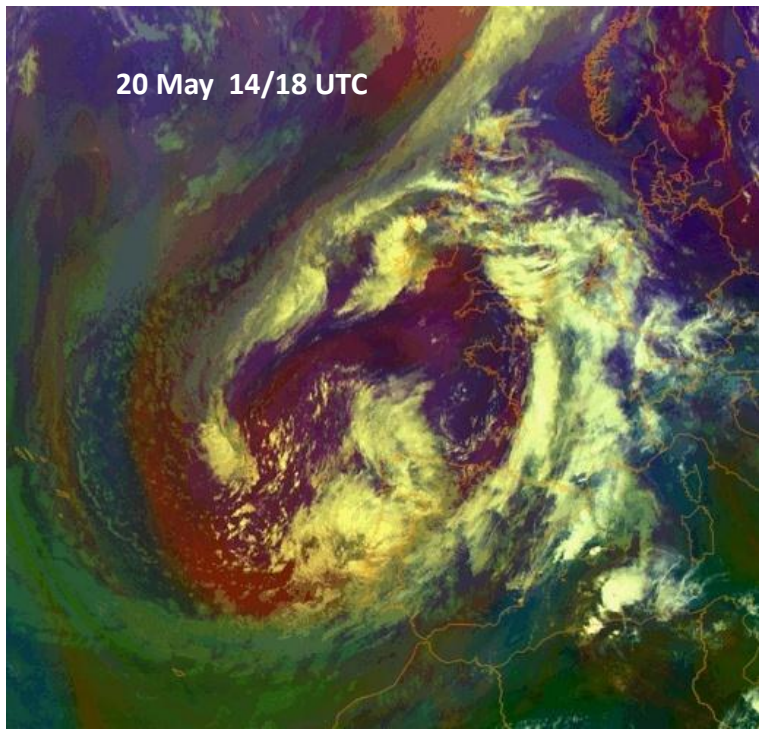
Cold Air Development (CAD)



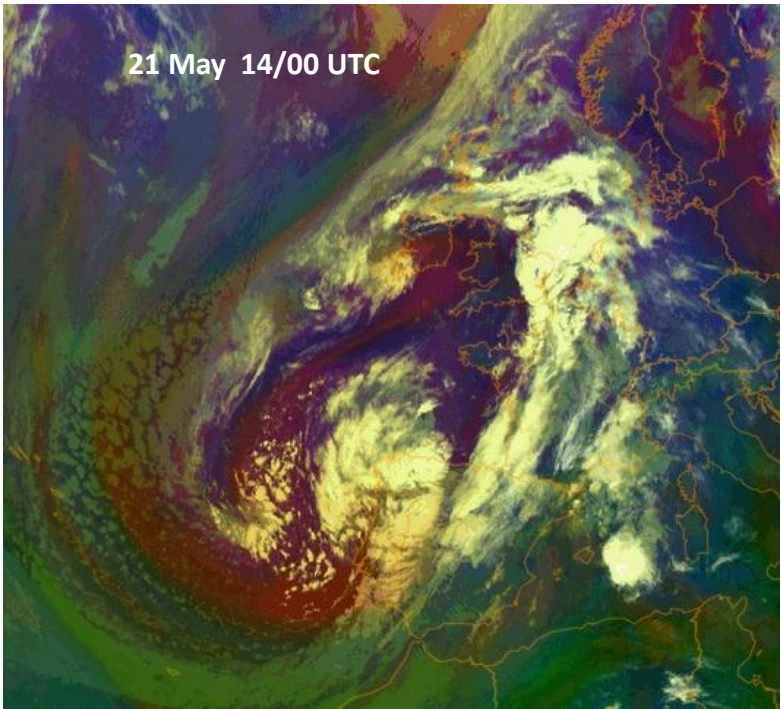


20 – 21
May 2014

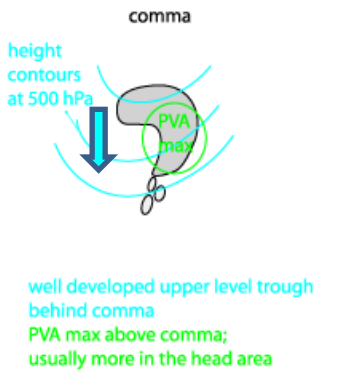
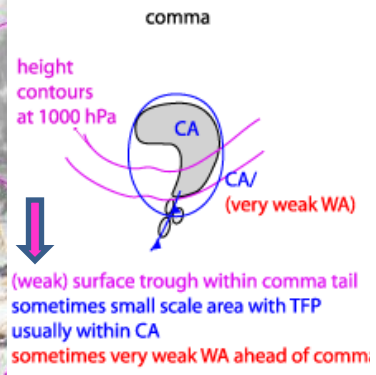
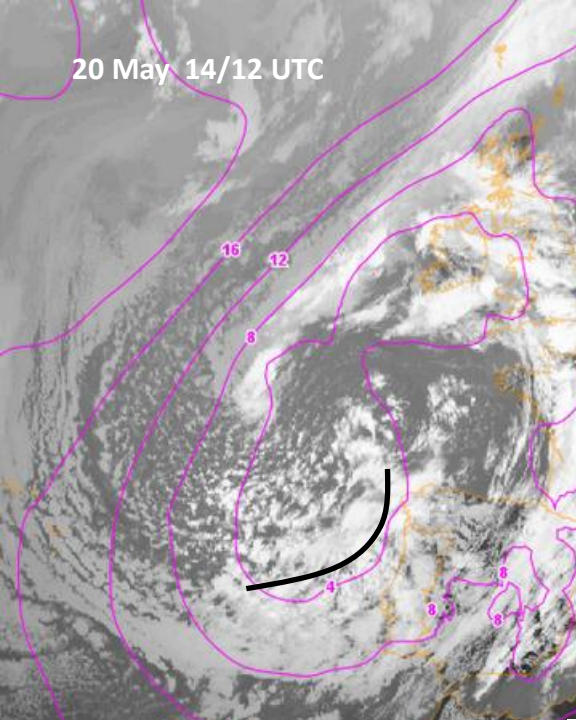
CAD
Cold Air
Development



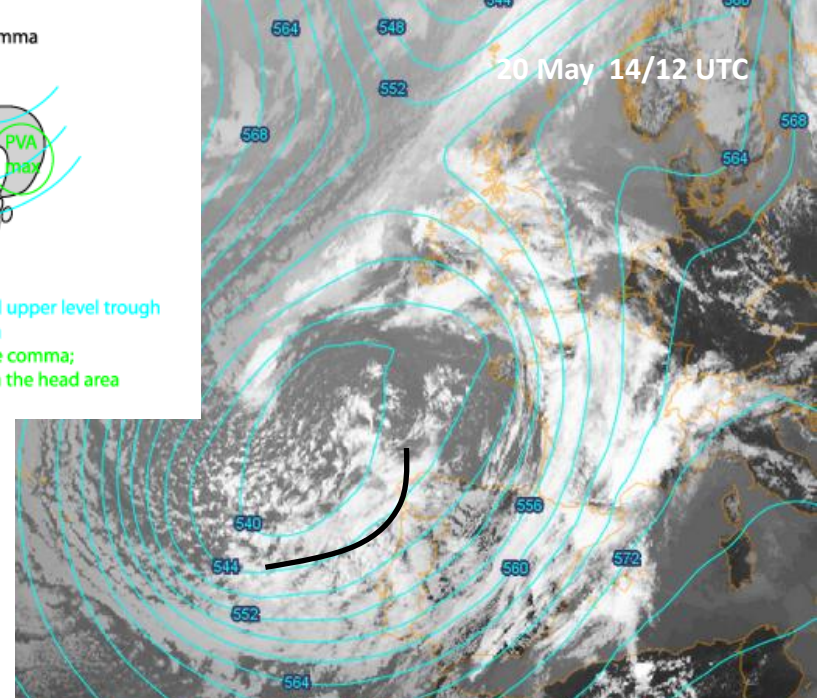
Air mass RGB



20 May 14/12 UTC

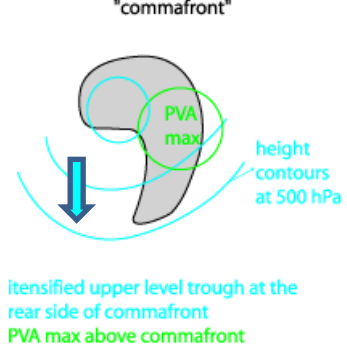
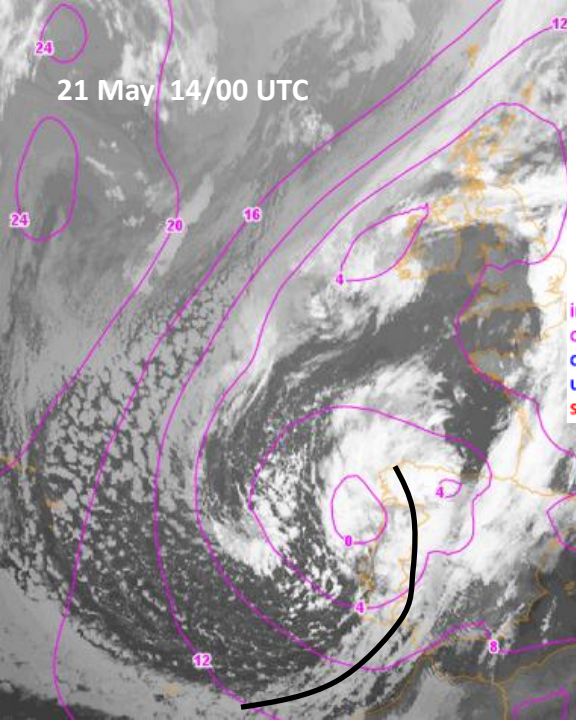


20 May 14/12 UTC

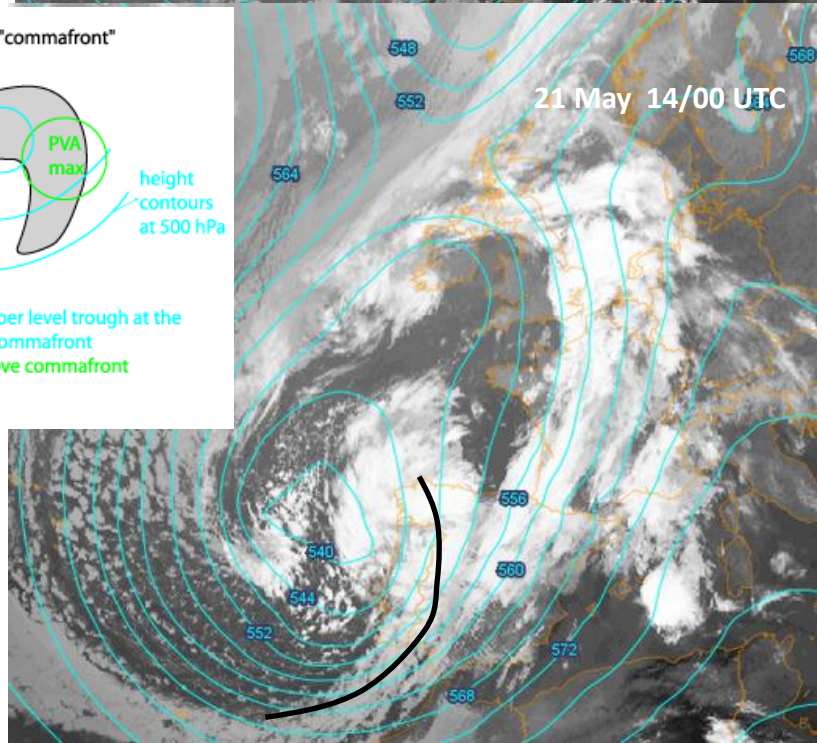


20 – 21
May 2014
CAD
Cold Air
Development

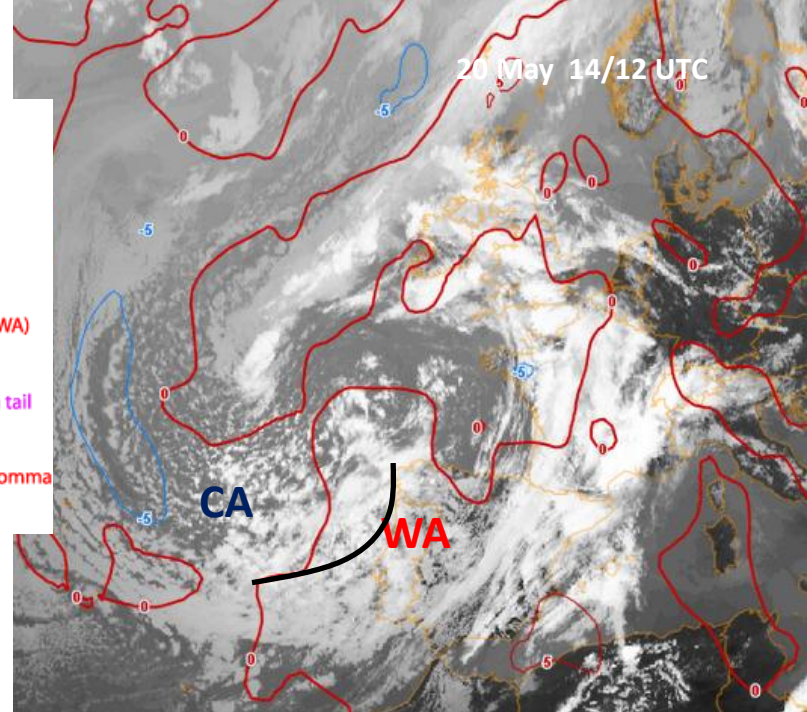
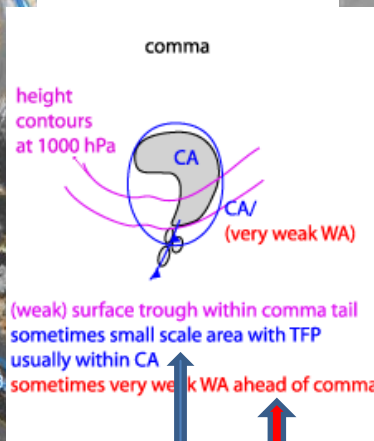
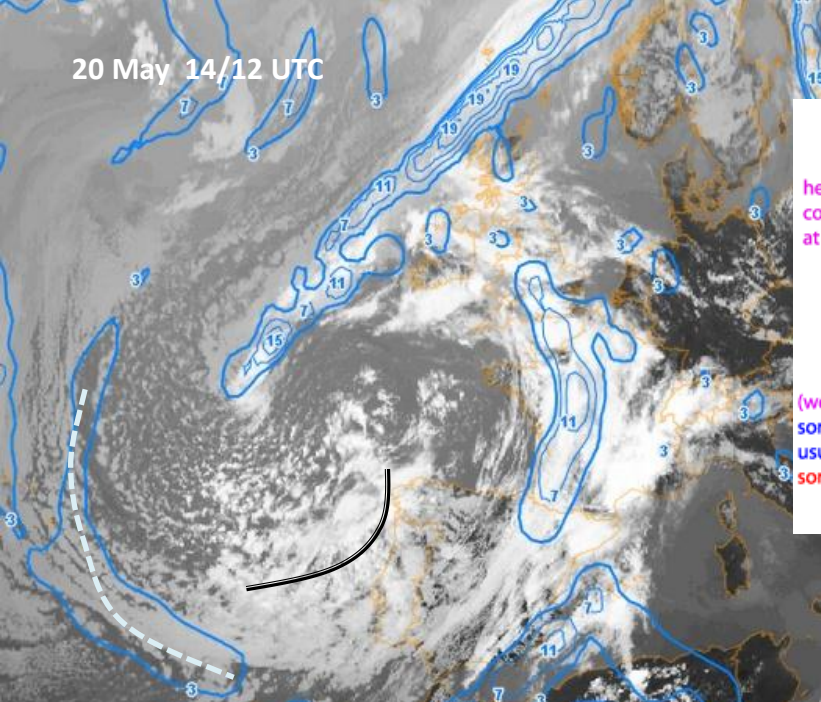
21 May 14/00 UTC



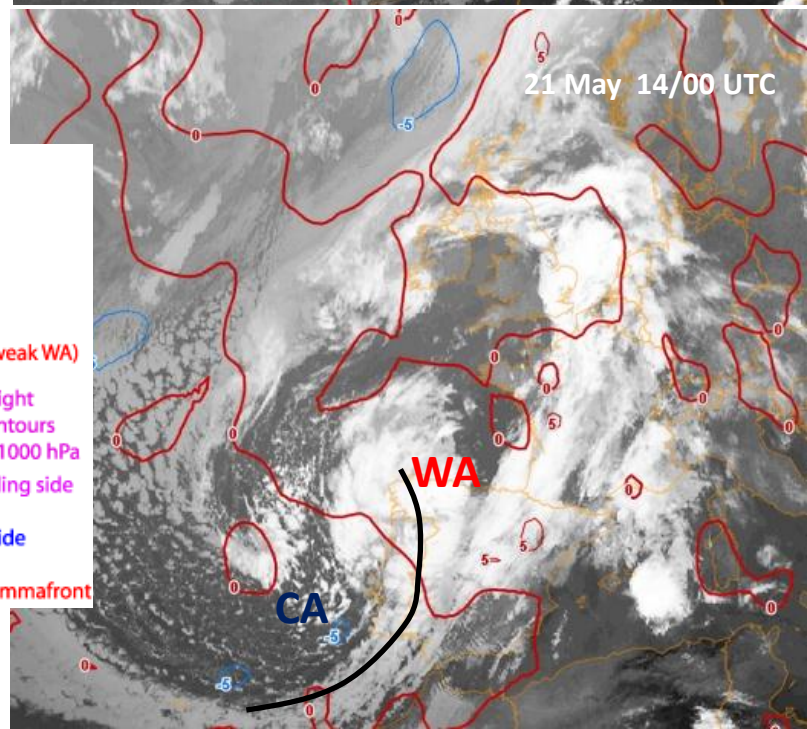
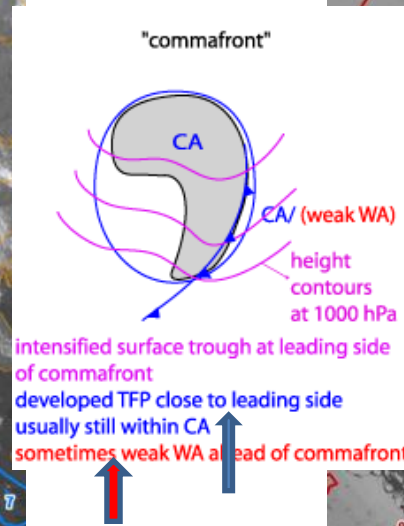
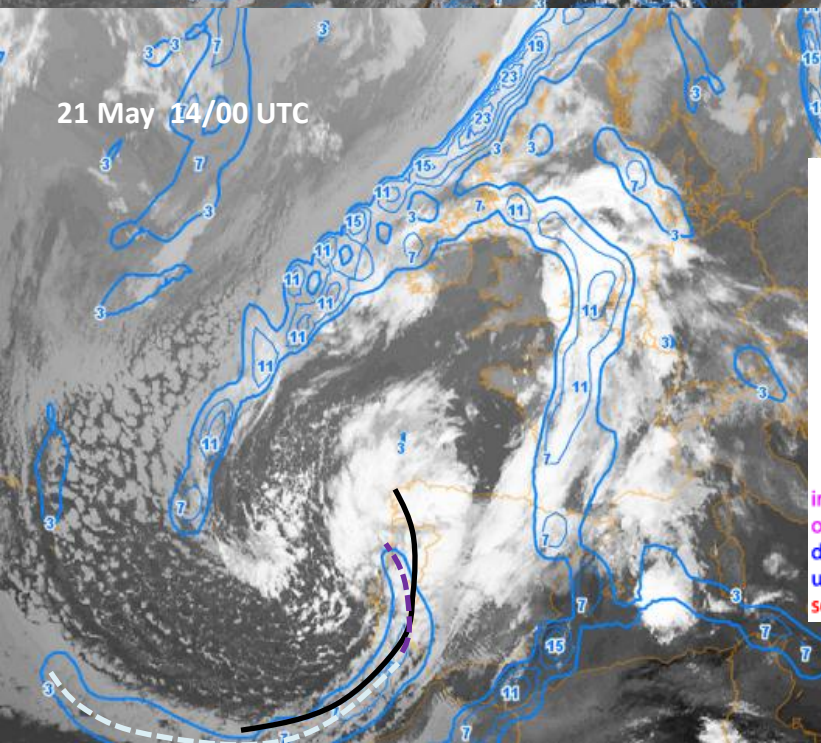
21 May 14/00 UTC



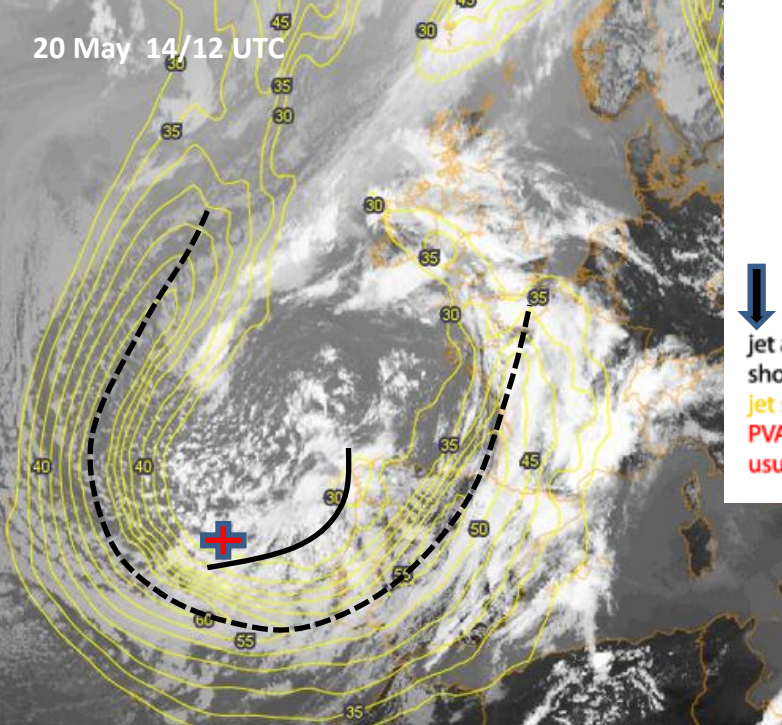
Cyan: height
500 hPa
Magenta:
height 1000 hPa



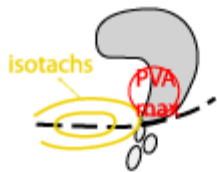
Left: TFP
Right:
temperature advection



20 May 14/12 UTC



comma



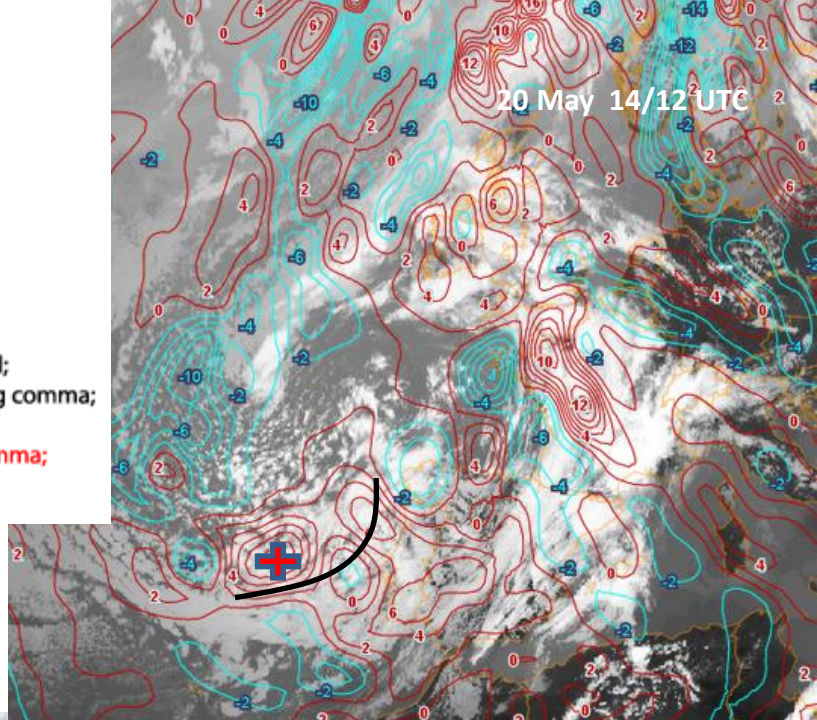
jet axis close to the comma tail;
shows tendency for overtaking comma;
jet streak approaches comma
PVA max at 300 hPa above comma;
usually more over the tail



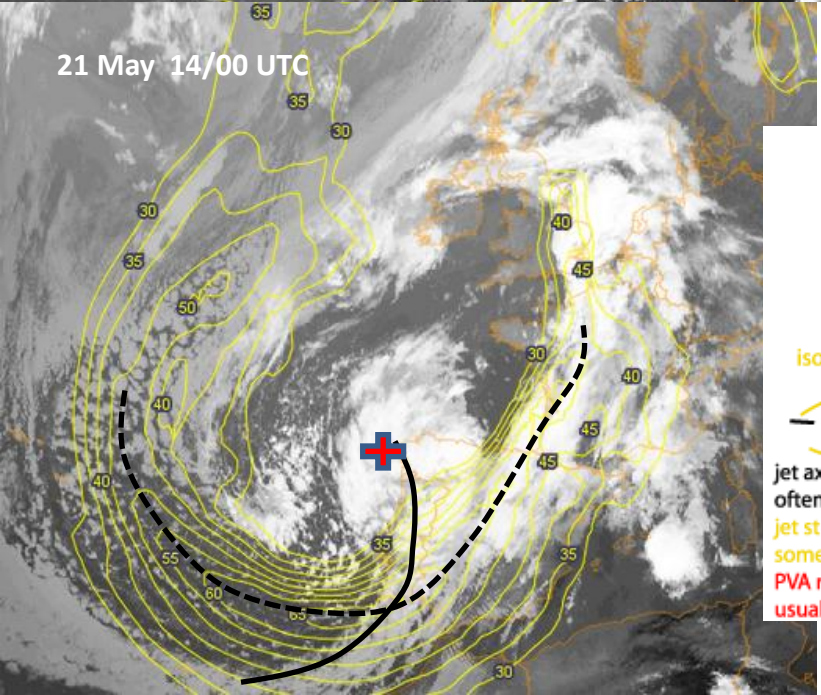
Left: isotachs
300

Right: vorticity
advection 300

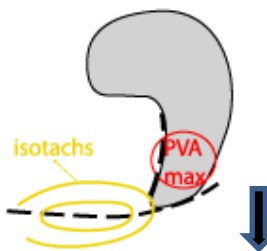
20 May 14/12 UTC



21 May 14/00 UTC



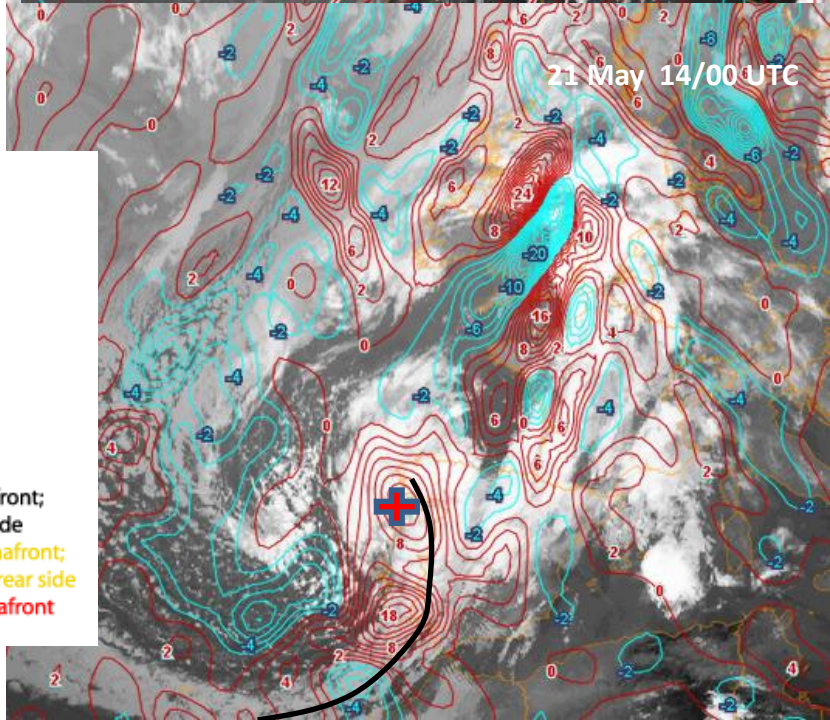
"commafront"



jet axis usually overtakes commafront;
often second branch along rear side
jet streak usually overtakes commafront;
sometimes second branch along rear side
PVA max at 300 hPa above commafront
usually in the southern part

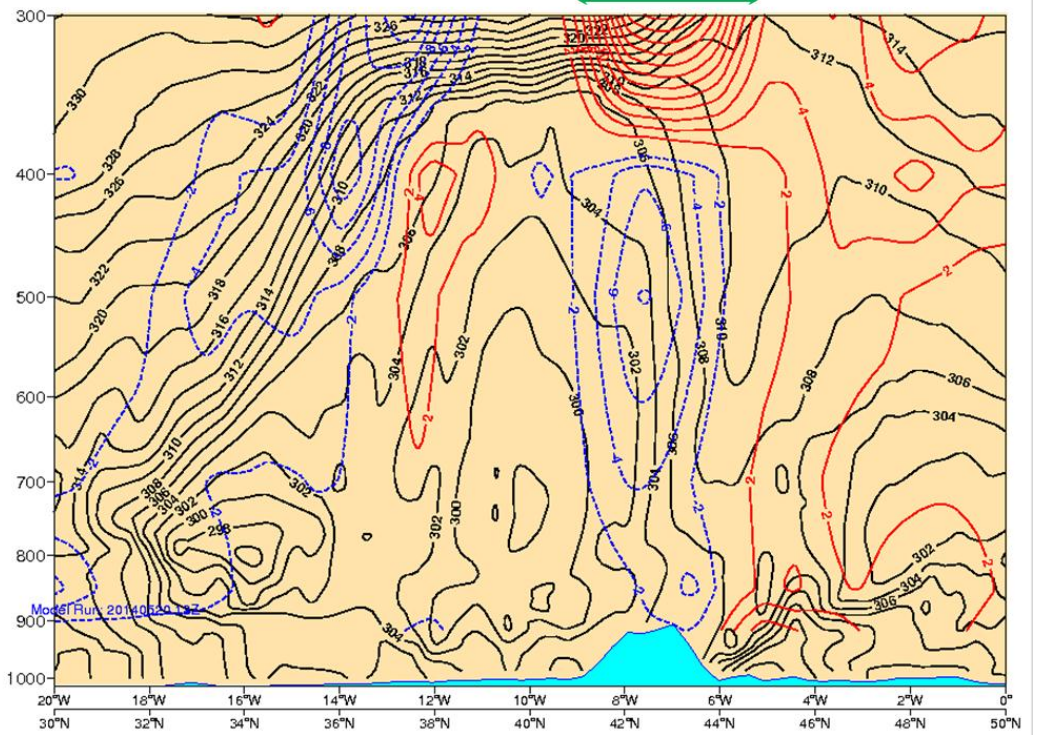
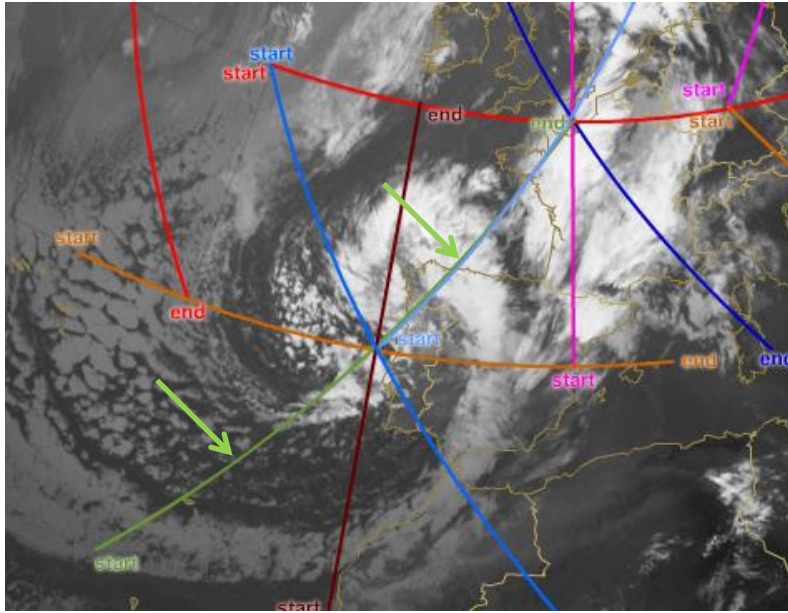
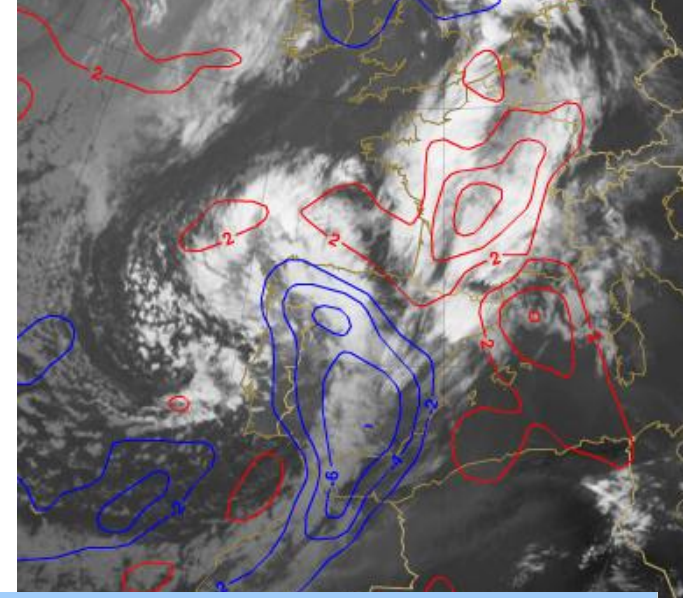
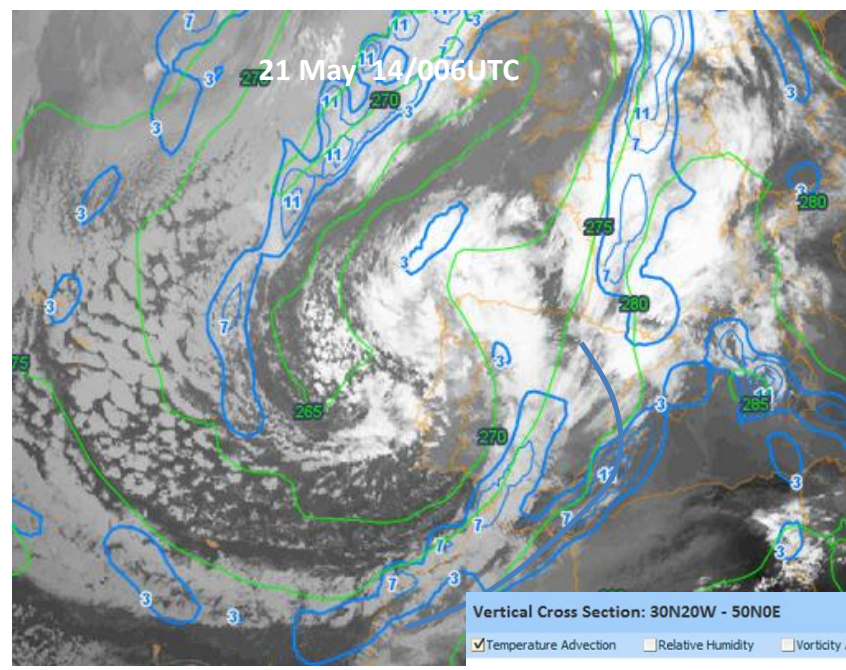


21 May 14/00 UTC



Fully developed stage of a CAD:

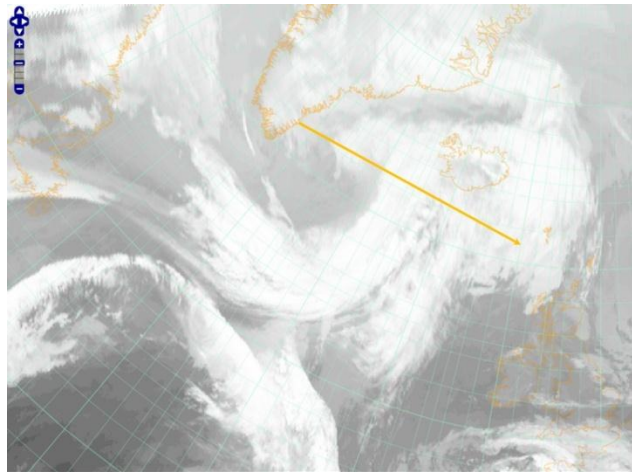
Lack of WF



Enough for today!

Homework:

- Go to EUMeTrain e-port
- Look into the case of 1 July 2014/ 06
 - The image of 1 July 2014/06 UTC is supposed to be a development stage of a rapid cyclogenesis. Prove if this is true or not and if yes describe the material (satellite images, numerical parameters and VCS) which supports your findings.
 - You can use for this the template
- In the homework ppt in the course page you find additional VCSs and hints



- Don't hesitate to start a discussion about new, interesting or difficult topics with me

Enough for today!

**Any Questions already now?
The Students Forum is open for all
questions coming up later!**

**Thank you for your attention!
Meet you in the Students Forum and
in Langen!**