

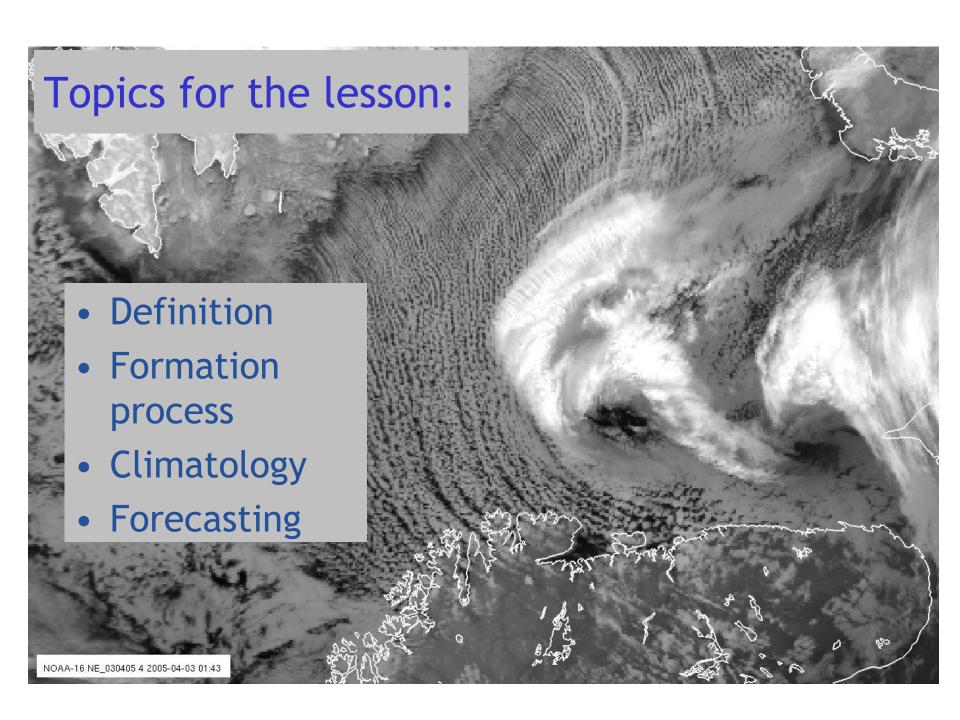
An introduction to polar lows

Gunnar Noer

The Norwegian Meteorological Institute

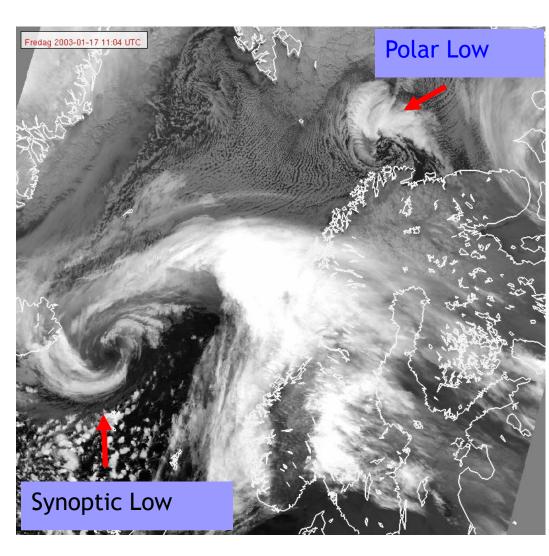
Forecasting division of Northern Norway

12.12.2013



The definition of polar lows by the European polar low working group:

- 'A small, but fairly intense low in maritime regions'
- In cold air outbreaks, well north of the polar front.
- Diameter 200 600km
- Cyclonic curvature

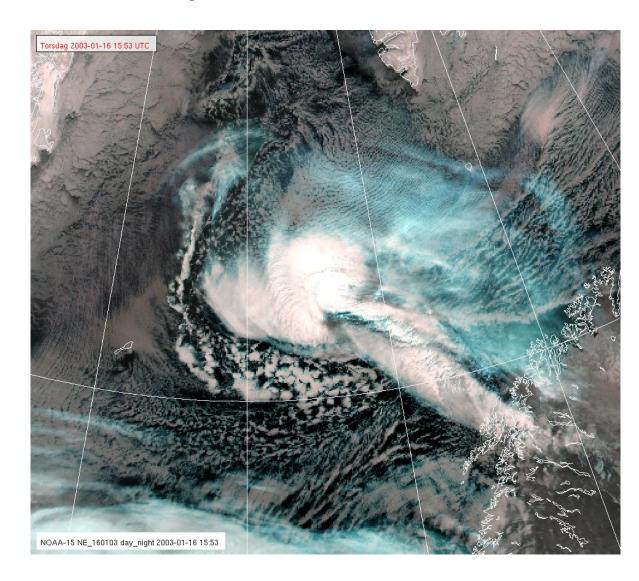


Key processes:

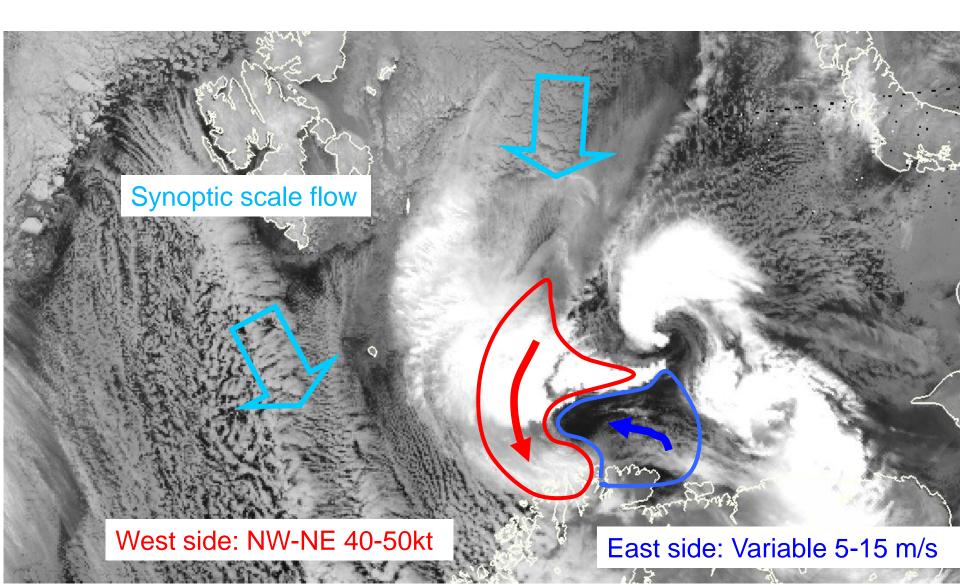
- Polar Lows develop from areas of baroclinic or convective instability:
 - Occlusions, troughs, convergence lines, etc.
- Destabilisation at low levels:
 - Cold arctic air is advected out over warmer sea surface.
 - Transfer of sensible and latent heat from the sea surface
- Further destabilisation of the layers above:
 - Passing cold core, unobstructed static instability up til 500-400hPa
 - Trough at mid troposphere, e.g. as seen from Z500 hPa

Characteristics of the polar low

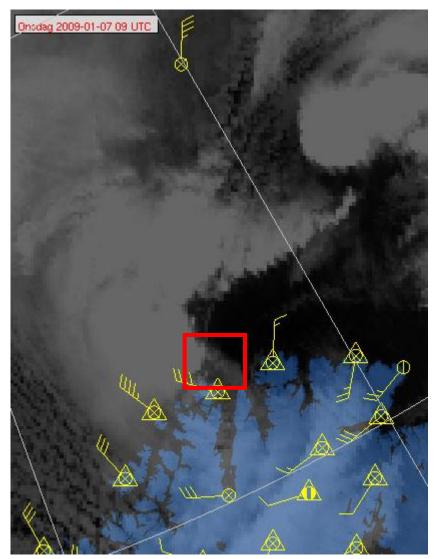
- Average max obs. windspeed 42kt
- About 30% have more than 50kt
- Rapid changes in wind speed and direction
- Showers of snow or hail, visibility less than 100m
- Avalanche danger
- Waves



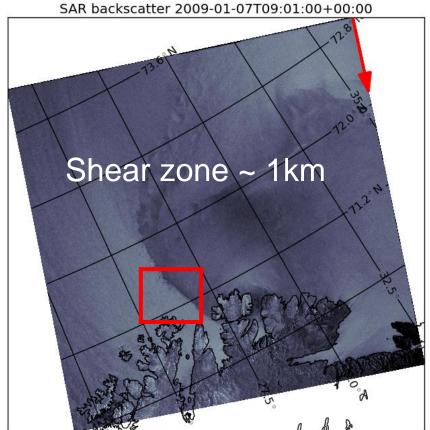
Typical wind pattern:



The wind from SAR-images:



Honningsvåg 7th January, 2009

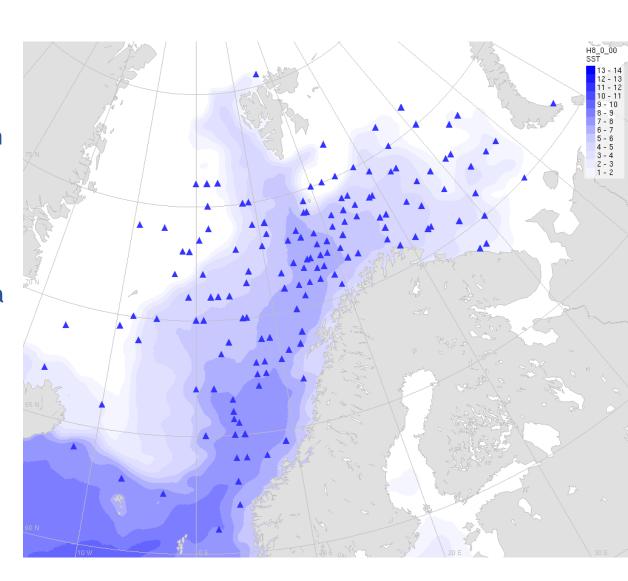


Kilde: Furevik et. al. 2012

1.0 -0.5 -0.0 -0.5 -1.0 [gp] (embis) bolo -1.5 bolo -2.0 bolo -2.5 -3.0 -3.5

Geographical extent

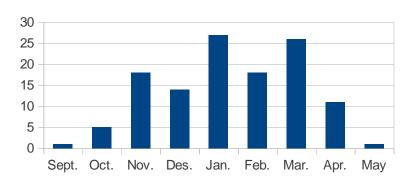
- Points of development
- Temperature gradients in the Golf current
- Areas of convergence, Spitsbergen
- Less cases in the eastern and southern part of area
- 190 events 2000-2013



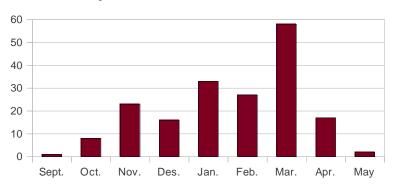
Seasonal variation:

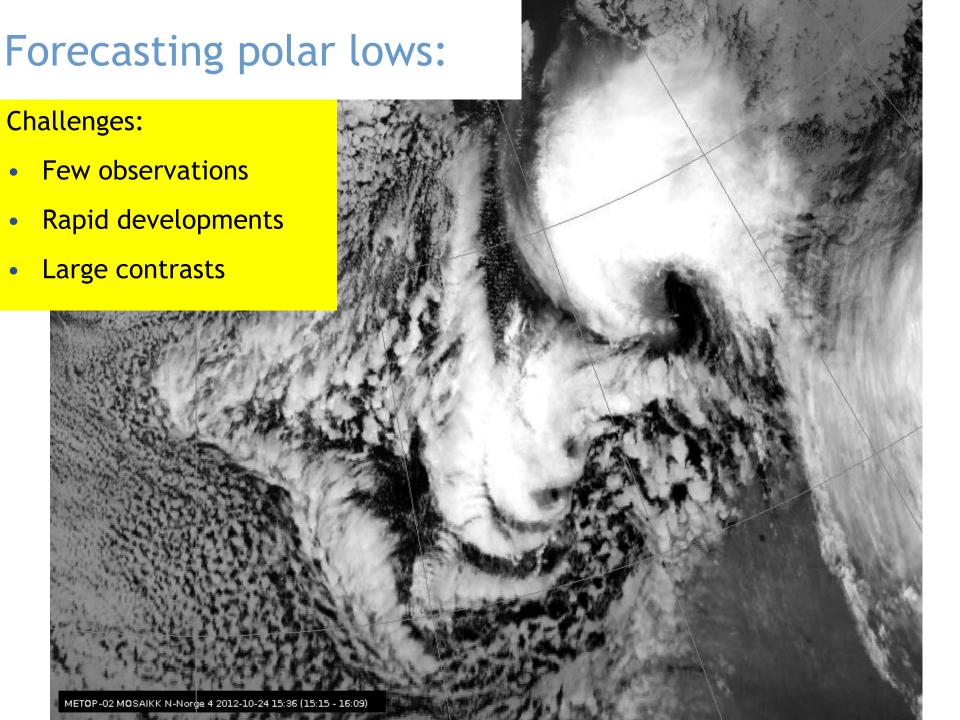
- On average 13 events pr. Year
- October til May
- Max in January and March, Local nadir in February
- Interannual variation in November and March
 - 2000-2009: On average 2,8 pr month in March
 - In 2010, 2011 and 2013: 8 to 12 PL's in March

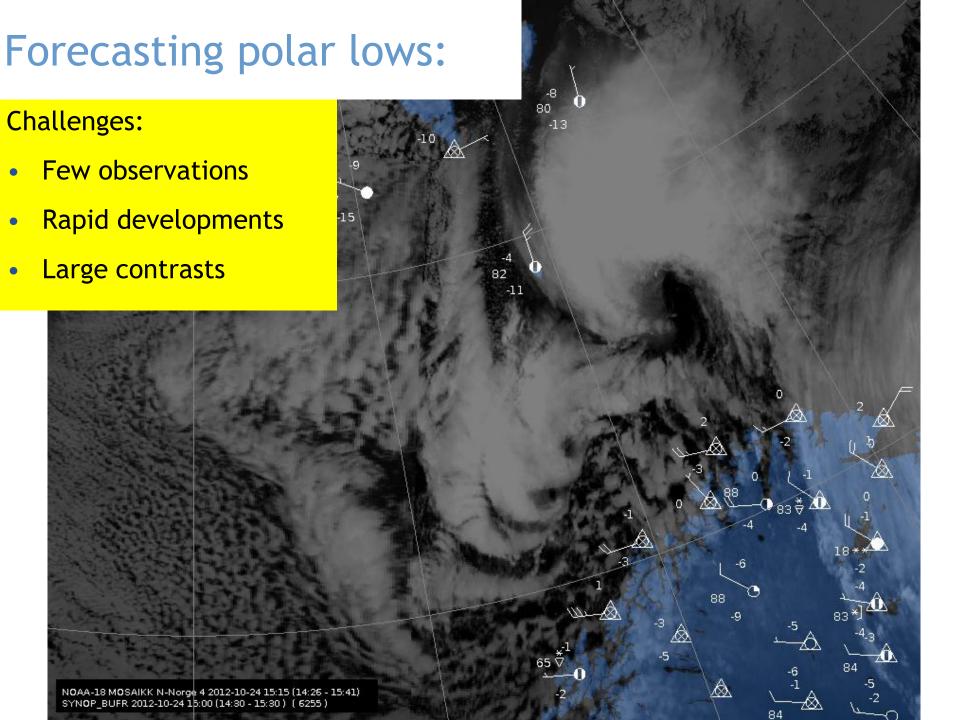
Monthly variation of Polar Lows 2000-2009

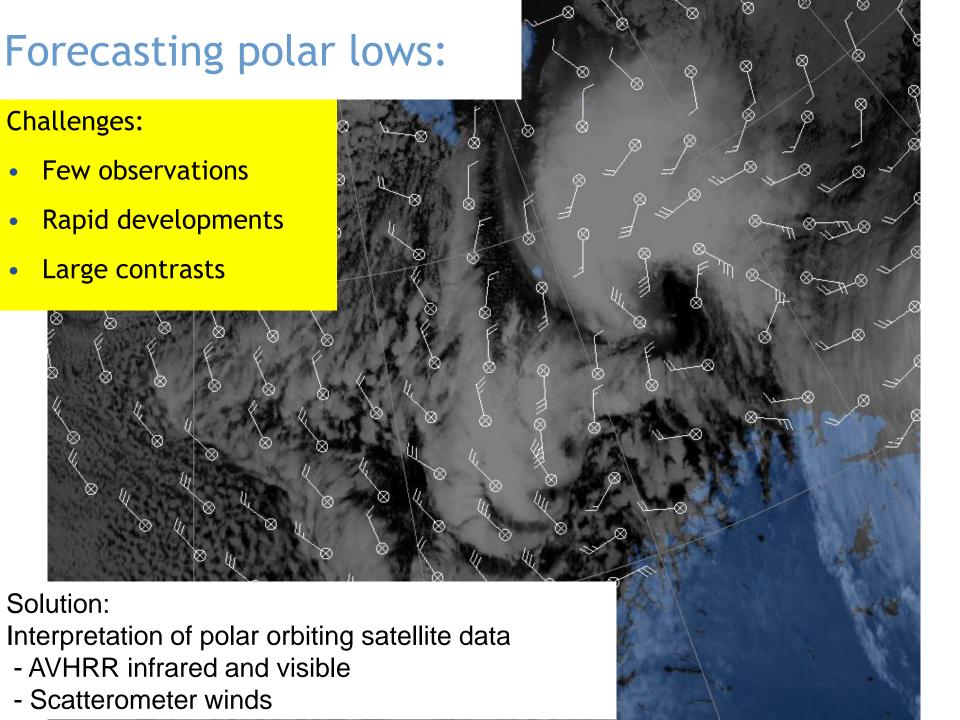


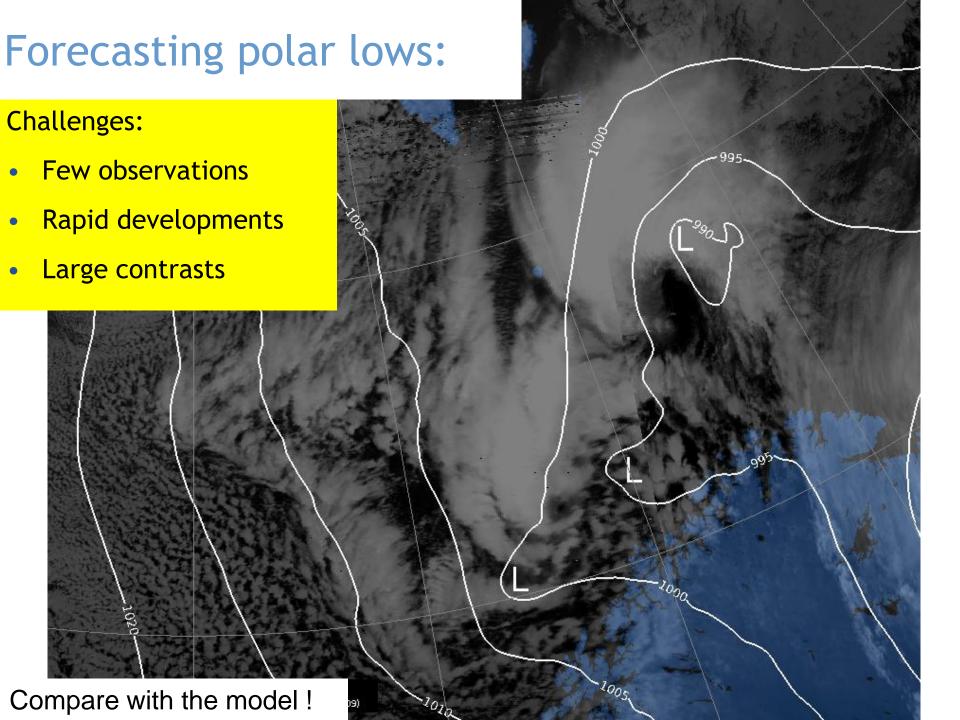
Monthly variation of Polar Lows 2000-2013

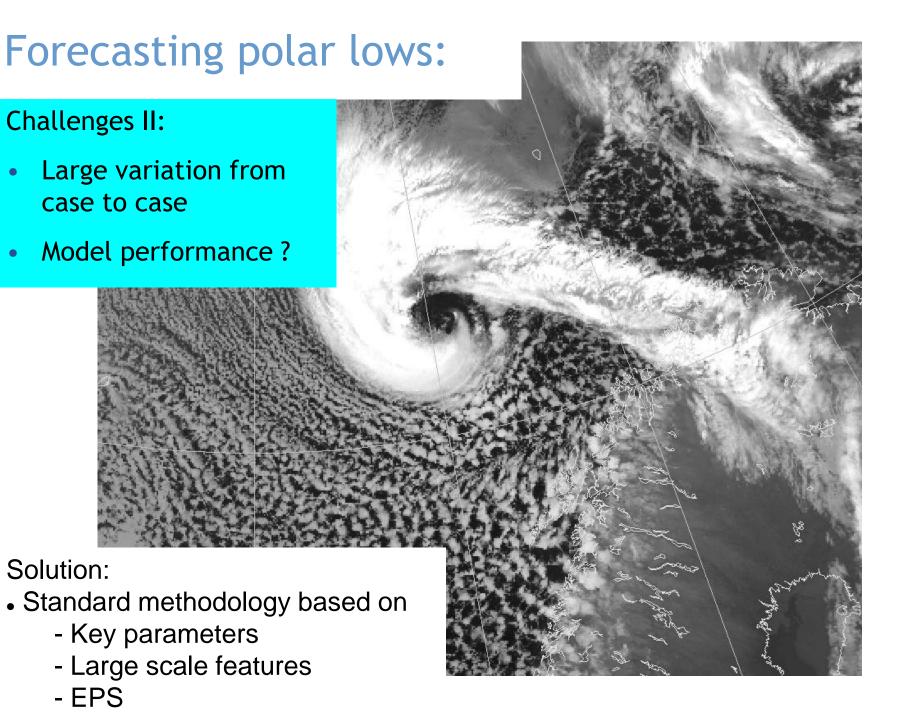










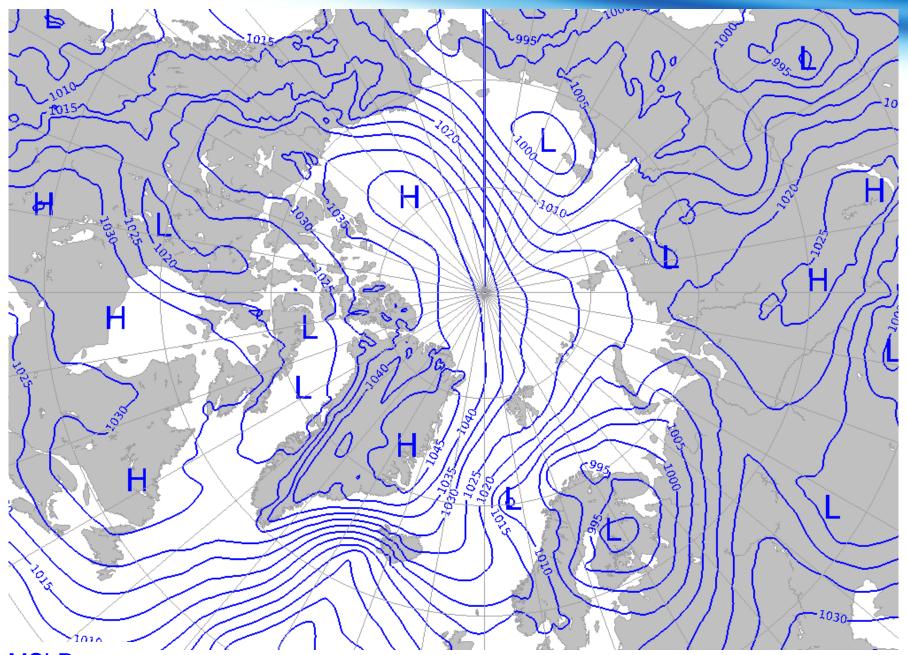


Standard methodology:

- Synoptic scale:
- Cold air outbreak at low levels
- Trough in the mid. Trophosphere, e.g. Z@500hPa
- The temperature potential; SST-T500

- Meso scale:
- -Low level source of instability:
- -Troughs, convergence lines, fronts, etc...

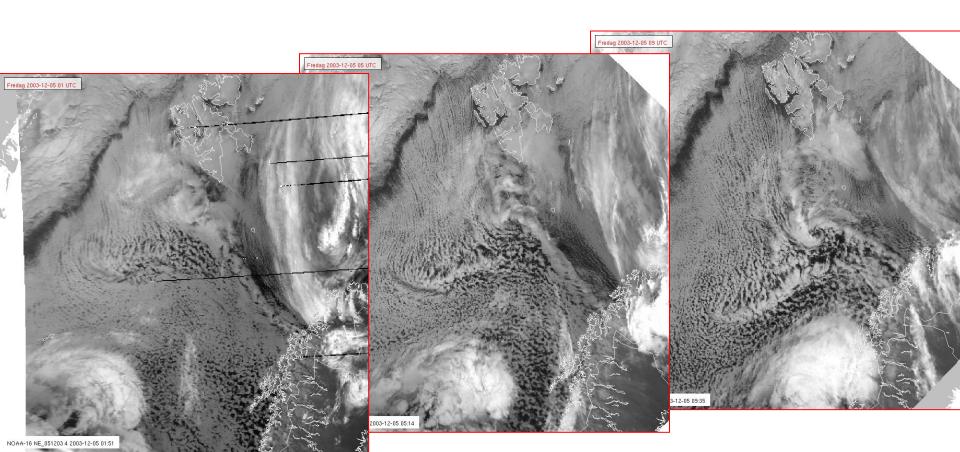
The cold air outbreak

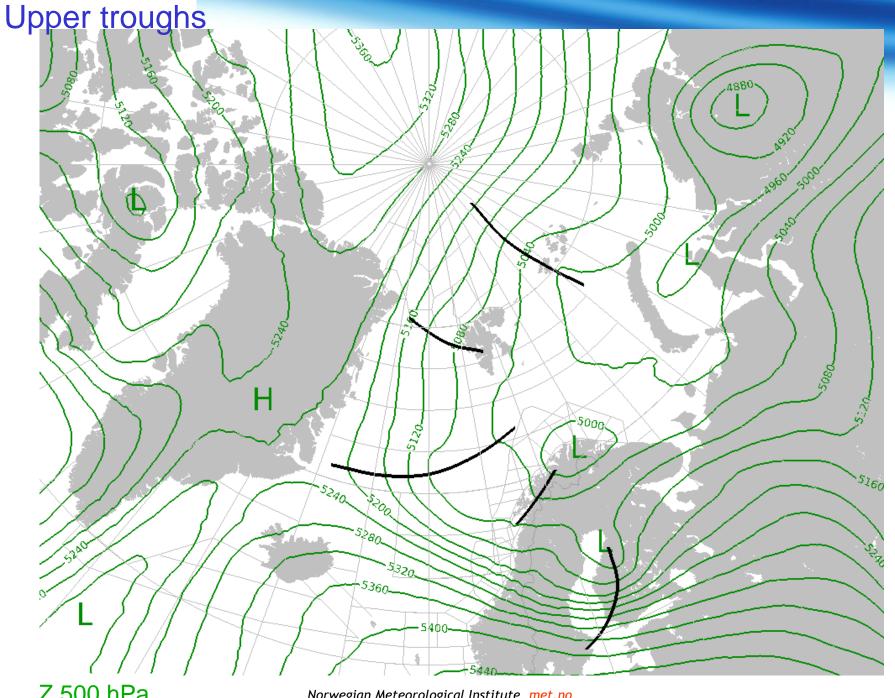


The cold air outbreak STARC.10_070126_00, MSLP (+6) 2007-01-26 06 UT/ NOAA-15 ATL_070126 4 2007-01-26 06:04 **MSLP** Potential temperature Relative humidity

Low level areas of instability

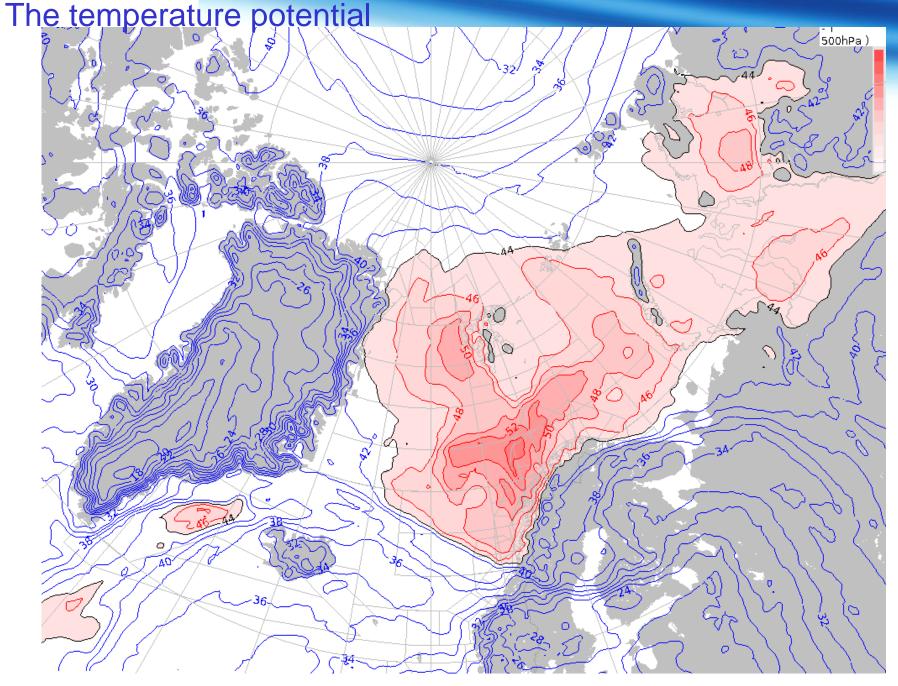
Convergence lines, arctic fronts, troughs, CB-clusters, old fronts, etc.





Z 500 hPa

Norwegian Meteorological Institute met.no

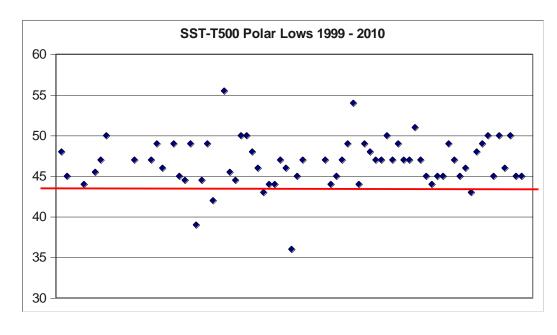


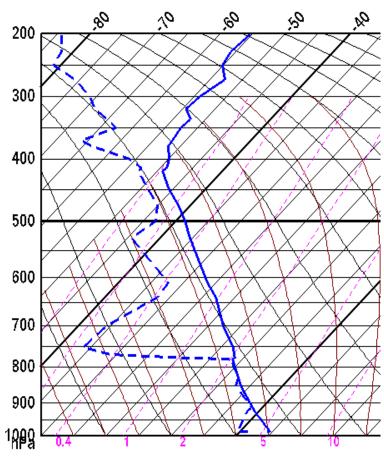
The temperature potential

How big a difference?

$$SST - T_{500} \ge 44^{\circ}C$$

- with exeptions!

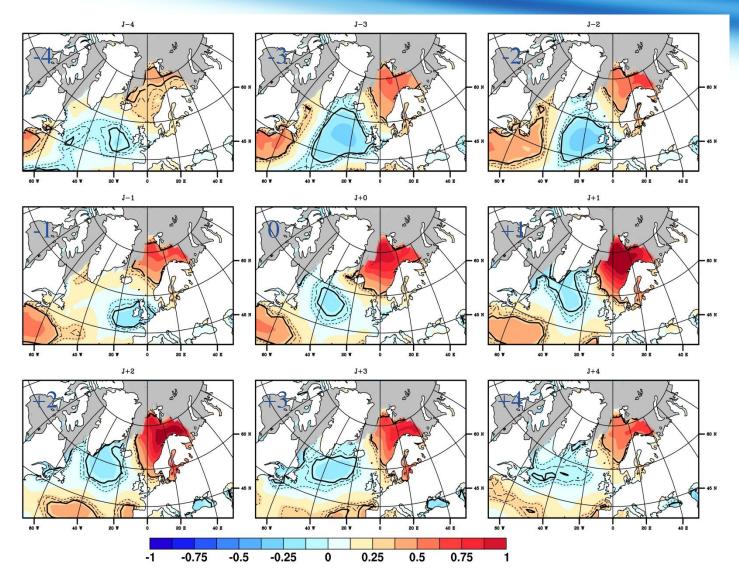




Time dependency of the temperature potential:

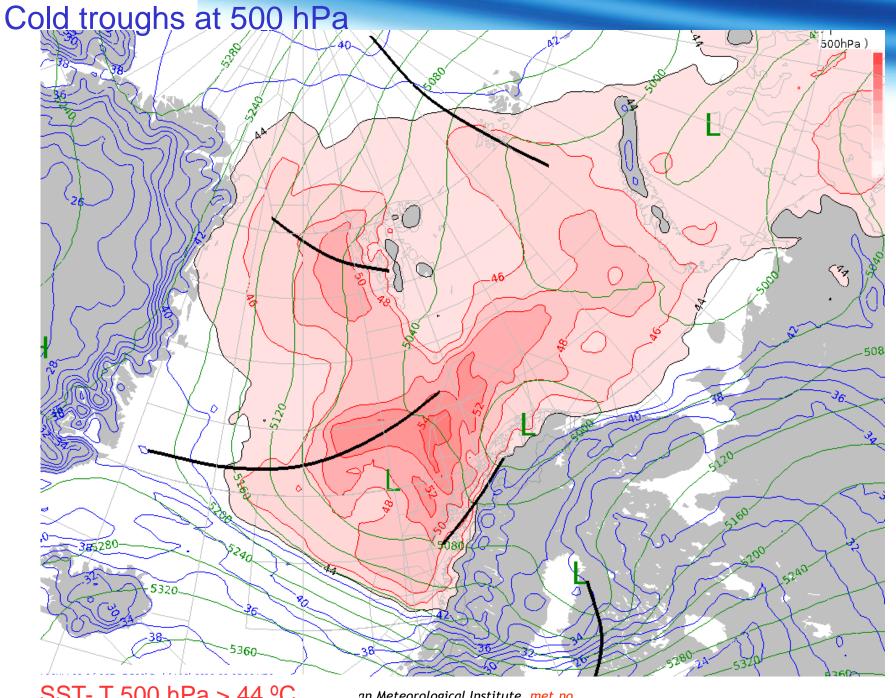
Standard deviation of the temperature potential

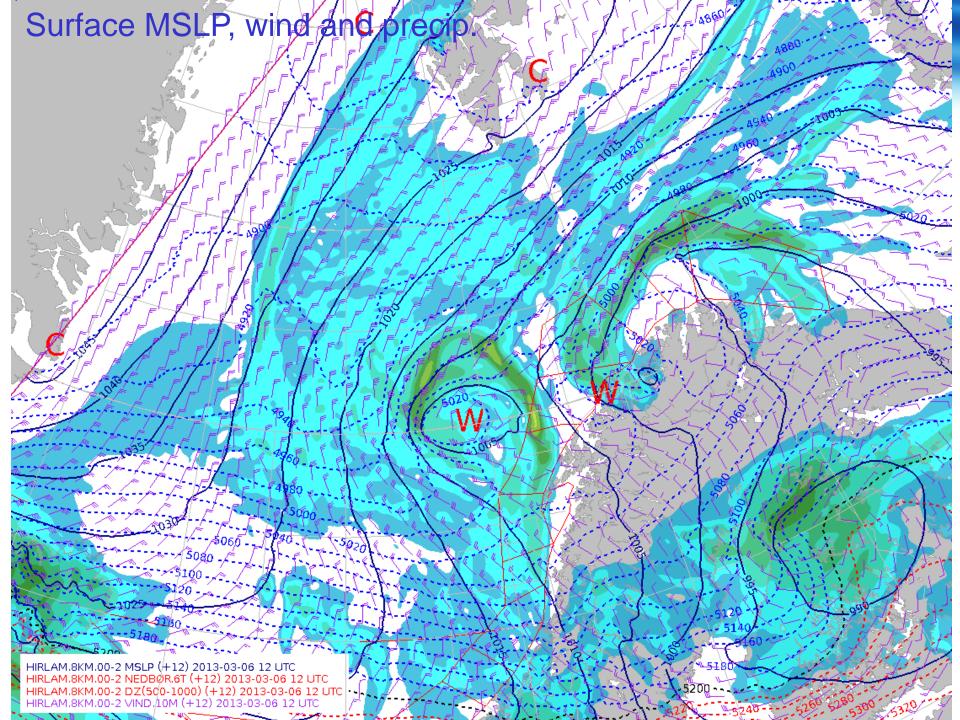
J+0 is day of development

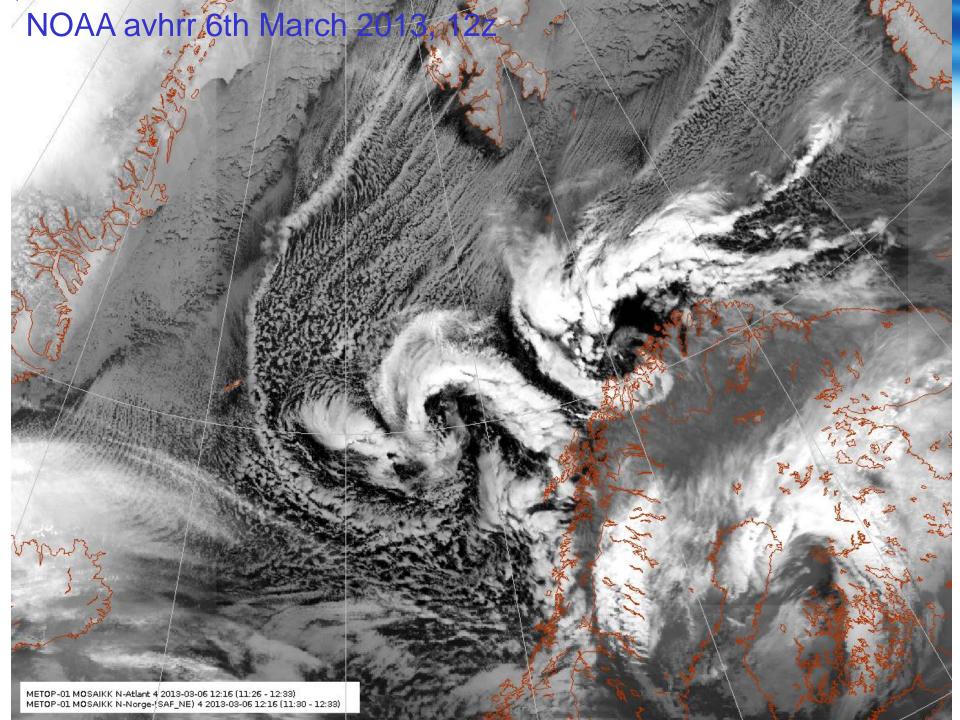


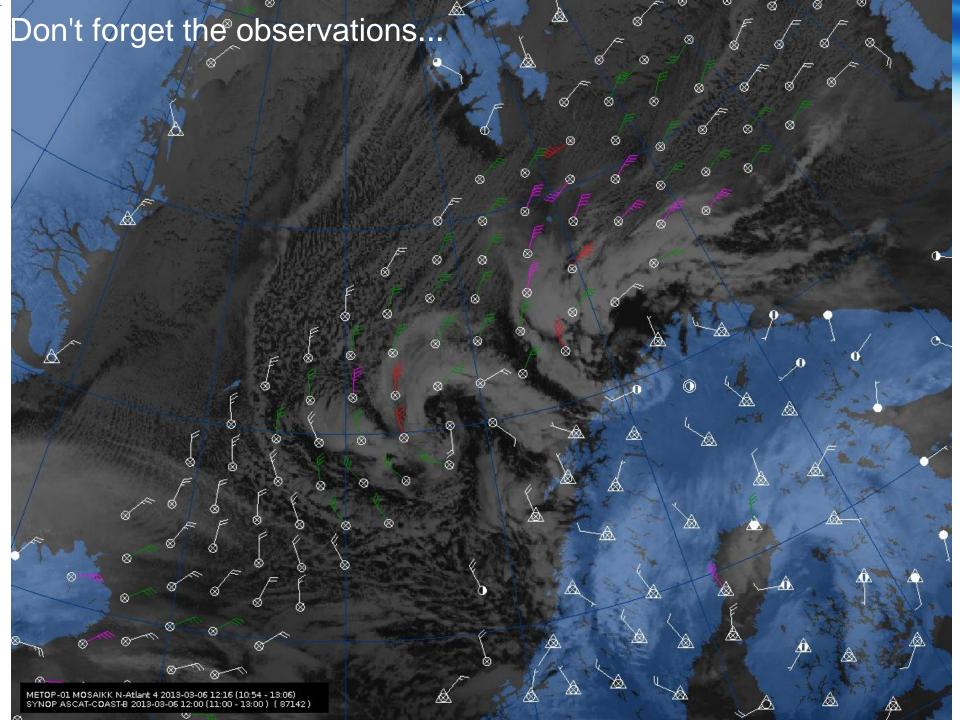
Polar lows over the Nordic and Labrador Seas: Synoptic circulation patterns and associations with North Atlantic-Europe wintertime weather regimes.

Mallet, Claud, Cassou, Noer, and Kodera, 2012





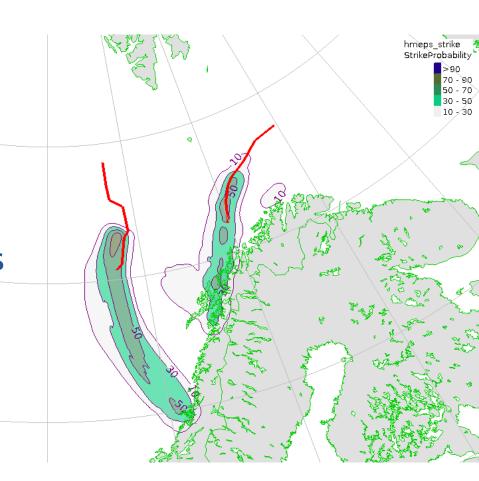




Practical use of the temperature potential:

Polar low tracking:

- EPS, downscaling EC-EPS by Arome 2,5km
- Tracking of low centers
- Filtering by
 - Temperature potential
 - Diameter
 - Duration



Why worry?



ENTC TROMSO/LANGNES -- NO

Raw surface observations for station ENTC

```
ENTC 061950Z 05009KT CAVOK M05/M08 Q1000 NOSIG RMK WIND 2600FT 03012KT=
ENTC 061920Z 06008KT 9999 FEW010 SCT065 M05/M08 Q0999 NOSIG RMK WIND 2600FT 04020KT=
ENTC 061850Z 05015KT 9000 -SHSN FEW015 SCT045 M04/M07 Q0999 NOSIG RMK WIND 2600FT 03021KT=
ENTC 061820Z 36016G26KT 9999 VCSH SCT011 BKN045 M04/M06 Q0998 NOSIG RMK WIND 2600FT 02026KT=
ENTC 061750Z 36018KT 2000 SN VV012 M04/M05 Q0998 BECMG 9999 FEW010 SCT030 RMK WIND 2600FT 36039KT=
ENTC 061720Z 31018KT 2200 -SHSN VV017 M03/M06 Q0998 RESHSN TEMPO VRB08KT 0500 +SHSNGS VV004 RMK WIND 2600FT 31049KT=
ENTC 061650Z 33019KT 1500 R01/1200VP2000N R19/P2000N SHSN VV009 M04/M06 Q0999 TEMPO VRB08KT 0500 +SHSNGS VV004 RMK WIND 2600FT 33040KT=
ENTC 061620Z 31026KT 0800 R01/1000VP2000N R19/P2000D SHSN VV003 M04/M06 Q0999 TEMPO VRB08KT 0500 +SHSNGS RMK WIND 2600FT 33046KT=
ENTC 061520Z 30021G31KT 1000 R01/P2000D R19/1000VP2000D SHSN VV003 M04/M07 Q0999 TEMPO VRB08KT 0500 +SHSNGS RMK WIND 2600FT 32046G56KT=
ENTC 061450Z 31020KT 2500 -SHSN VV016 M04/M06 Q0999 TEMPO 0500 +SHSNGS VV004 RMK WIND 2600FT 33051KT=
ENTC 061420Z 33021KT 9999 -SHSNGR FEW014 SCT026 BKN037 M04/M07 Q0999 TEMPO 0500 SHSN VV004 RMK WIND 2600FT 34053G69KT=
ENTC 061350Z 22021KT 2000 -SHSNGR VV010 M03/M07 Q0999 TEMPO 0500 SHSN VV004 RMK WIND 2600FT 28003KT=
ENTC 061350Z 28012G24KT 9999 VCSH FEW006 SCT021 BKN041 M04/M10 Q1000 TEMPO 0500 SHSN VV004 RMK WIND 2600FT 28003KT=
ENTC 061250Z 28012G24KT 9999 VCSH FEW006 SCT021 BKN041 M04/M10 Q1000 TEMPO 0500 SHSN VV004 RMK WIND 2600FT 28003KT=
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Tromsø airport mountain observation at 2600ft:

13:20z: 290 03kt (Westerly force 1)

ENTC 061220Z 27012KT 9000 -SHSN FEW008 SCT020 BKN040 M05/M10 01001 TEMPO 2000 SHSN VV010 RMK WIND 2000FT 32005KT=

13:50z: 340 53kt gusting 69kt (Northwesterly storm force 10)





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